

# Status of Assimilating Satellite Data at Deutscher Wetterdienst (DWD)



R. Hess, C. Köpken, C. Gebhardt, W. Gräsele

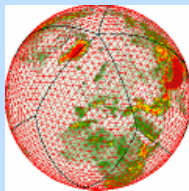
Deutscher Wetterdienst, Germany

E-Mail: reinhold.hess@dwd.de

## Model System: GME 40/40

### GME (Global Model):

hydrostatic, icosahedral-hexagonal grid, mesh size 40km  
terrain following hyb. coordinate,  
40 layers, bottom layer at 10m, top 10 hPa,  
7 soil layers, freezing and melting  
of soil water  
forecast range: 174h initial dates 00, 12 UTC  
and 48h for 18 UTC  
prognostic cloud ice, prognostic sea ice



### Analysis:

OI optimal interpolation,  
3-hourly intermittent analysis,  
observation window: +/- 1.5h,  
observations: conv., AIREP, AMDAR, ACARS,  
BUFR AMV (NOAA 15, 16),  
SATOBS (GOES E and W, GMS, Meteosat 5 and 7), PAOB  
MODIS-Winds, pseudo-temps  
cutoff: 2h30 (for forecasts)

## Monitoring of ATOVS (AMSU-A)

Monitoring with routine GME available on internet in real time:

<https://www.dwd.de/wm/Products/amsu/amsu.html>

Observed minus background  
(o-b) statistics for ATOVS and  
different areas.

Diurnal cycles in statistics  
because of different model  
climates of GME and IFS  
(using pseudo-temps).

Pseudo-temps cool down GME  
at 0 UTC leading to negative  
o-b differences. During the next  
24 model hours the  
analysis-forecasting system  
warms up again to reach for  
its own climatic equilibrium.

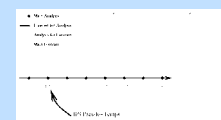


## Use of ECMWF-Pseudo-Temps

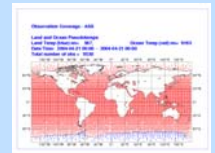
(operational since Dec 17 2003)

Idea: use profiles of IFS/4D-Var analyses (temperature, humidity and wind) of ECMWF  
and assimilate them as temps in OI/GME.

In this way satellite data that is assimilated at ECMWF is used for GME giving  
a significant boost in forecast quality especially in the southern hemisphere.



Assimilation is done only one time per day  
for 0 UTC into update (main) analysis of GME.  
In this way an almost independent  
analysis-forecast system is still maintained.  
Pseudo-temps are to be substituted with  
the assimilation of radiances (ATOVS/IASI)  
with 1D-Var or directly with 3D-Var  
once the forecast skill of pseudo-temps is  
reached.



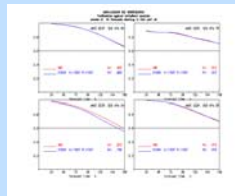
Data coverage of pseudo temps:  
about 100 km over sea and Antarctica.

Humidity is assimilated only above  
700 hPa to not affect the humidity  
equilibrium of GME in the lower  
troposphere, which is different to  
the IFS climate.

## Development: 1D-Var for AMSU-A (NOAA 15, 16 and Aqua)

### Setup:

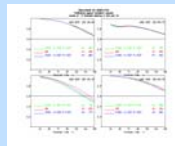
- Receive level 1c ATOVS from UKMO (NOAA 15, 16 and Aqua)
- Radiative transfer computation: RTTOV-7
- Use IFS/ECMWF forecasts for radiative transfer in stratosphere
- Channel choice: AMSU-A 5-14 (HIRS deferred)
- Assimilation of temperature only
- Bias-Correction: UKMO-Style (scan angle and mass, A4 and A9)
- MW Rain/ice-detection (Kelly and Bauer-ECMWF)
- Assimilation as SATEMS in OI (geopotential thicknesses),  
Assimilate temperature increments to avoid spurious biases
- B-Matrix based on GME-Statistics (NMC-Method)



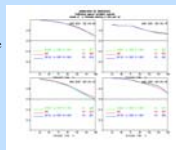
### Status:

Pre-operational experiments:  
In northern hemisphere and Europe almost equal to use of  
pseudo-temps, however still significant lag in southern hemisphere.

Comparison to SATEMS:  
Significant improve-  
ment of 1D-Var in  
southern hemisphere.



Use of AMUA-Aqua  
(not introduced in above  
statistics):  
Small but  
significant  
improvement  
in southern  
hemisphere.



## Tuning of observation and background errors

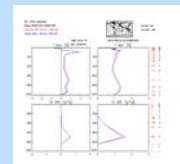
Independent data is required to estimate observation  
and background errors simultaneously

Idea: use pseudo-temps:

- Tune magnitudes of R and B for the correct size  
of the 1D-Var cost function at its minimum
- Tune relation of R and B for best collocation of retrievals  
with pseudo-temps

Reduced observation errors:

Channel	B	R	B	R	B	R	B	R	B
15	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
16	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
17	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
18	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
19	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
20	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
21	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
22	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
23	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
24	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
25	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
26	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
27	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
28	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
29	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
30	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
31	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
32	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
33	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
34	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
35	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
36	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
37	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
38	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
39	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001
40	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001	0.001

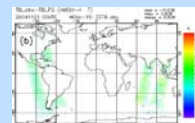


## Tuning of horizontal thinning

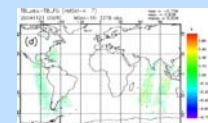
Apply methods from information and image theory: keep most information with fewest fields of view, e.g.:

- estimation method: take that point that can be derived worst from its neighbours
- cluster method: take most representative field of view
- etc.

Compare analyses and subsequent forecasts to define best method; work is in progress.



ATOVS data coverage from regular (naïve) thinning



data coverage resulting from estimation method

## Development of PSAS (3D-Var)

(physical space assimilation system, dual space, observation space)

### Minimisation in Observation Space rather than Model Space

Conventional 3D-VAR (MSAS, Model Space Assimilation System)

solve:

$$\mathbf{c}_o = \mathbf{c}_a + [\mathbf{B}^{-1} + \mathbf{R}^{-1}]^{-1} \mathbf{R}^{-1} \mathbf{y} - \mathbf{R}^{-1} \mathbf{y}_o$$

PSAS (OSAS, Observation Space Assimilation System)

solve:

$$\mathbf{c}_o = \mathbf{c}_a + \mathbf{B} \mathbf{R}^{-1} [\mathbf{R} \mathbf{B}^{-1} + \mathbf{I}]^{-1} \mathbf{y} - \mathbf{R}^{-1} \mathbf{y}_o$$

### Comparison of PSAS with conventional 3D-Var:

- PSAS has more flexibility in definition of B  
flow dependent background errors can be introduced more easily.
- Observation space is smaller than model space  
(however high resolution infrared sounders may change situation)
- When using the flexibility specifying B, the minimisation costs become  
quadratic in number of observations  
(thus 1D-Var retrievals remain interesting even with 3D-Var)

## Adaptive Error Correlations derived from NMC-statistics

(depending e.g. on vorticity, geopotential)

Forecast error correlations depend on the synoptic situation  
(as derived from NMC-method)

Left: Vertical error covariances for good weather (red line) are much  
smaller than for bad weather (blue line).

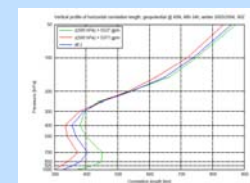
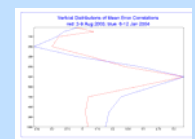
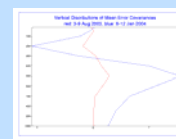
Right: Vertical error correlations for good weather are sharper than  
for bad weather

Good and bad weather is defined by vorticity in this case.

### Strategy:

- Improve analysis with adjusted background error specification
- Improve retrieval of vertical resolution of high spectral resolution sounders

Also horizontal correlation lengths depends on synoptic situation, right:



**Open question:** How much can you improve the analysis/forecast  
with adaptive variances and correlations?