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# Beyond optimal estimation: sensitivity of analysis error to mis-specification of background error

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# Beyond optimal estimation: sensitivity of analysis error to mis- specification of background error

- Motivation
- Theory of analysis/retrieval error
  - optimal estimation
  - sub-optimal case
- Illustration
  - scalar case
  - IASI example
- Conclusions and further work



# Motivation (1)

What improvements are needed to exploit advanced IR sounder data more fully in NWP?

- Efficient processing of the full spectrum
- Observation errors, including correlations
- Residual biases
- Surface properties over land and ice
- **Background error statistics**
- Treatment of cloud



## Motivation (2)

- Optimal estimation (OE) theory
  - ... assumes the error covariances are known.
  - In practice, they are not known.
- 2 ways forward:
  - improve estimates of covariances – continuing work
  - make assimilation/retrieval robust against our **inevitable** lack of knowledge
- Applies to both background and obs error covs, B and R
  - in this presentation, **only B considered**



## Motivation (3)

- Why is B **inevitably** in error?
  - global averages can be estimated quite accurately
  - ... but large spatial/temporal variability.
- We need to understand our sensitivity to B and its **inevitable** mis-specification,
  - particularly for satellite radiances ...
    - non-local observations (→ Fiona Hilton's paper)



## Motivation (4)

- Advanced IR sounders have **vertical resolution**  $\sim 1$  km
  - sensitive with low error to scales  $\gg 1$  km
  - not sensitive to scales  $\ll 1$  km
  - sensitive to scales  $\sim 1$  km, but with errors comparable to background errors
- Need to understand  $B$  - its magnitude on different scales
  - determines how measurements and prior information are weighted **on each scale**
- ... and effects of mis-specifying  $B$  **on each scale**



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# Theory (1)

## GENERAL CASE

Analysis equation (linearised):  $x^a = x^b + K \cdot (y^o - H[x^b])$

Analysis error equation:  $\varepsilon^a = \varepsilon^b + K \cdot (\varepsilon^o - H \cdot \varepsilon^b)$

$$\varepsilon^a = (I - K \cdot H) \cdot \varepsilon^b + K \cdot \varepsilon^o$$

Analysis error covariance:  $A = (I - K \cdot H) \cdot B \cdot (I - K \cdot H)^T + K \cdot R \cdot K^T$

## OPTIMAL CASE

assumed value  $B_A =$  true value  $B$

$$K = B_A \cdot H^T \cdot (H \cdot B_A \cdot H^T + R)^{-1}$$

$$A_{\text{OPT}} = (I - K \cdot H) \cdot B_A \cdot (I - K \cdot H)^T + K \cdot R \cdot K^T$$

$$= (I - K \cdot H) \cdot B_A$$

$$A_{\text{OPT}}^{-1} = B_A^{-1} + H^T \cdot R^{-1} \cdot H$$



## Theory (2)

Projecting on to the eigenvectors of  $B_A$  :

$V$  = eigenvectors of  $B_A$ ;  $\Lambda$  = eigenvalues of  $B_A$

$$A_{\text{OPT}}^{-1} = B_A^{-1} + H^T.R^{-1}.H$$

$$V^T.A_{\text{OPT}}^{-1}.V = V^T.B_A^{-1}.V + V^T.H^T.R^{-1}.H.V$$

$$V^T.A_{\text{OPT}}^{-1}.V = \Lambda^{-1} + V^T.H^T.R^{-1}.H.V$$

### Why $B_A$ ?

- because this is what we use – the “filter” within the DA system
  - Met Office 4D-Var performs vertical analysis in this eigen-space





# Theory (3)

## OPTIMAL

$$A_{\text{OPT}}(\mathbf{B}) = (\mathbf{I} - \mathbf{K} \cdot \mathbf{H}) \cdot \mathbf{B} \cdot (\mathbf{I} - \mathbf{K} \cdot \mathbf{H})^T + \mathbf{K} \cdot \mathbf{R} \cdot \mathbf{K}^T$$

$$\mathbf{K}(\mathbf{B}) = \mathbf{B} \cdot \mathbf{H}^T \cdot (\mathbf{H} \cdot \mathbf{B} \cdot \mathbf{H}^T + \mathbf{R})^{-1}$$

GENERAL / SUB-OPTIMAL, which means  $\mathbf{B} \neq \mathbf{B}_A$ ,  $\mathbf{K} = \mathbf{K}(\mathbf{B}_A)$

$$A(\mathbf{B}) = (\mathbf{I} - \mathbf{K}(\mathbf{B}_A) \cdot \mathbf{H}) \cdot \mathbf{B} \cdot (\mathbf{I} - \mathbf{K}(\mathbf{B}_A) \cdot \mathbf{H})^T + \mathbf{K}(\mathbf{B}_A) \cdot \mathbf{R} \cdot \mathbf{K}(\mathbf{B}_A)^T$$

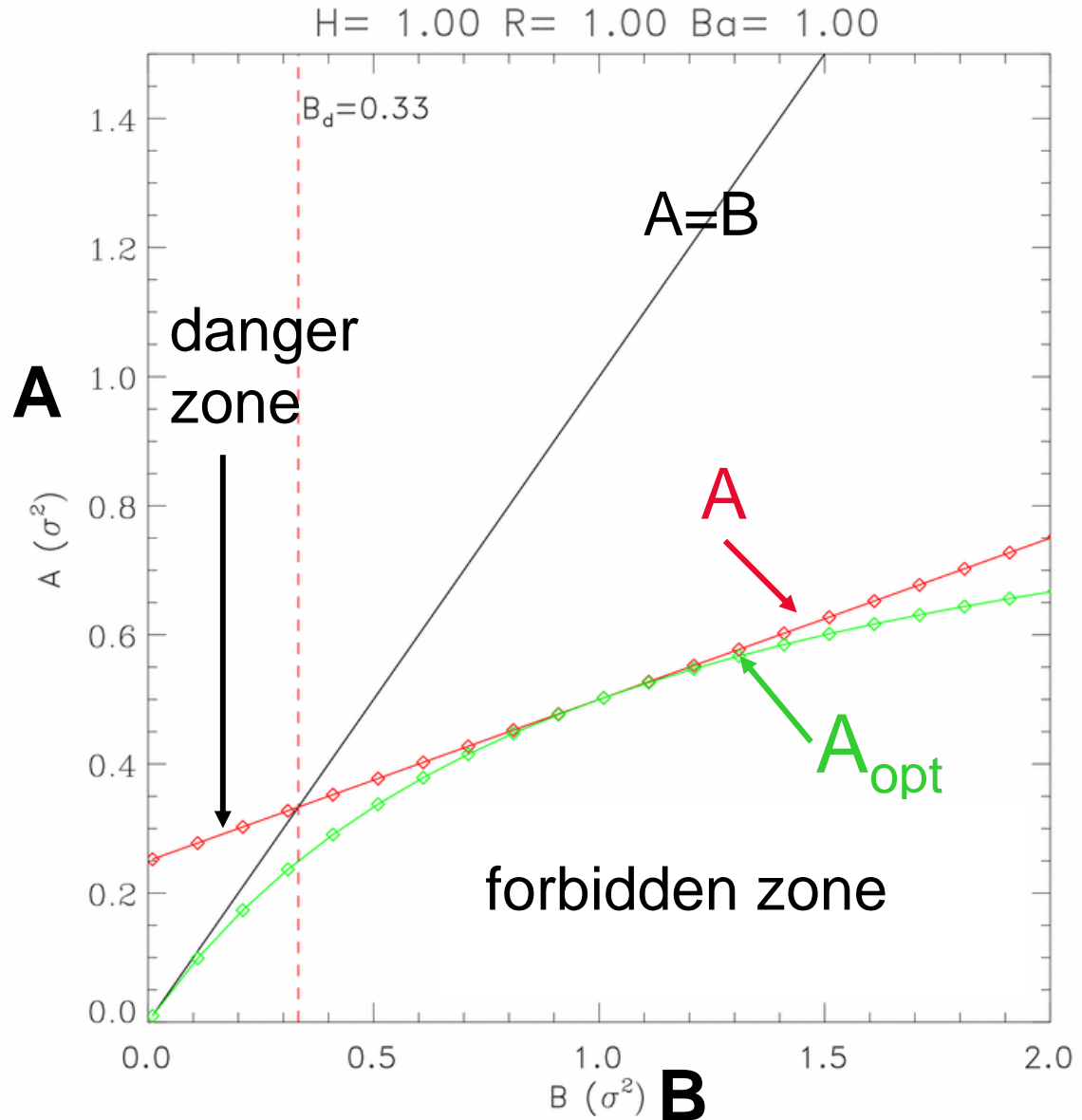
$$A(\mathbf{B}) = A_{\text{OPT}}(\mathbf{B}_A) + (\mathbf{I} - \mathbf{K}(\mathbf{B}_A) \cdot \mathbf{H}) \cdot (\mathbf{B} - \mathbf{B}_A) \cdot (\mathbf{I} - \mathbf{K}(\mathbf{B}_A) \cdot \mathbf{H})^T$$

Note: linear in  $\mathbf{B}$



# Illustration – scalar case (1)

$H = 1$   
 $R = 1$   
 $B_A = 1$



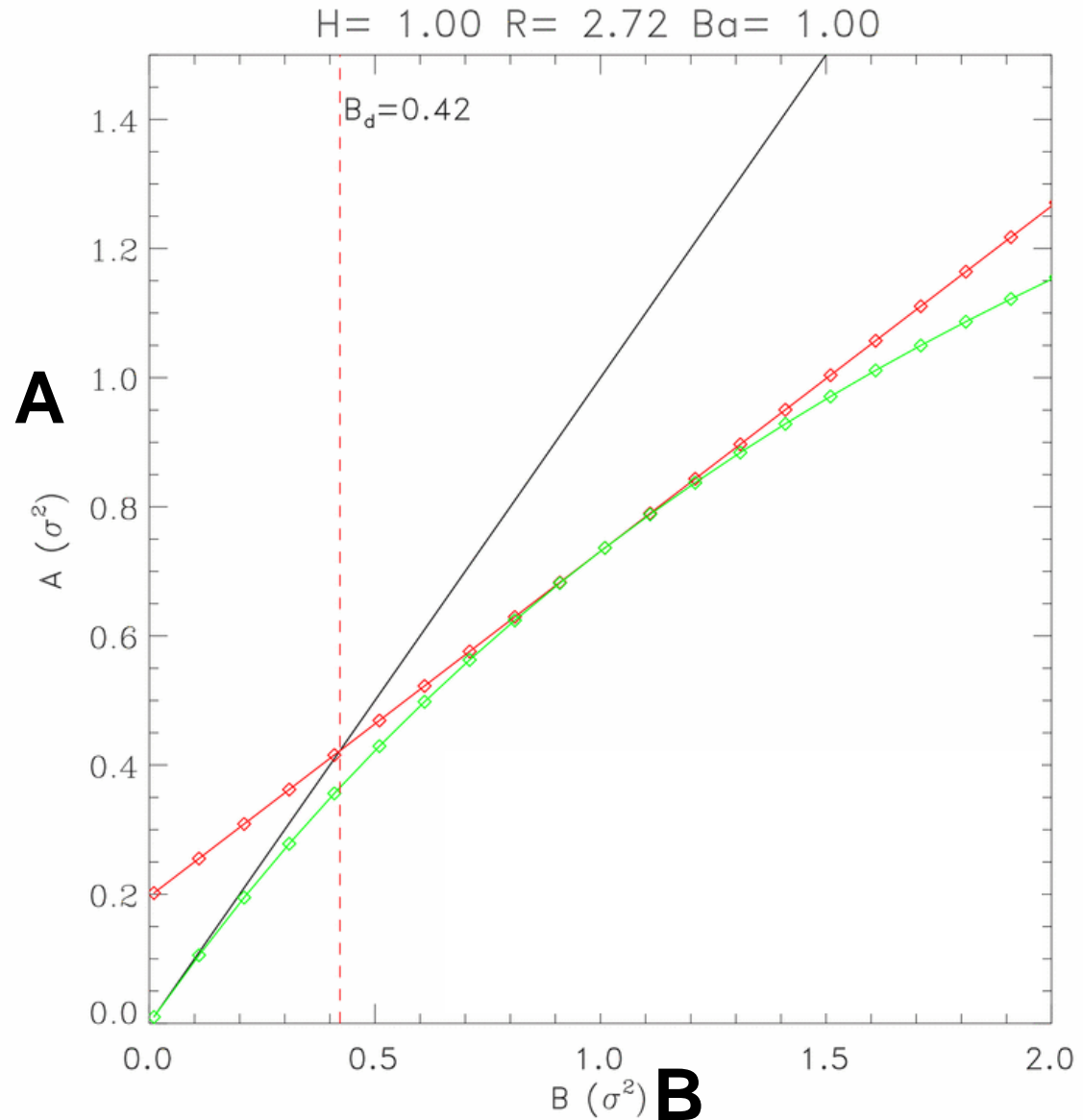


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# Illustration – scalar case (2)

$H = 1$   
 $R = 2.72$   
 $B_A = 1$

higher  
observation  
error





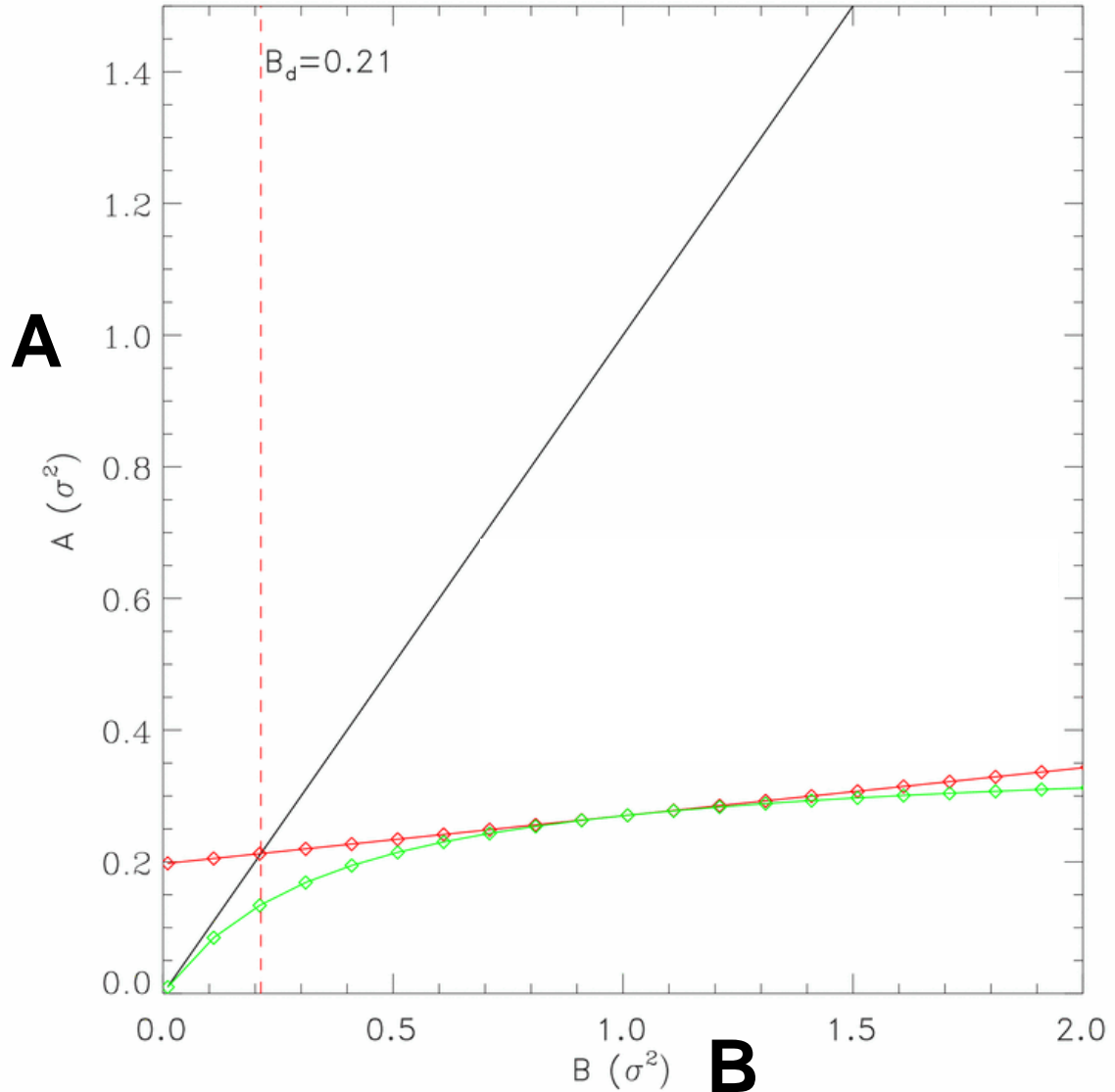
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# Illustration – scalar case (3)

$H = 1$   
 $R = 0.37$   
 $B_A = 1$

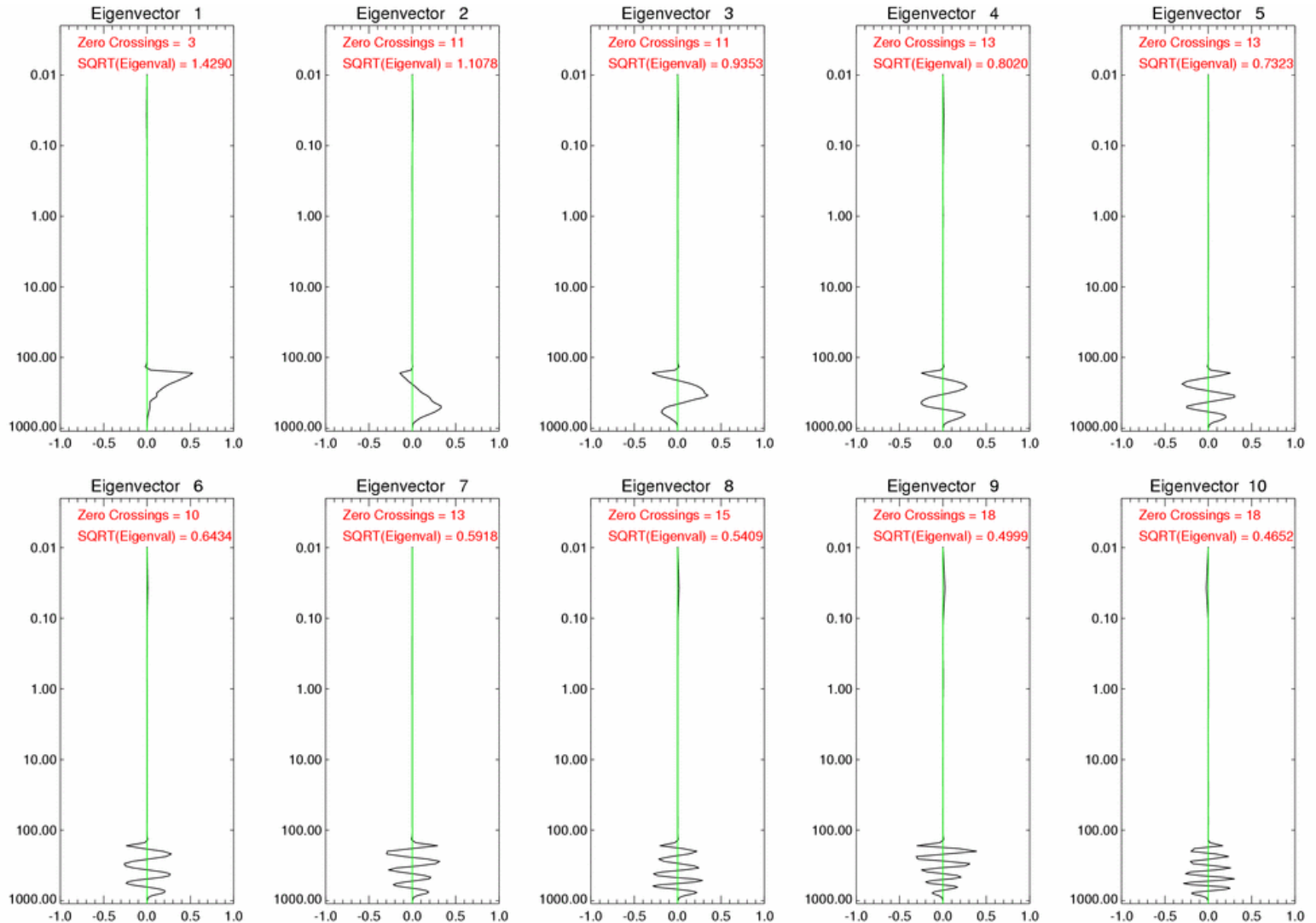
lower  
observation  
error

$H = 1.00$   $R = 0.37$   $B_a = 1.00$



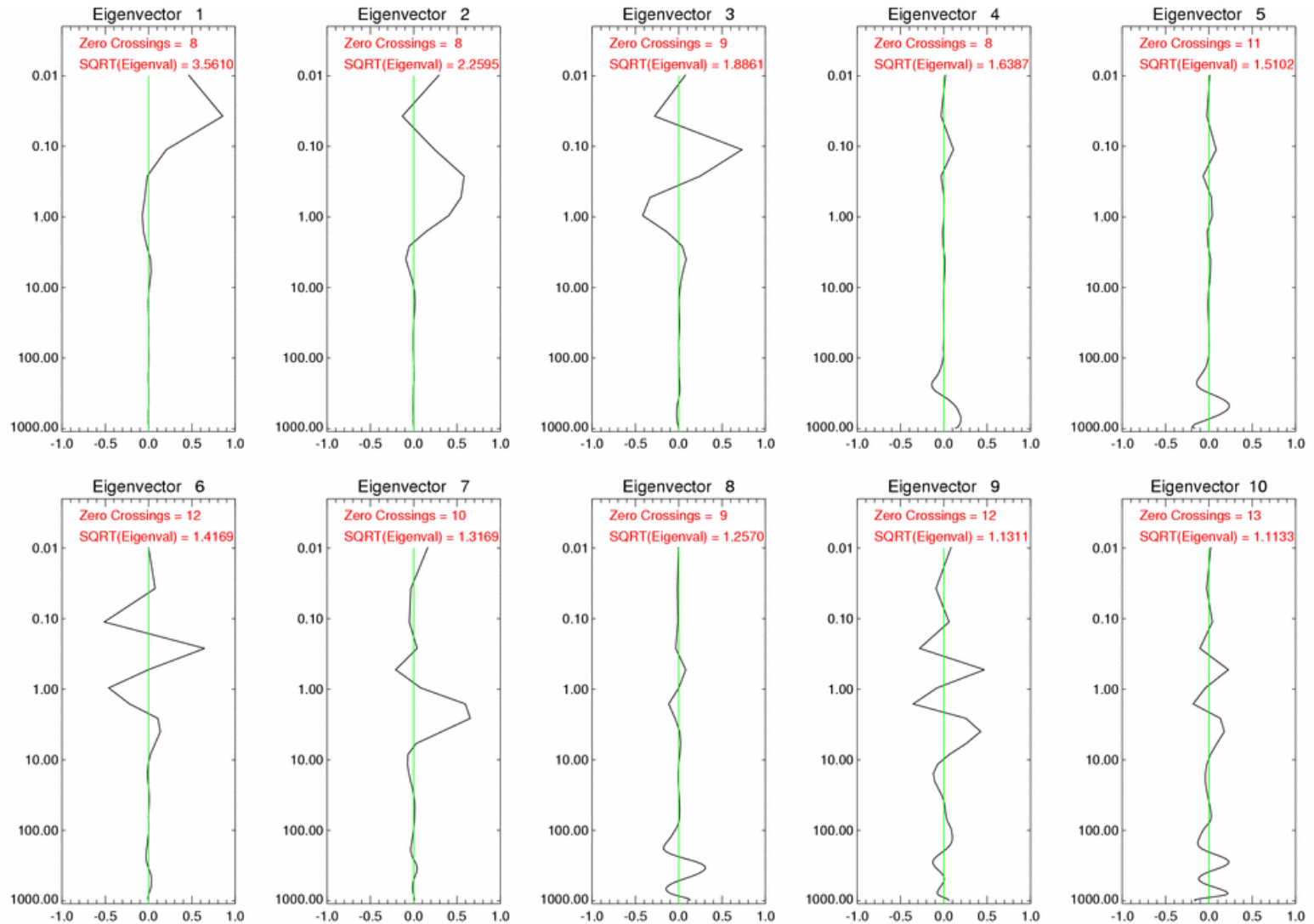


# Leading eigenvectors of $B_A$ MetO 70-level model, $\ln(q)$ (vectors 1-10)



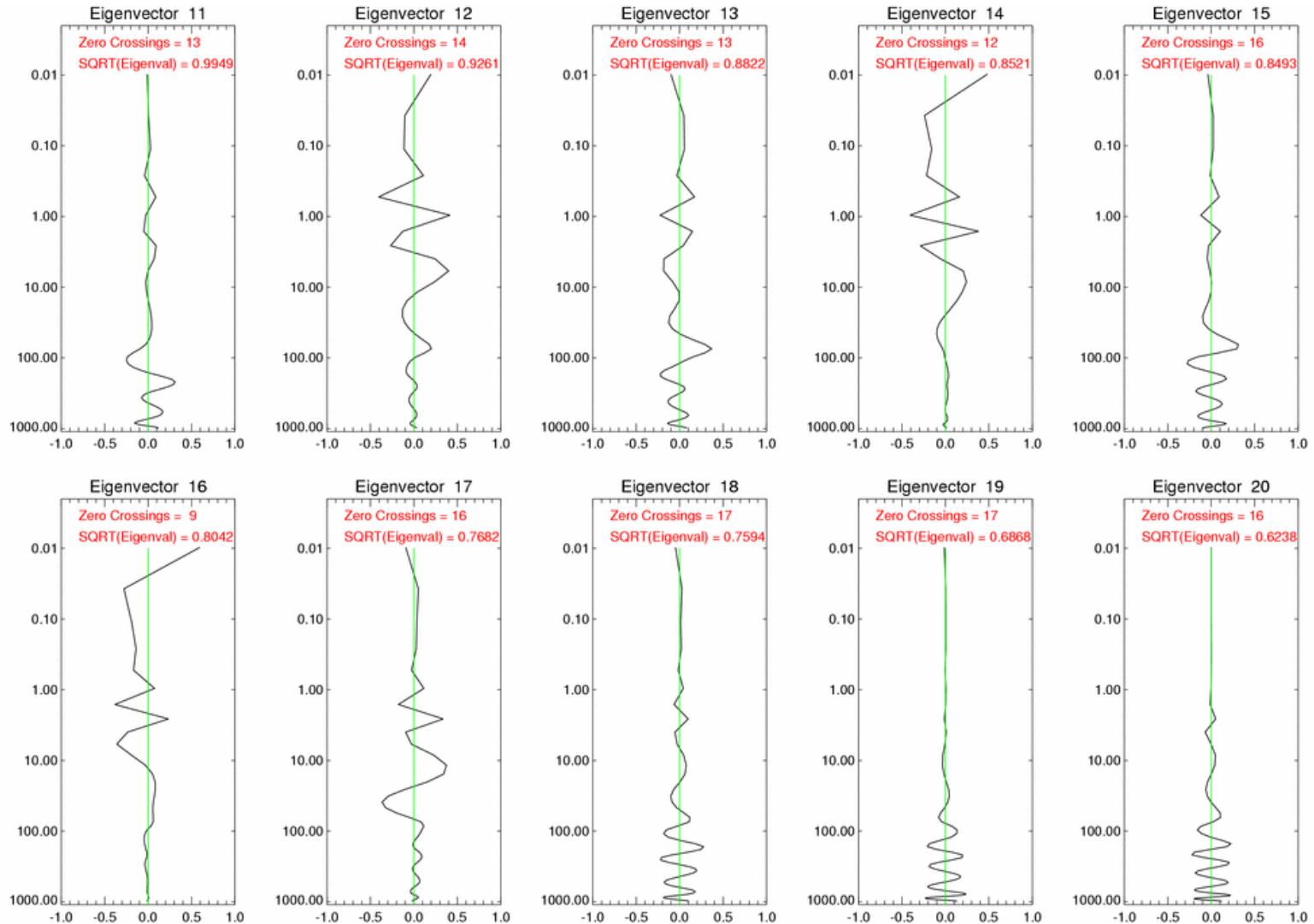


# Leading eigenvectors of $B_A$ MetO 70-level model, temp. (vectors 1-10)





# Leading eigenvectors of $B_A$ MetO 70-level model, temp. (vectors 11-20)





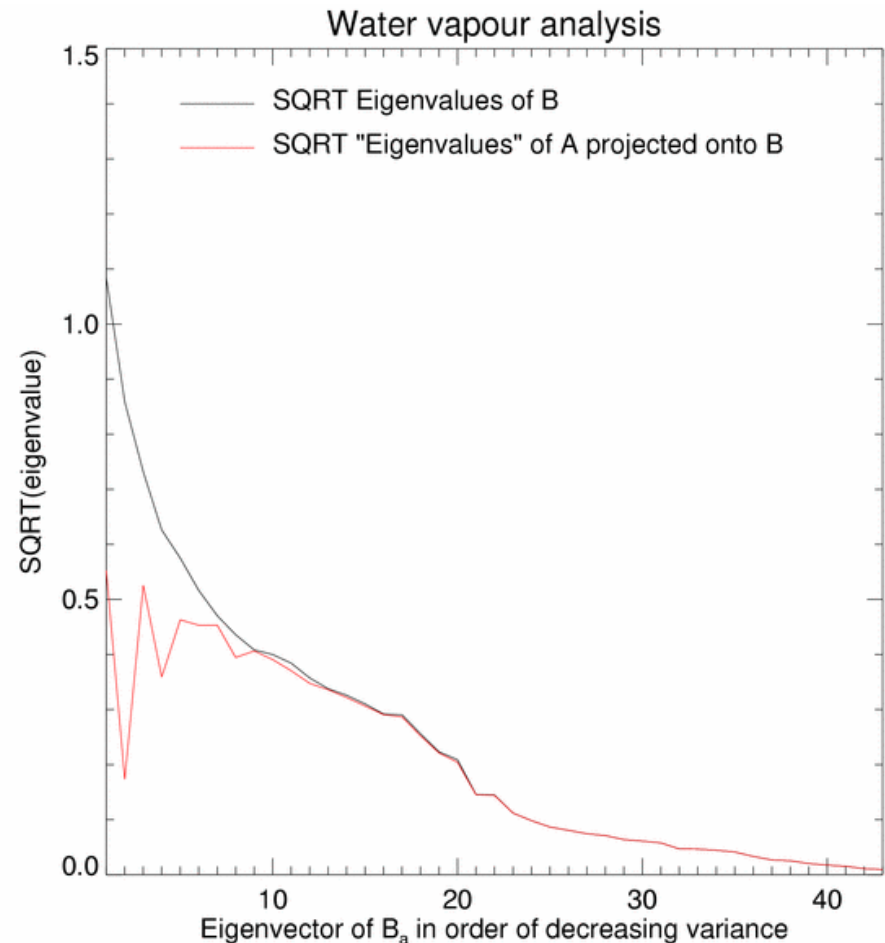
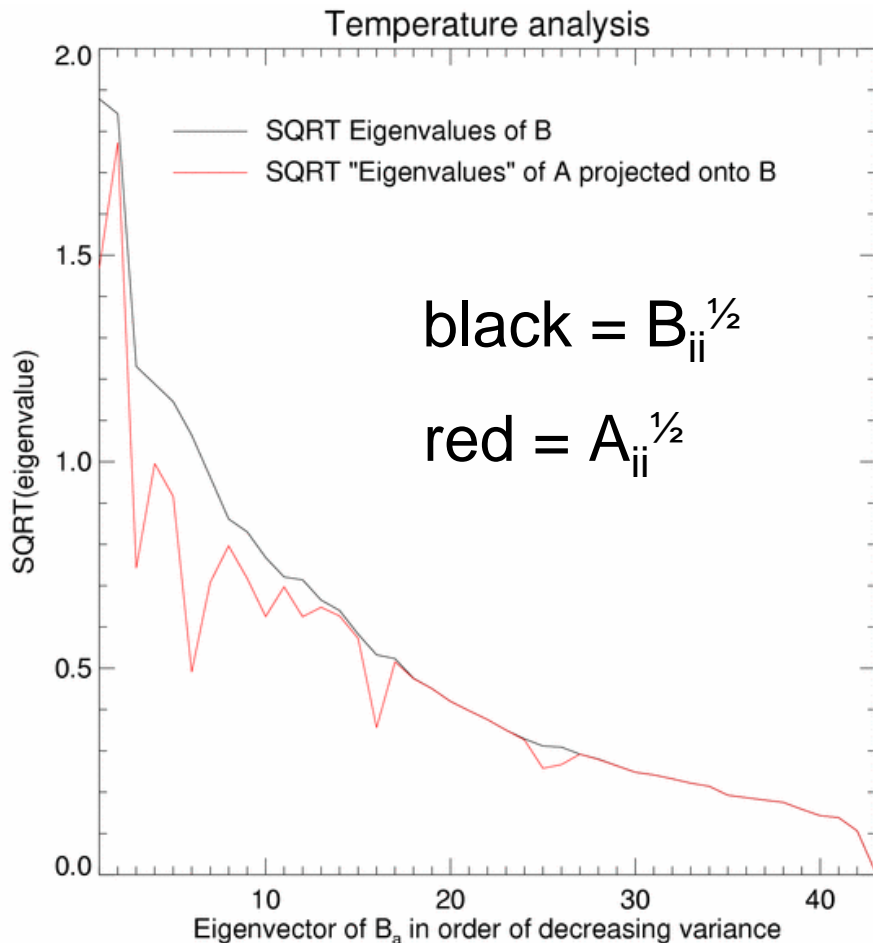
# Illustration – IASI (1)

- Met Office operational 1D-var channel selection
  - 183 channels, of which 31 in water vapour band
- Observation error
  - instrument noise, or
  - instrument noise + forward model error of 0.2K + extra for unmodelled trace gas
- Analysis on 43 RTTOV levels using Jacobians from US standard atmosphere



# Diagonal of analysis error mapped to eigenvectors of $B_A$

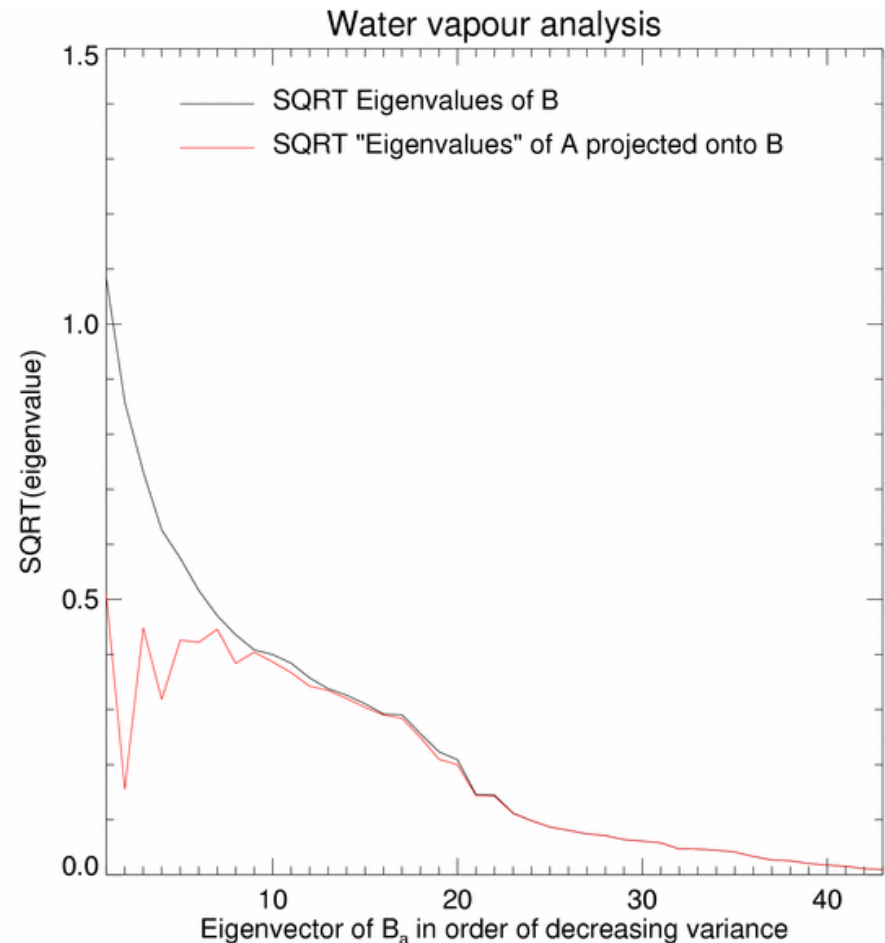
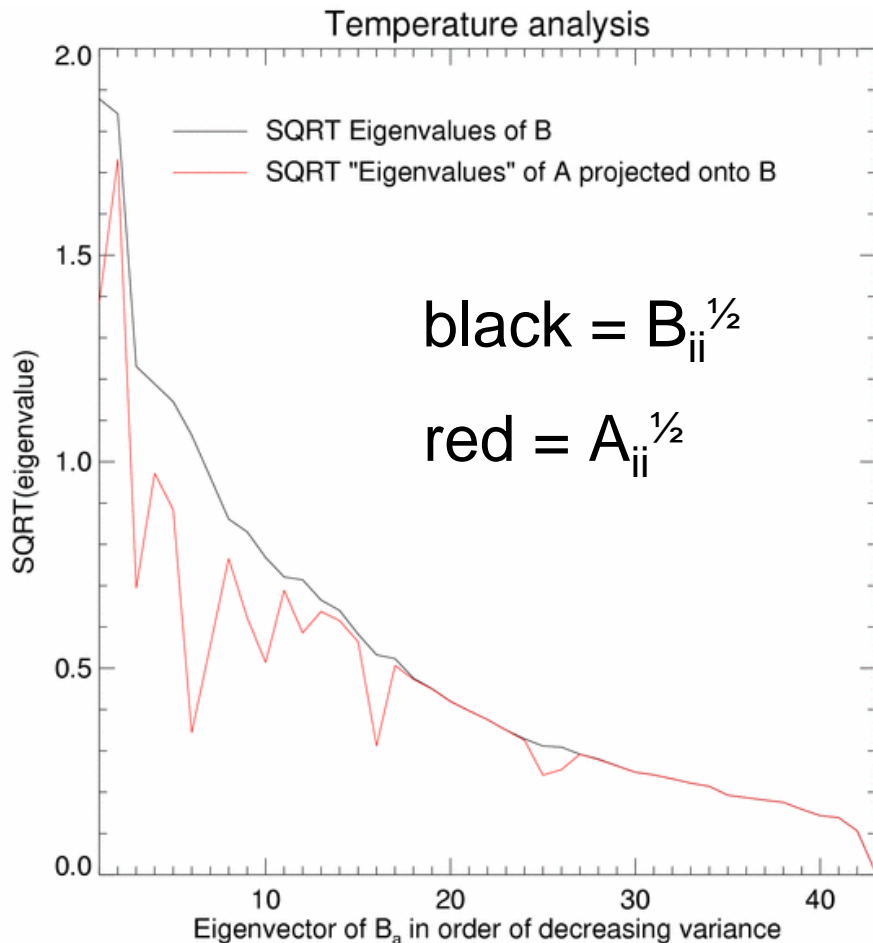
observation error = instrument noise + forward model error





# Diagonal of analysis error mapped to eigenvectors of $B_A$

observation error = instrument noise





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## Conclusions so far ...

- Goal: make retrievals/analyses **robust** against **inevitable** errors in the background error covariance
- ... particularly for effective assimilation of satellite sounder data
- What is crucial for NWP? - structure of B **assumed** by the DA system,  $B_A$
- **Beware the danger zone!** – analysis errors higher than background errors
- Current problems with Met Office 4D-Var B-matrix for temperature
- (provisional result) Some real IASI information is currently filtered out by the assimilation system



# Further work

- Further work needed:
  - to perform a more complete error analysis for IASI
  - to understand  $B_A$  on each scale | good idea, in general
  - ... and to improve it
  - to make  $B_A$  robust against inevitable errors



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Thank you! Questions?

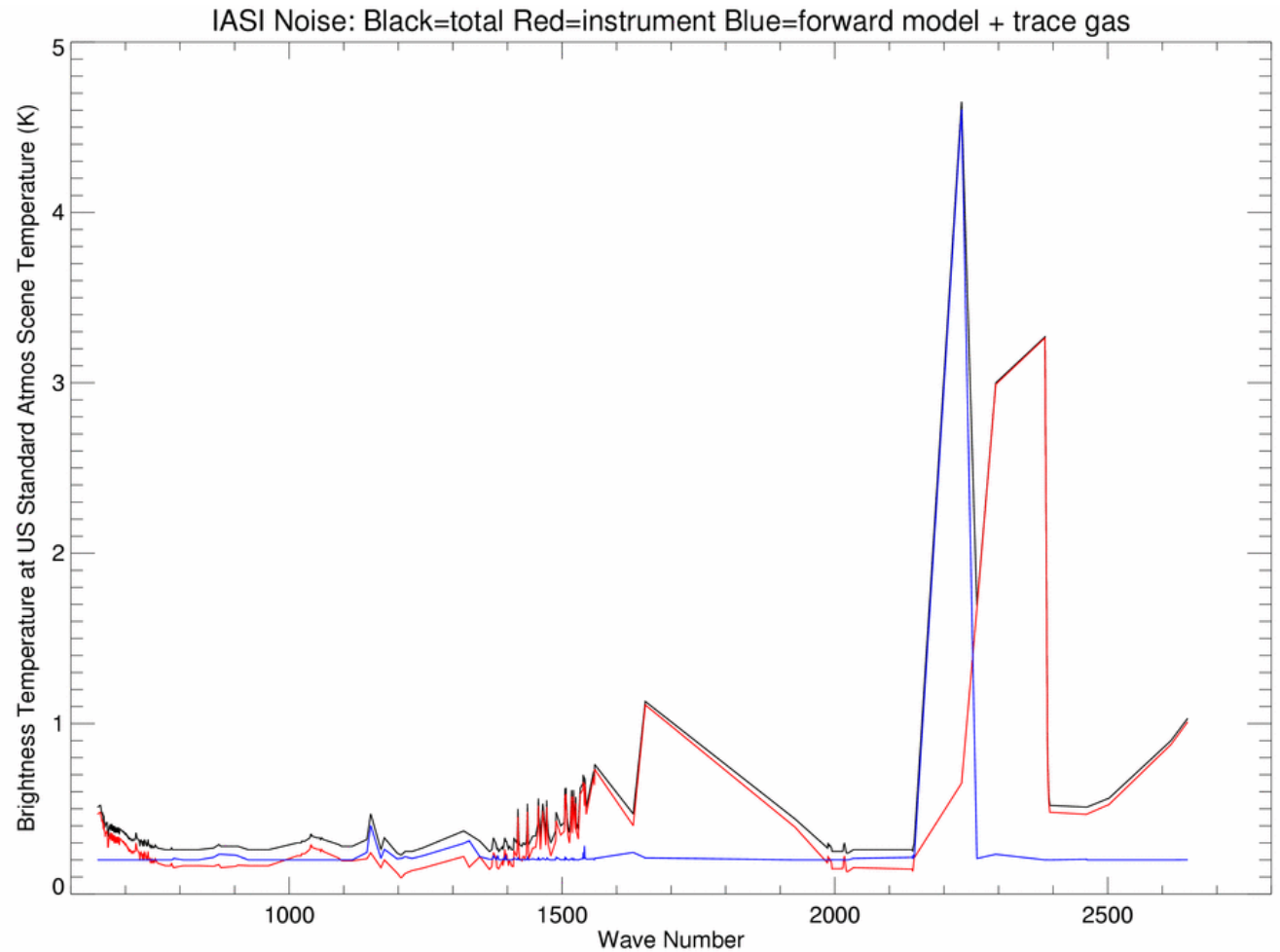


# IASI noise

black: total

red: instrument

blue: forward model + "trace gas noise"





# Goal

- To exploit the improved vertical resolution of advanced IR sounders
- ... whilst retaining the (usually accurate) information from the NWP model on sharp vertical structures
  - e.g. PBL top, tropopause



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# Prior work on mis-specification of errors

- O.N.Strand 1977 The Annals of Statistics
- R.Daley 1991 Atmospheric Data Assimilation
  - R.Seaman 1977 MWR
  - R.Seaman et al. 1983 Aus. Met. Mag.
  - R.Franke 1985 MWR
- P.Watts and A.McNally 1988 Proc. ITSC-IV
- A.McNally 2000 QJRMS
- Others?

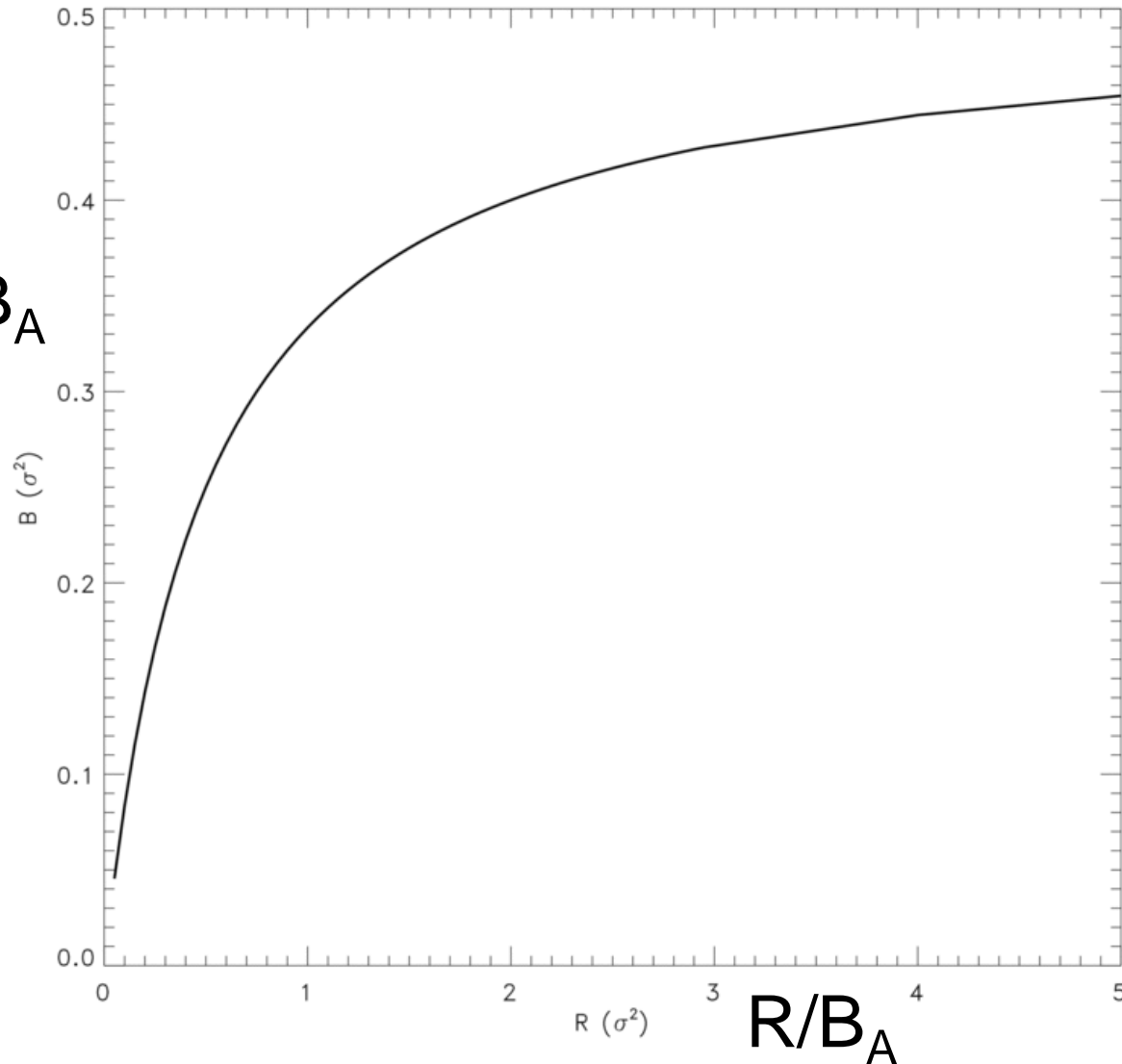




# Scalar case – the danger zone

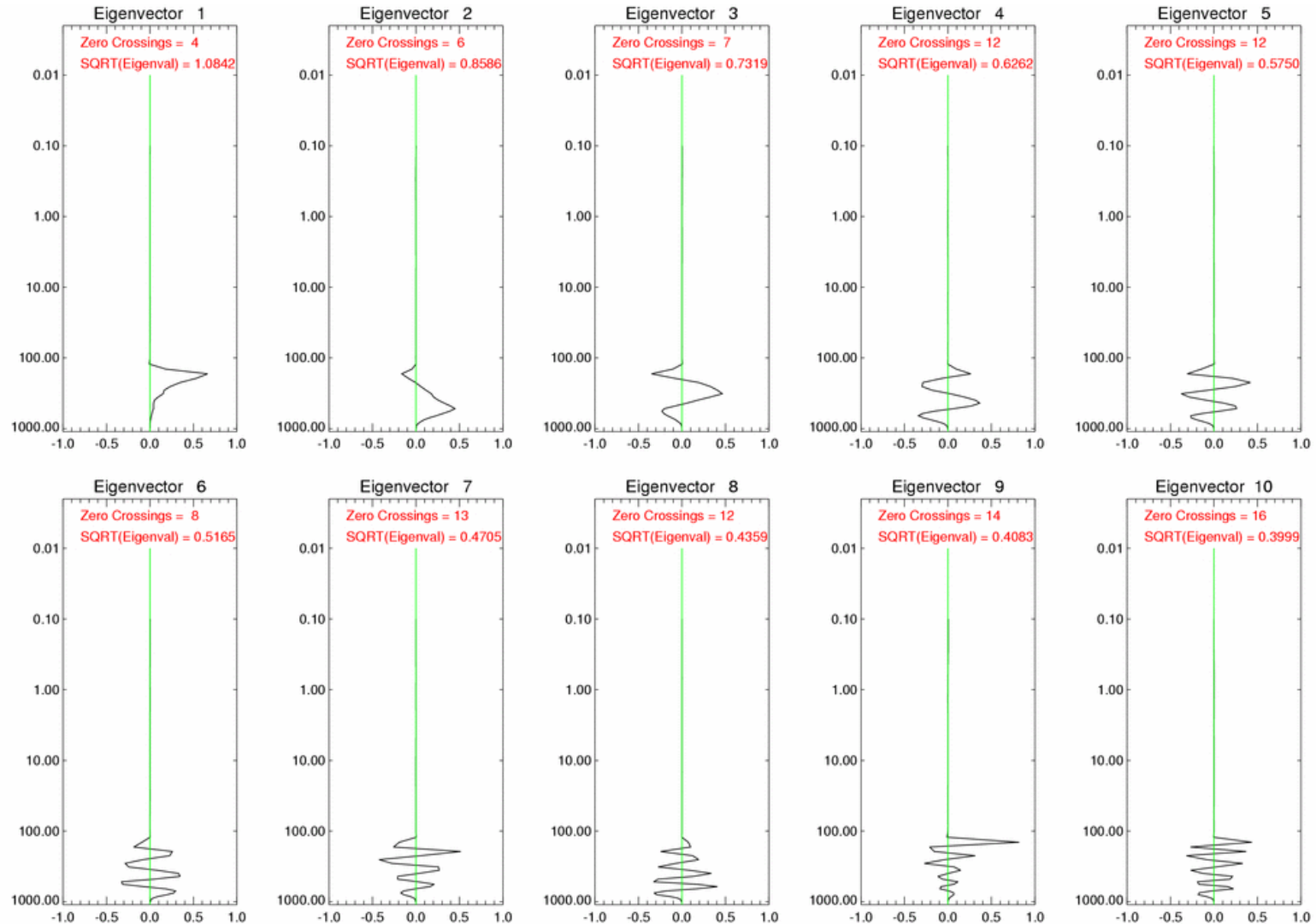
$B_d$ , defining the "Danger Zone" where  $A > B$ ;  
for  $H = 1.00$   $Ba = 1.00$   $K = B/(B+R)$

$B(\text{danger})/B_A$



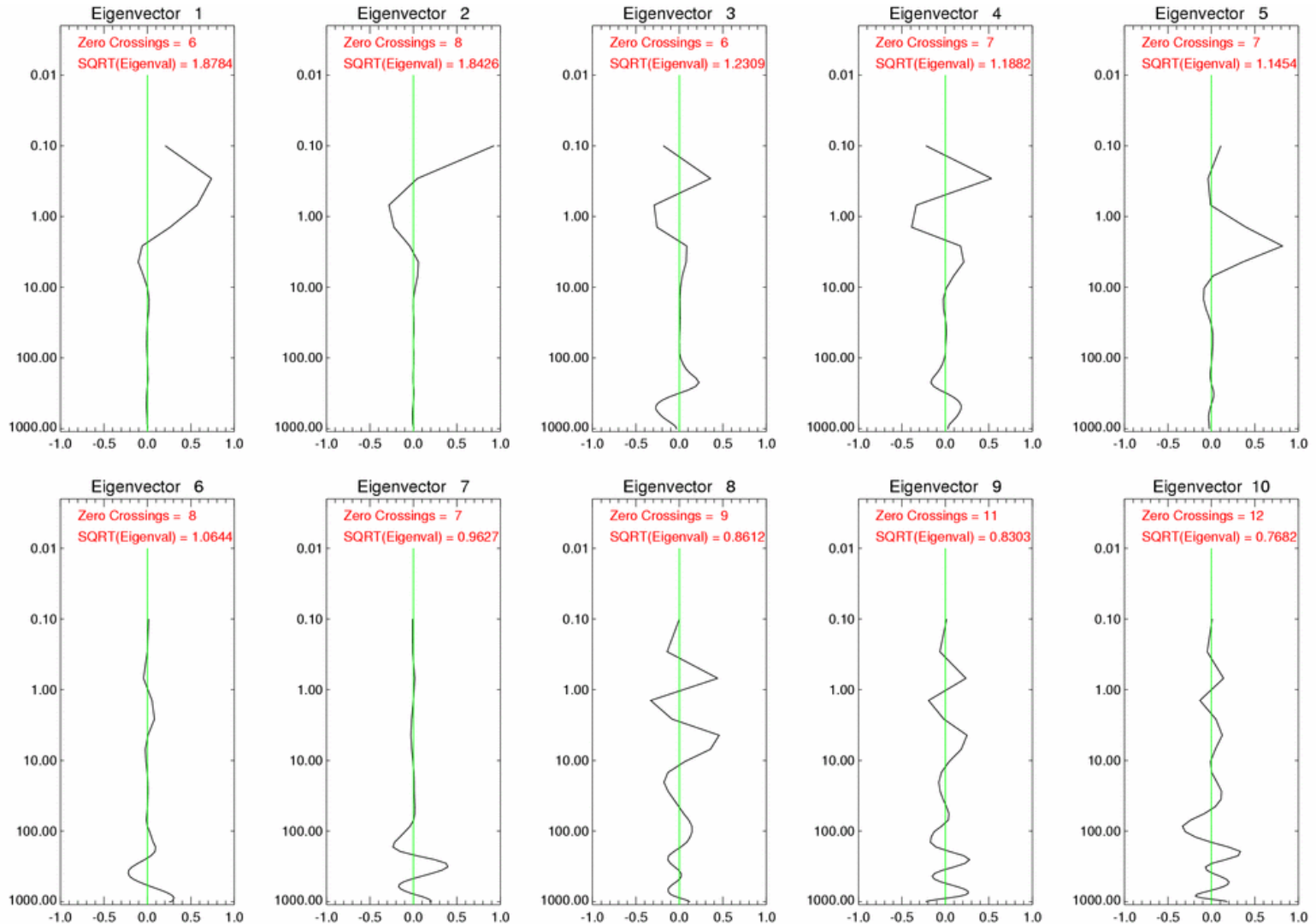


# Leading eigenvectors of $B_A$ 43 RTTOV levels $\ln(q)$ (vectors 1-10)





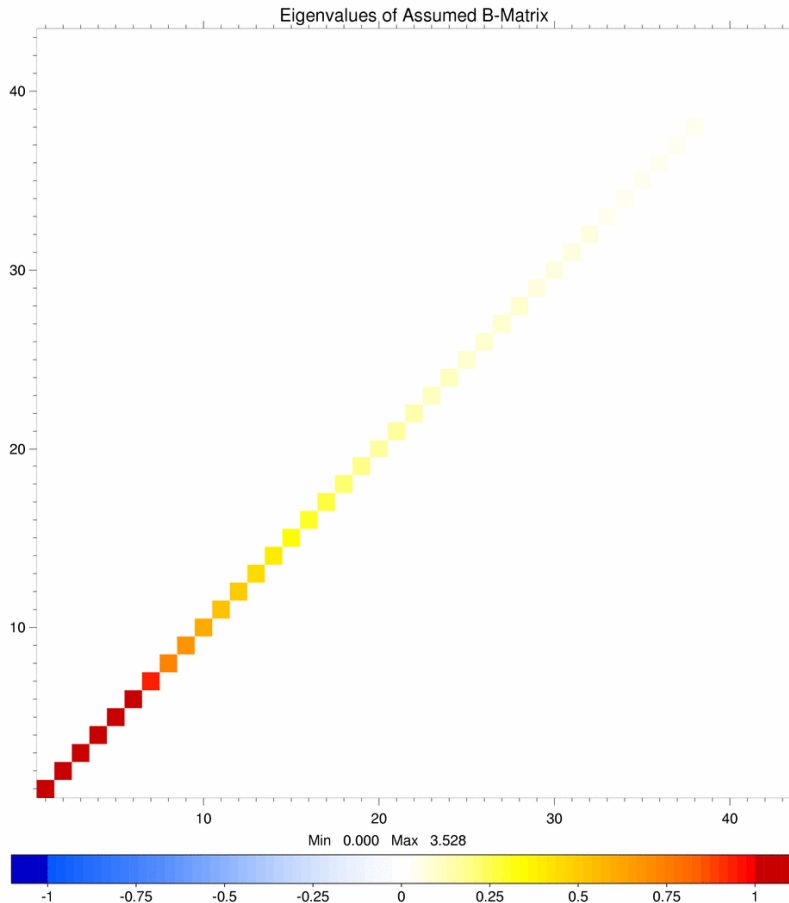
# Leading eigenvectors of $B_A$ 43 RTTOV levels temperature (vectors 1-10)



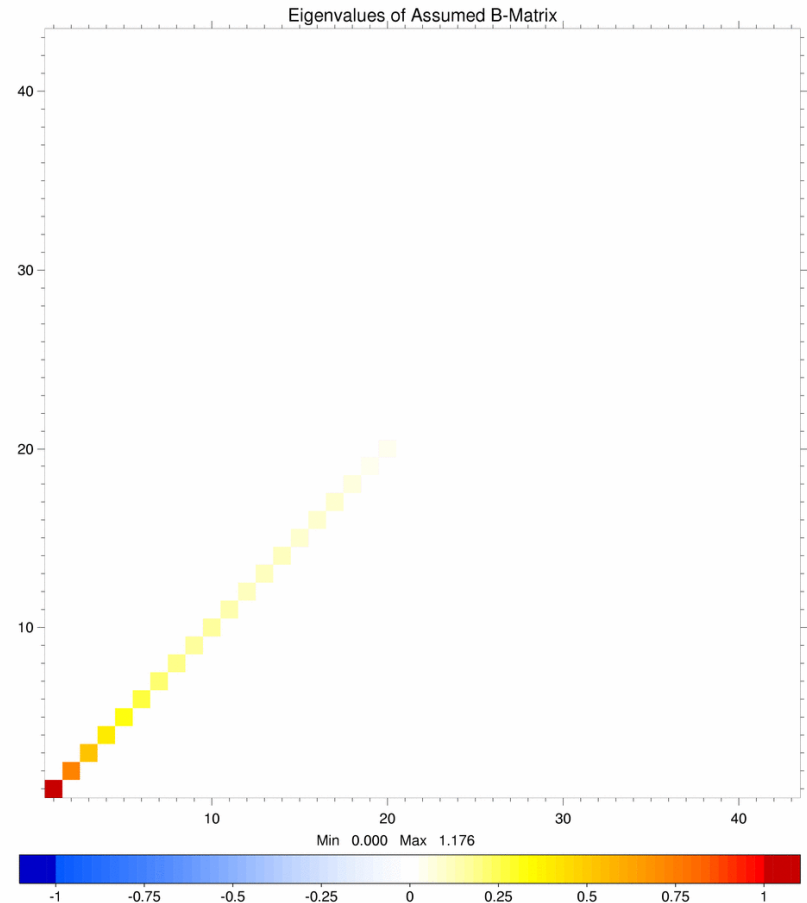


# Eigenvalues of $B_A$ : background errors in the eigenspace of $B_A$

Temperature



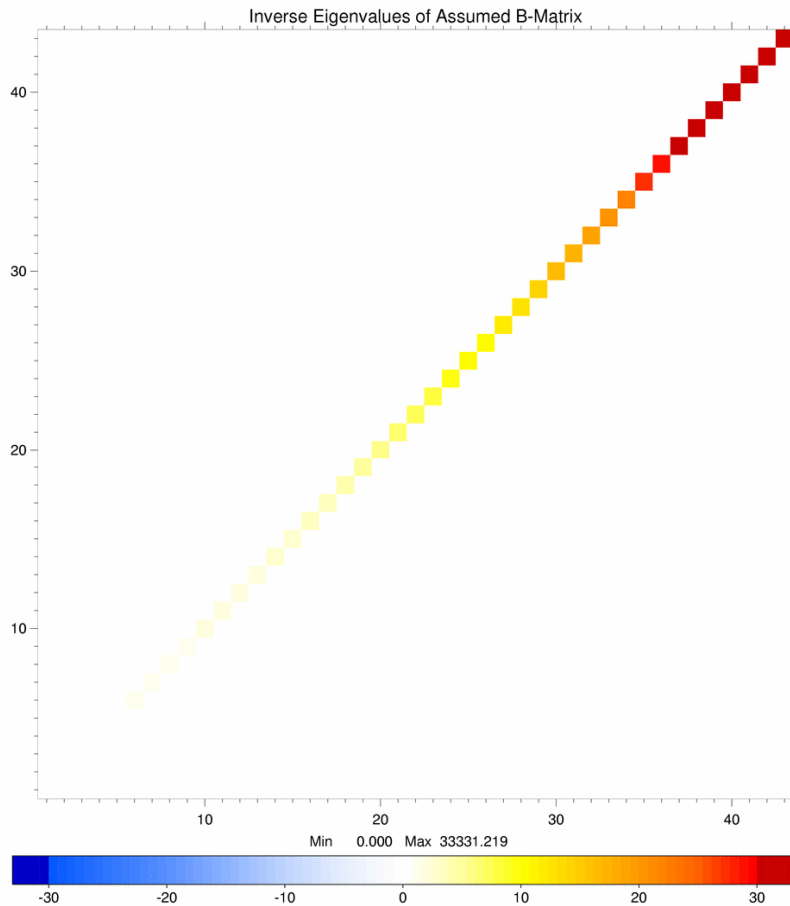
$\ln(q)$



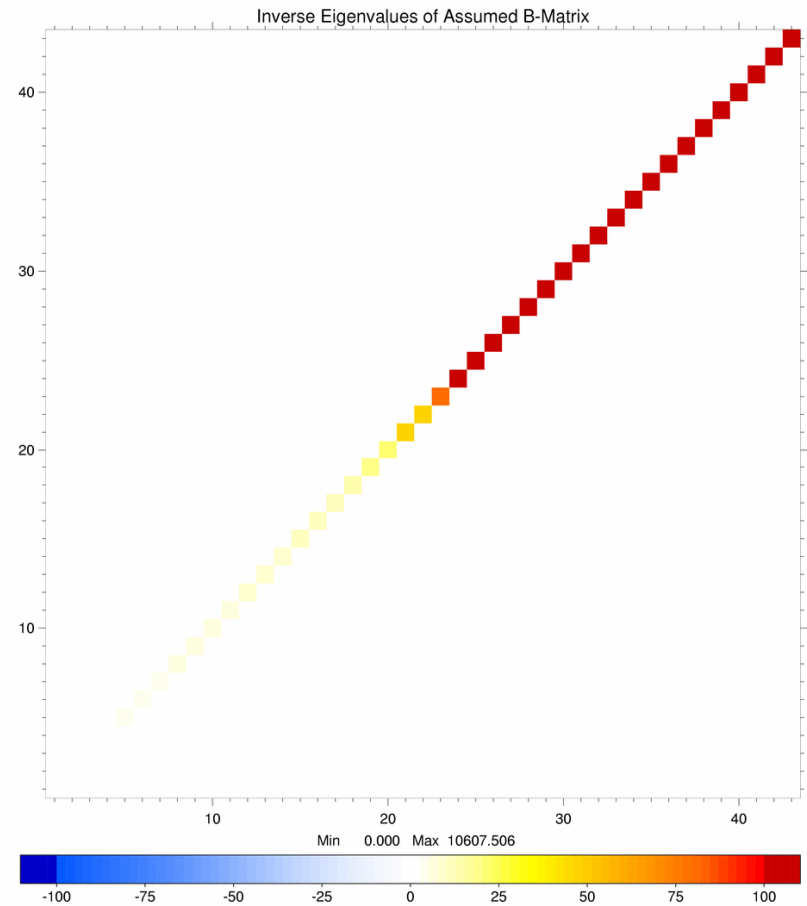


# Inverse eigenvalues of $B_A$

## Temperature



## $\ln(q)$





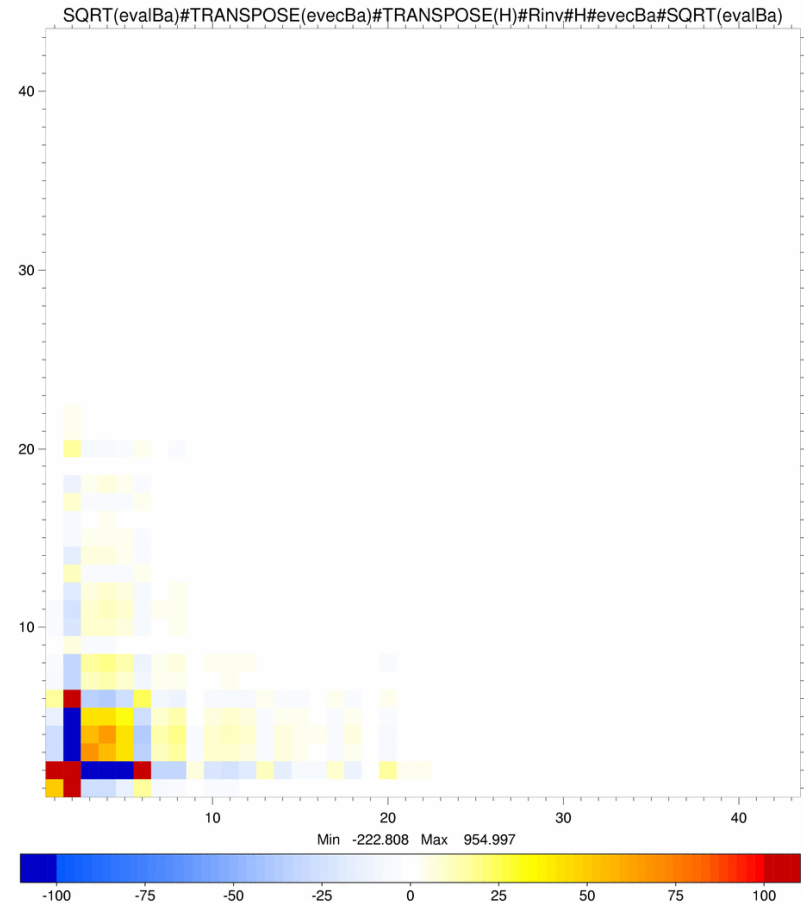
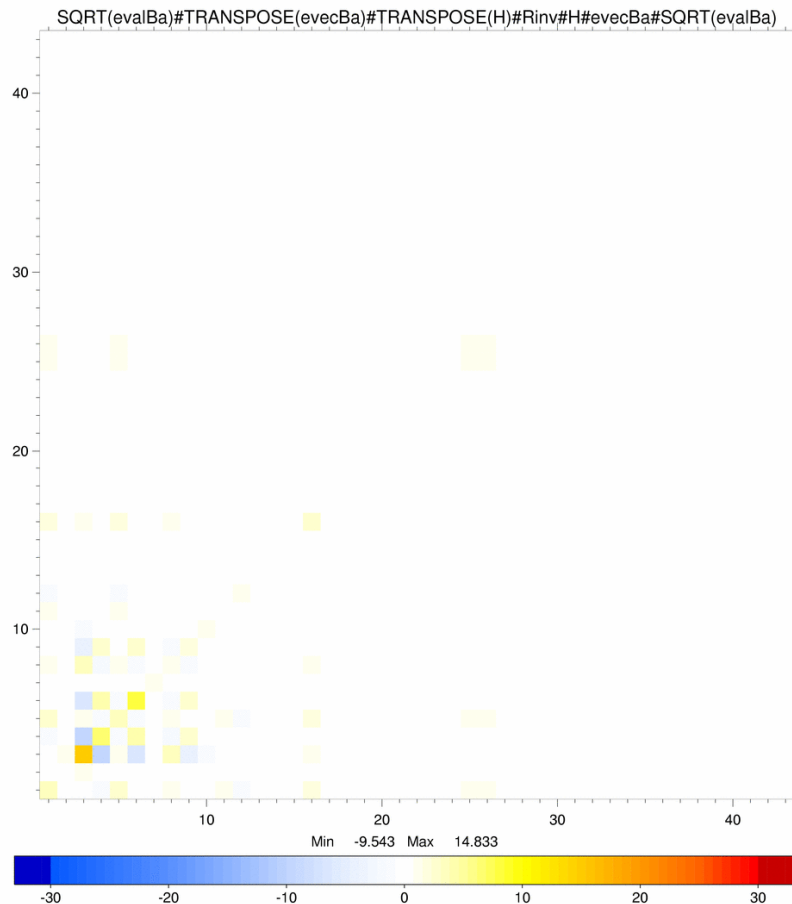
# IASI information mapped to $B_A$ eigenvectors and normalised by $B_A$ eigenvalues:

$$\Lambda^{1/2} \cdot V^T \cdot H^T \cdot R^{-1} \cdot H \cdot V \cdot \Lambda^{1/2}$$

Temperature

Instrument noise + forward model error

$\ln(q)$





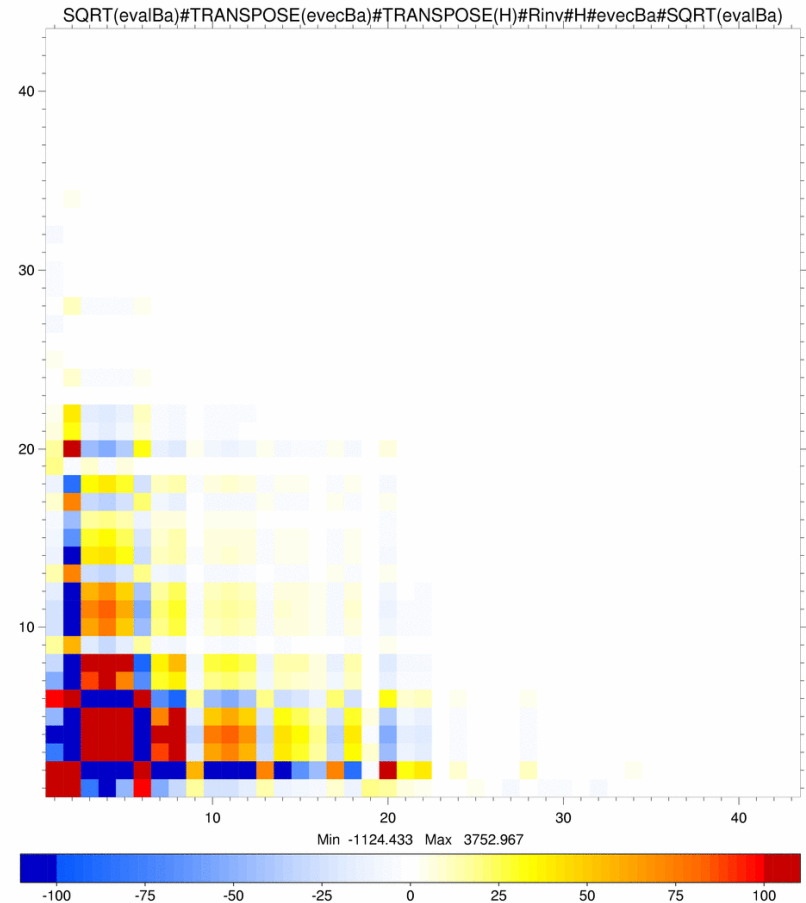
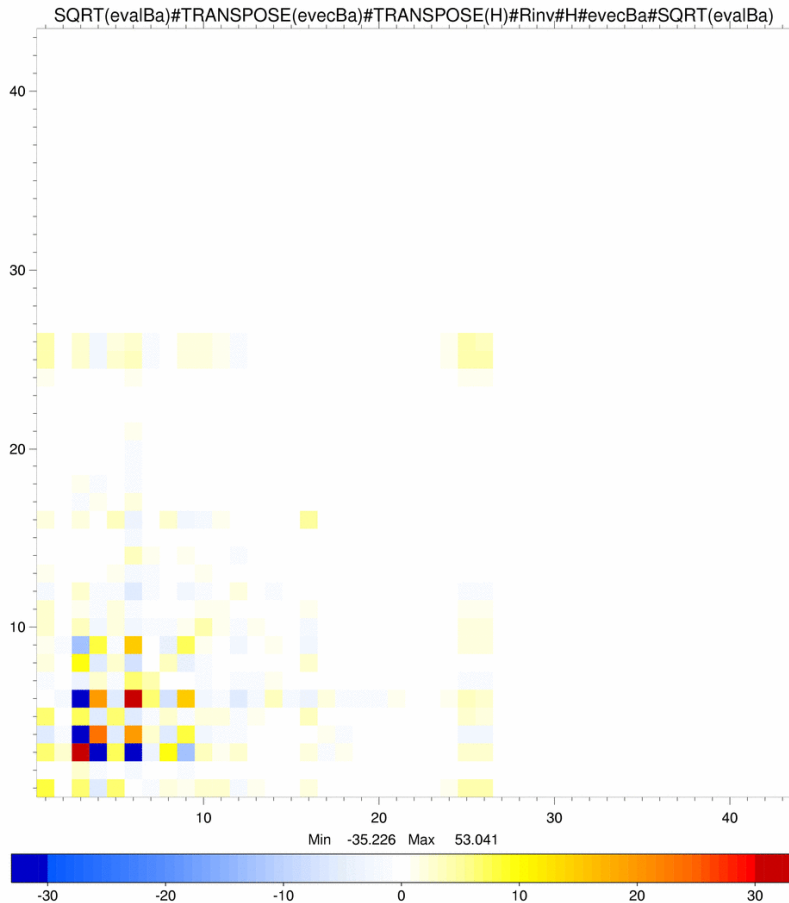
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$$\Lambda^{1/2} \cdot V^T \cdot H^T \cdot R^{-1} \cdot H \cdot V \cdot \Lambda^{1/2}$$

Temperature

Instrument noise only

$\ln(q)$

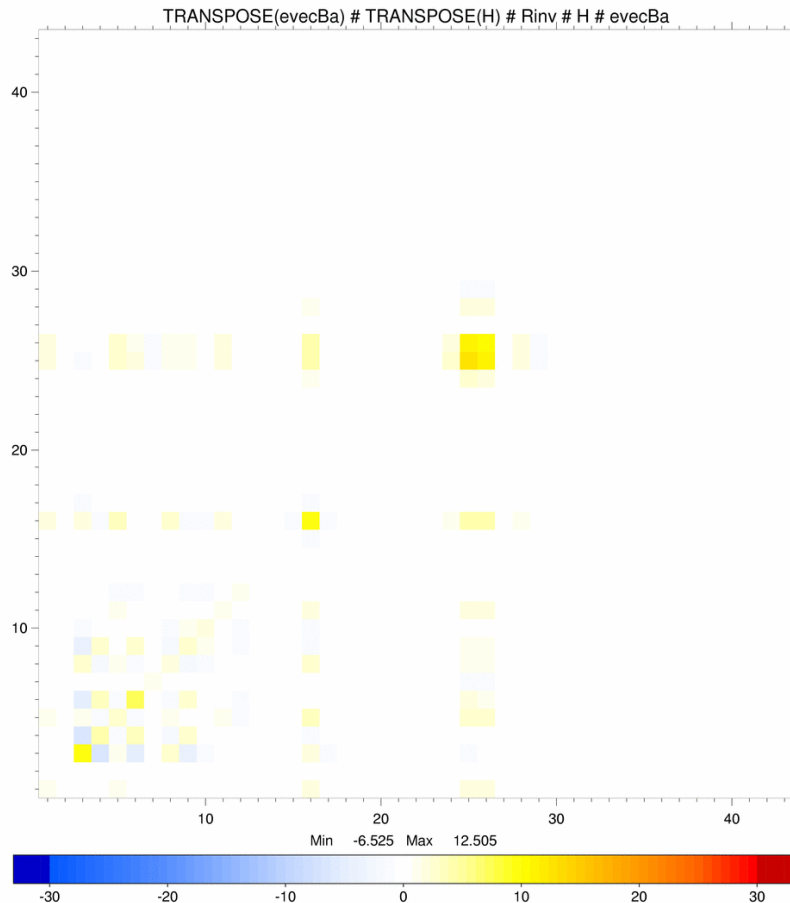




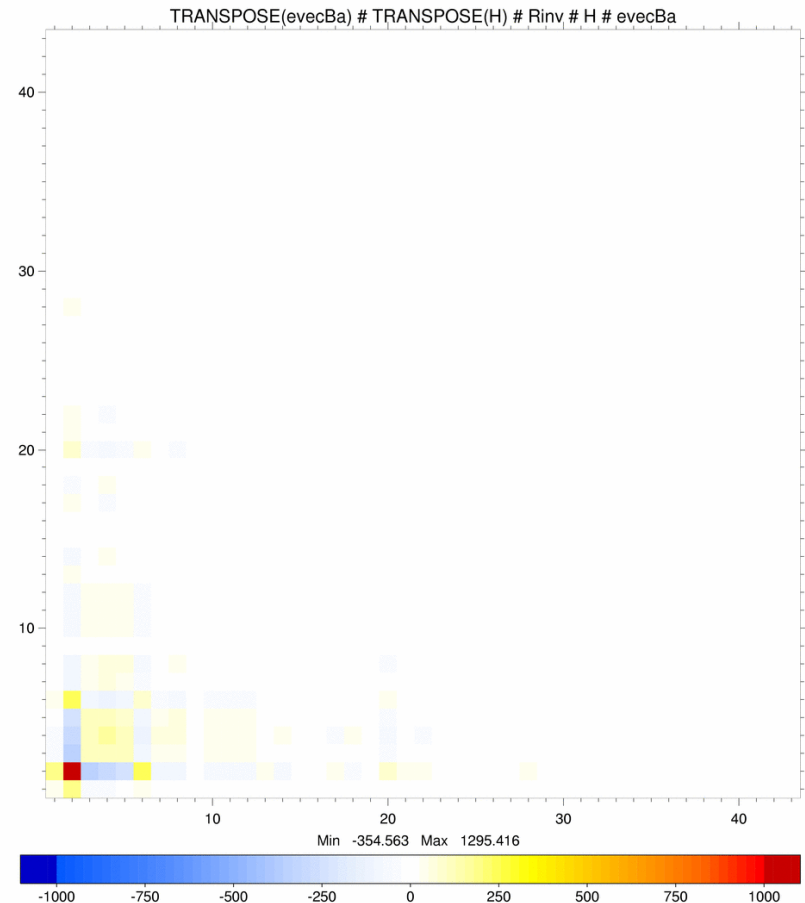
# IASI observation error mapped to eigenvectors of $B_A$ : $V^T \cdot H^T \cdot R^{-1} \cdot H \cdot V$

observation error = instrument noise + forward model error

temperature



ln(q)



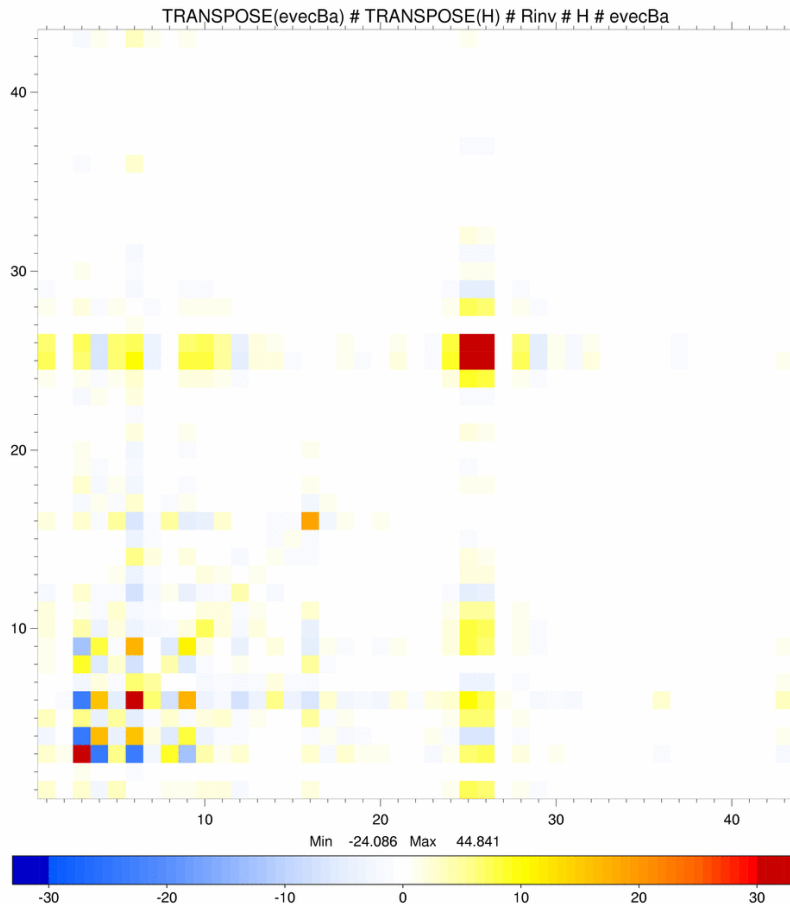




# IASI observation error mapped to eigenvectors of $B_A$

observation error = instrument noise only

temperature



ln(q)

