Reg Coef

Recent Development of ATOVS usage in Korea Meteorological Administration

Sang-Won Joo, Eun-Ju Lee, and Seung-On Hwang Korea Meteorological Administration, Seoul, Korea



History of (A)TOVS assimilation at KMA

2000: 1dVar for TOVS + 3dOI

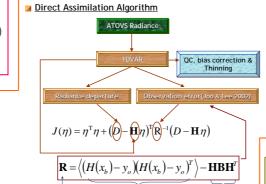
3dVar Formulation

 $\eta = L^{-1}(x - x_b), B = LL^T$

2002: 1dVar for ATOVS + 3dOI

2004: Direct assimilation of ATOVS in 3dVa

 $J(\eta) = \eta^{\mathrm{T}} \eta + (\mathbf{H} \eta - D)^{\mathrm{T}} R^{-1} (\mathbf{H} \eta - D)$



First estimates of Derber and Wu (1999)

 $D = (y - H(x_b))$ • Control variables: $\eta = L^{-1} [\delta \zeta, \delta D_u, \delta (T, P_s)_u, \delta \ln q]_s^{\sigma}$

• Model variables : $x_b = [\zeta, D, T, P_s, q]_s^\sigma$

• Observed variables : $y = [\phi, u, v, T, q]^p$, $P_s, T_s, u_s, v_s, q_s, Rad$

Details of the H operator for radiance

Scan angle and air mass bias correction is done before 3DVAR(Joo and Okamoto, 2000)

Bias correction depends on the cloud condition and latitude No bias correction for the stratospheric channels

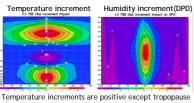
H operator is linearized in the background states and no update within the inner loop

Level 1d data of HIRS 1-8,10-15, AMSU-A 5-15 channels are used at the same time

Reg Coef

Adjoint check and Normalized Cost Function $(Hdx)^{T*}(Hdx)$ dx^TH^THdx sounding 0.025767163 0.025767163 0.514098014 0.514098014 0.744231604 0.744231604 0.864059397 0.864059397 0.758332559 0.758332559 0.557038524 0.557038524 0.146815814 0.146815814

One point observation experiment Put an ATOVS observation of which innovation is 3 deg in all channels at the point of 30N / 125E



and humidity is decreased. It is what we expected!

Background temperature Schematic diagram explains error correlation the negative correlation

Mass-wind relationship Mass and wind balance is maintained by the linear

balance equation and it makes the wind increment by

RTM error and instrument error

The negative temperature analysis increment near the tropopause is caused by the negative background error correlation of temperature.

the ATOVS assimilation in 3dVar.

Effect of fixed TL in basic state

In order to get the 3dVa results in a reasonable time, basic state of the H operator is not updated

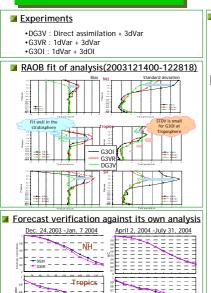
Is it true?

1dVar retrievals with and without H undate are compared



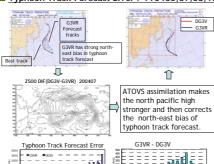
Fixed basic state of H is acceptable

| | Coef | Intercept | Slope | | COBI | Intercept | Slope | |
|--------|-------|-----------|-------|--------|-------|-----------|-------|--|
| SFC T | 1 | 0.481 | 0.998 | SFC q | 1 | 0.022 | 0.998 | |
| 1000 T | 1 | 0.624 | 0.998 | 1000 q | 1 | 0.027 | 0.997 | |
| 925 T | 1 | 0.569 | 0.998 | 925 q | 1 | 0.035 | 0.993 | |
| 850 T | 1 | 0.315 | 0.999 | 850 q | 0.999 | 0.033 | 0.987 | |
| 700 T | 1 | 0.003 | 1 | 700 q | 0.999 | 0.049 | 0.983 | |
| 600 T | 1 | -0.31 | 1.001 | 600 q | 0.999 | 0.06 | 0.985 | |
| 500 T | 1 | -0.837 | 1.003 | 500 q | 0.998 | 0.069 | 0.985 | |
| 400 T | 1 | -1.426 | 1.005 | 400 q | 0.992 | 0.06 | 0.982 | |
| 300 T | 0.998 | -1.384 | 1.005 | 300 q | 0.993 | 0.017 | 0.987 | |
| 250 T | 0.999 | -0.274 | 1.001 | MSLP | 1 | 0.683 | 0.999 | |
| 200 T | 1 | 0.395 | 0.998 | SFCU | 1 | -0.032 | 1.001 | |
| 150 T | 1 | -0.445 | 1.002 | SFC V | 1 | -0.007 | 1.001 | |
| 100 T | 1 | 0.086 | 1 | | | | | |
| 70 T | 1 | 0.141 | 0.999 | | | | | |
| 50 T | 1 | 0.103 | 1 | | | | | |
| 30 T | 1 | 0.033 | 1 | | | | | |
| 20 T | 1 | 0.011 | 1 | | | | | |
| 10 T | 1 | -0.005 | 1 | | | | | |

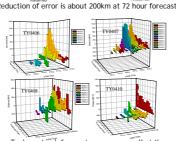


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Typhoon Track Forecast Error / TY0406,07,08,10



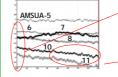
Reduction of error is about 200km at 72 hour forecast



Typhoon track forecast errors are small at the initial stage of typhoon development!

Problems in stratosphere

Time series of global averaged Innovation

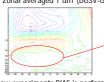


 KMA global model has strong warm bias in the stratosphere

•In order to correct the warm bias, the ATOVS stratospheric channel data is assimilated without bias correction.

· However the negative innovation of stratospheric channels are increased.

Zonal averaged T diff (DG3V-G3OI)



•The warm bias is corrected in the stratosphere.

·However, the warm bias is strongly increased from 300 to 100hPa

It is caused by the negative vertical temperature error

•The bias correction should not apply selectively

New experiments BIAS is performed to solve the problems in the stratosphere BIAS: DG3V + Bias correction is applied in the stratospheric channels depending on the latitude

One month averaged RMSE of 500hPa Bias correction is important to

Summary

- Direct assimilation of ATOVS improves forecast performance and it becomes operation at KMA in August 2004.
- Typhoon forecast errors are reduced by the ATOVS direct assimilation and it is assumed to be caused by the better analysis of the north pacific high.
- Careful investigation is needed to correct model bias by using the ATOVS data because of the vertical error correlation of model error

Plans

- FGAT for ATOVS direct assimilation
- ATOVS level 1C data assimilation in the Unified 3dVar which was developed to run the global and regional application with the same code