

*Cirrus mean effective ice crystal sizes
from Satellite
TIROS-N Operational Vertical Sounder (TOVS)
Observations*

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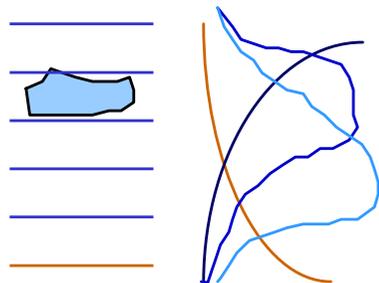
Desert Research Institute, Reno, USA

Longterm Satellite Observations

TOVS vertical sounders on polar satellites: since 1979

Path-B (*Scott et al., BAMS 1999, 1987-1995*)

- ◆ every 6 hours, 20 km resolution, averaged over 1°
- ◆ good spectral resolution (HIRS:19 IR,1VIS, MSU: 4 μw)



based on:

radiosonde measurements



4A radiative transfer

$R_{\text{clr}}, R_{\text{cld}}(\lambda_i, p_k, \theta)$

TOVS Initial Guess Retrieval dataset

3I Inversion

(*Chédin, Scott 1985*)

*temperature, water vapor profiles
+ cloud properties*

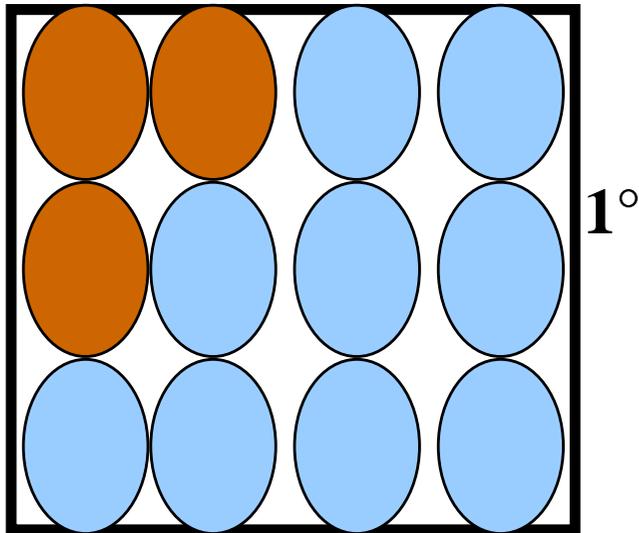
3I Cloud detection

◆ at HIRS resolution

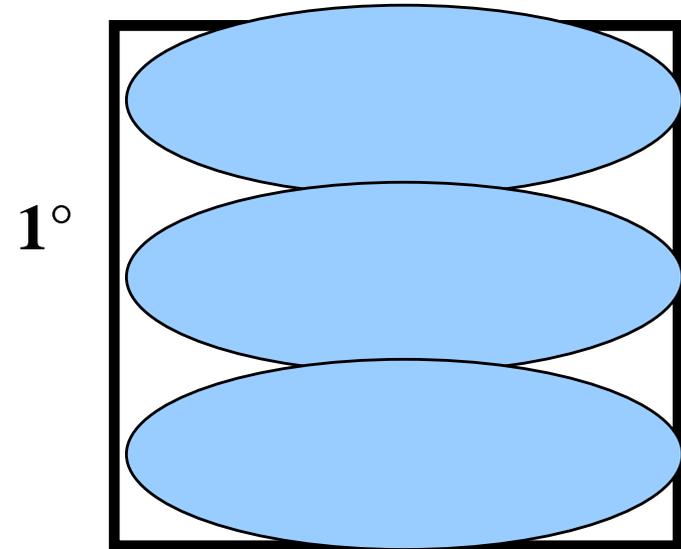
MSU predicts clear sky T_B

7/8 threshold tests

nadir
 1°

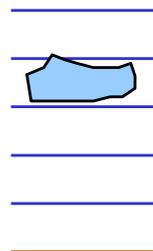


$\theta_v = 58^\circ$
 1°



3I Cloud property retrieval

◆ over averaged cloudy pixels

 4 channels in 14μm CO₂-band + 11 μm
max weights: 400-900 hPa surface

coherence of effective cloud amount

$$N\varepsilon(p_{cld}) \cong N\varepsilon(p_k, \lambda_i) = \frac{R_m(\lambda_i) - R_{clr}(\lambda_i)}{R_{cld}(p_k, \lambda_i) - R_{clr}(\lambda_i)} \quad \text{for } i=4,8$$

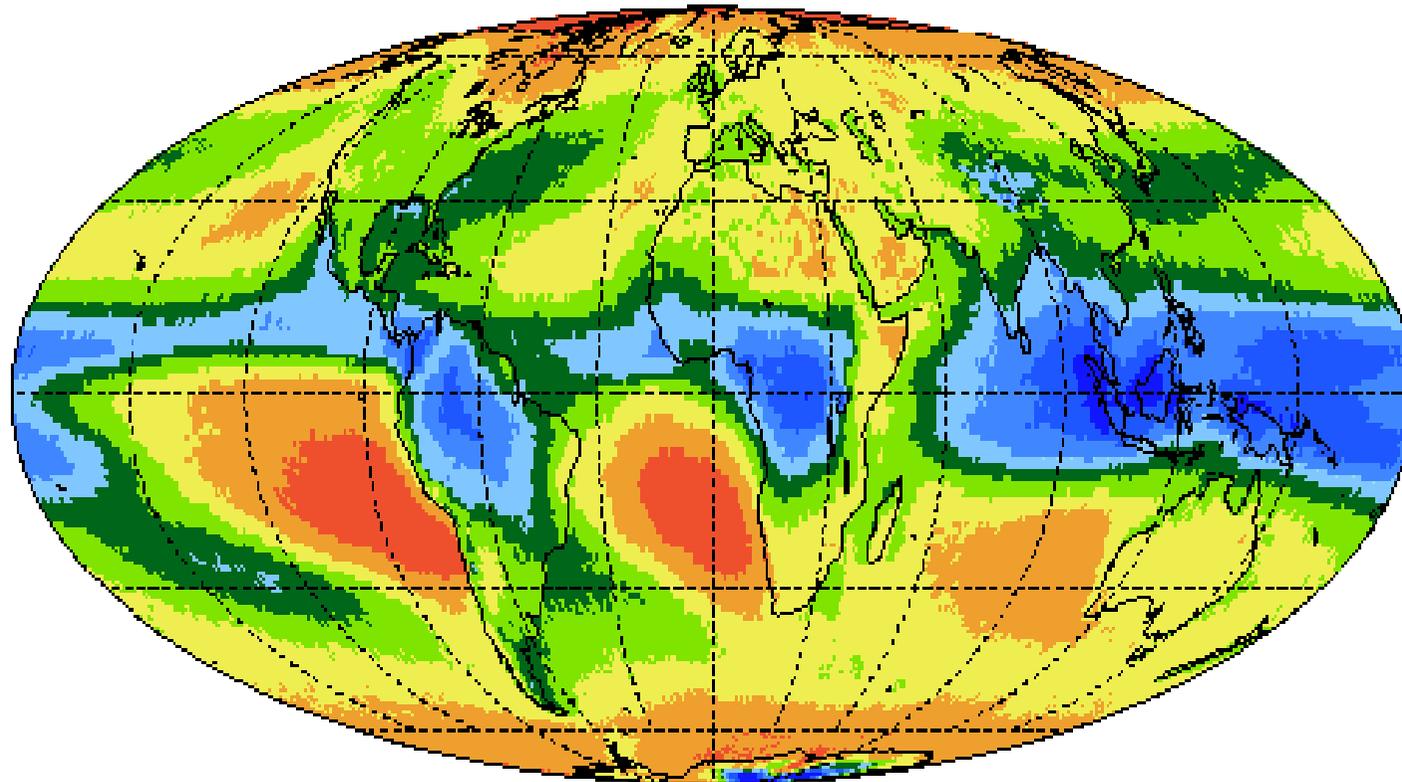
minimize weighted $\chi_w^2(p_k) \Rightarrow N\varepsilon, p_{cld}$

empirical weights \Leftrightarrow T profile uncertainty on $R_{cld} - R_{clr}$
(Stubenrauch et al., J. Climate 1999)

reliable cirrus clouds

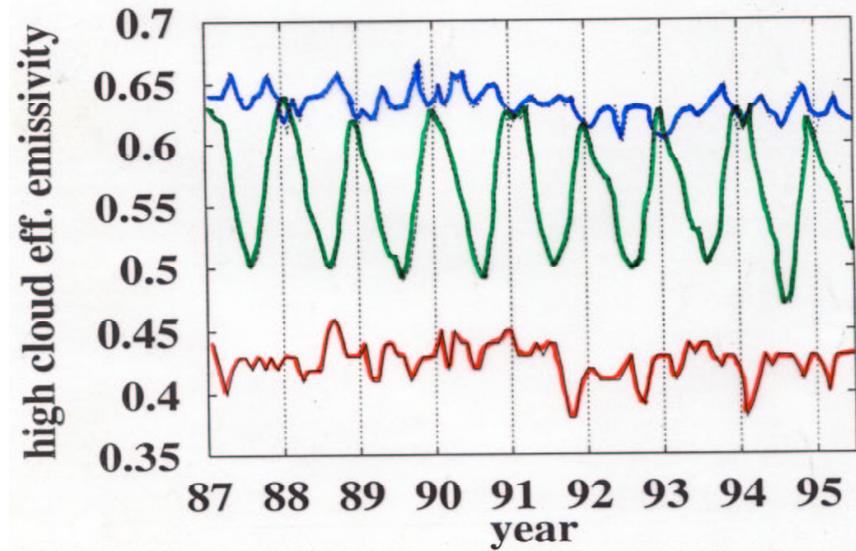
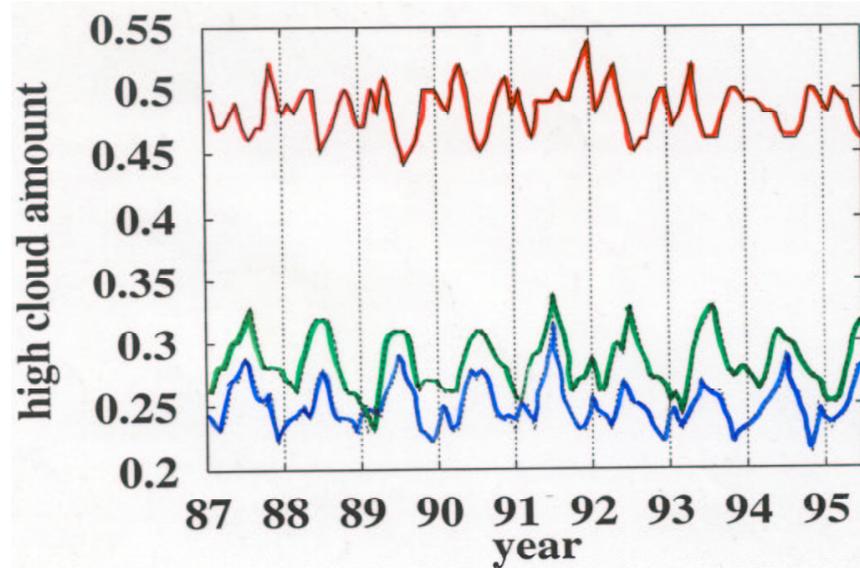
TOVS Path-B high cloud frequency

P_{cld} < 440 hPa, 1987-1991

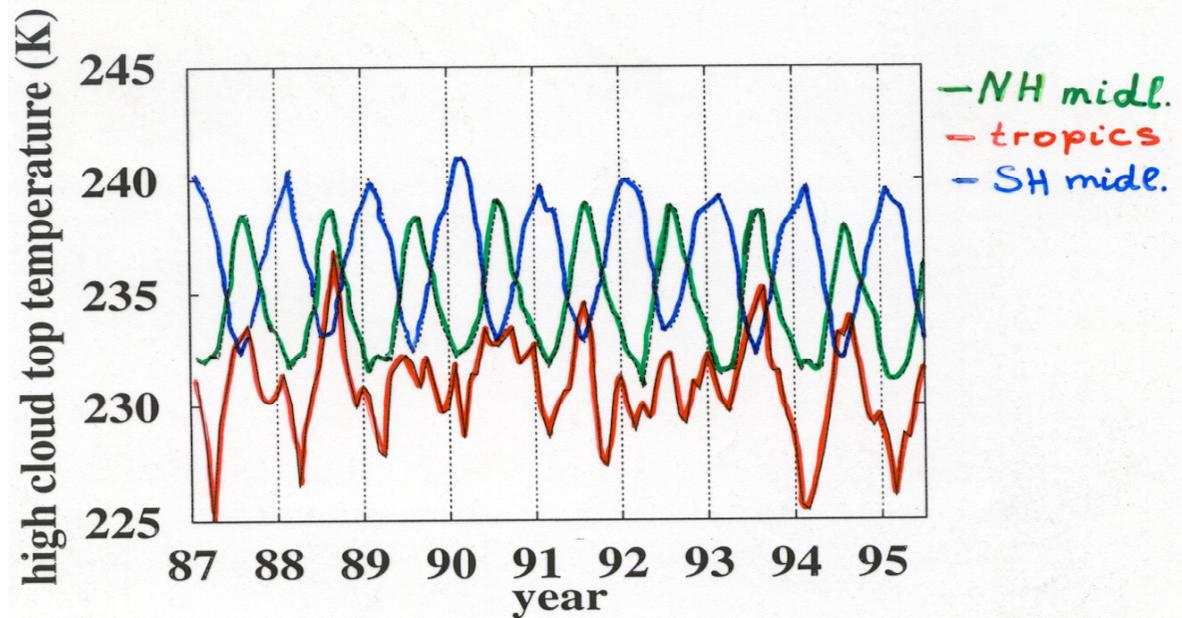


%

Time series of TOVS Path B high cloud properties

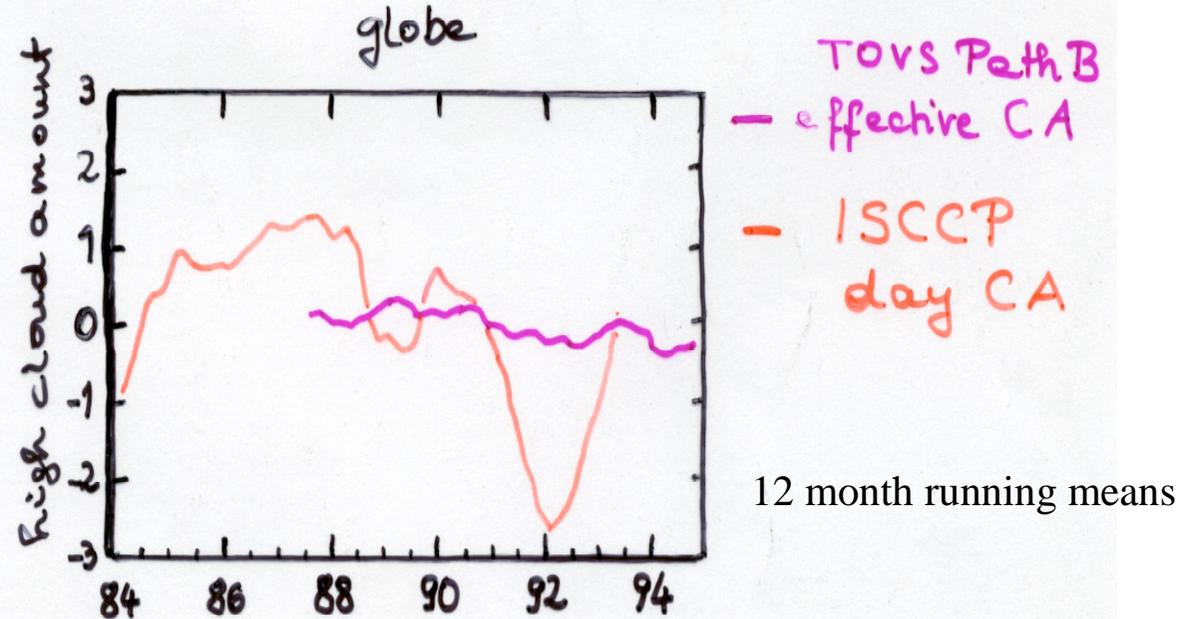


Global $CA_{\text{high}} \sim 0.3$



Study of high cloud amount evolution:

correlation with cosmic rays, volcanic eruptions?



Stubenrauch + Eddouia, CERN Proceedings 2001-007

Workshop on Ion-Aerosol-Cloud Interactions

◆ Satellite observations:

unique possibility to survey cld properties over long period

ISCCP

TOVS

1983-1995: $\langle CA \rangle \sim 67\%$ $\langle \text{eff. CA} \rangle \sim 53\%$

stable within 2% over globe

Svensmark's 'cosmic ray intensity - CA correlation' analysis could not be confirmed!

1991 Pinatubo eruption

aerosols -> slight overestimation of ISCCP τ

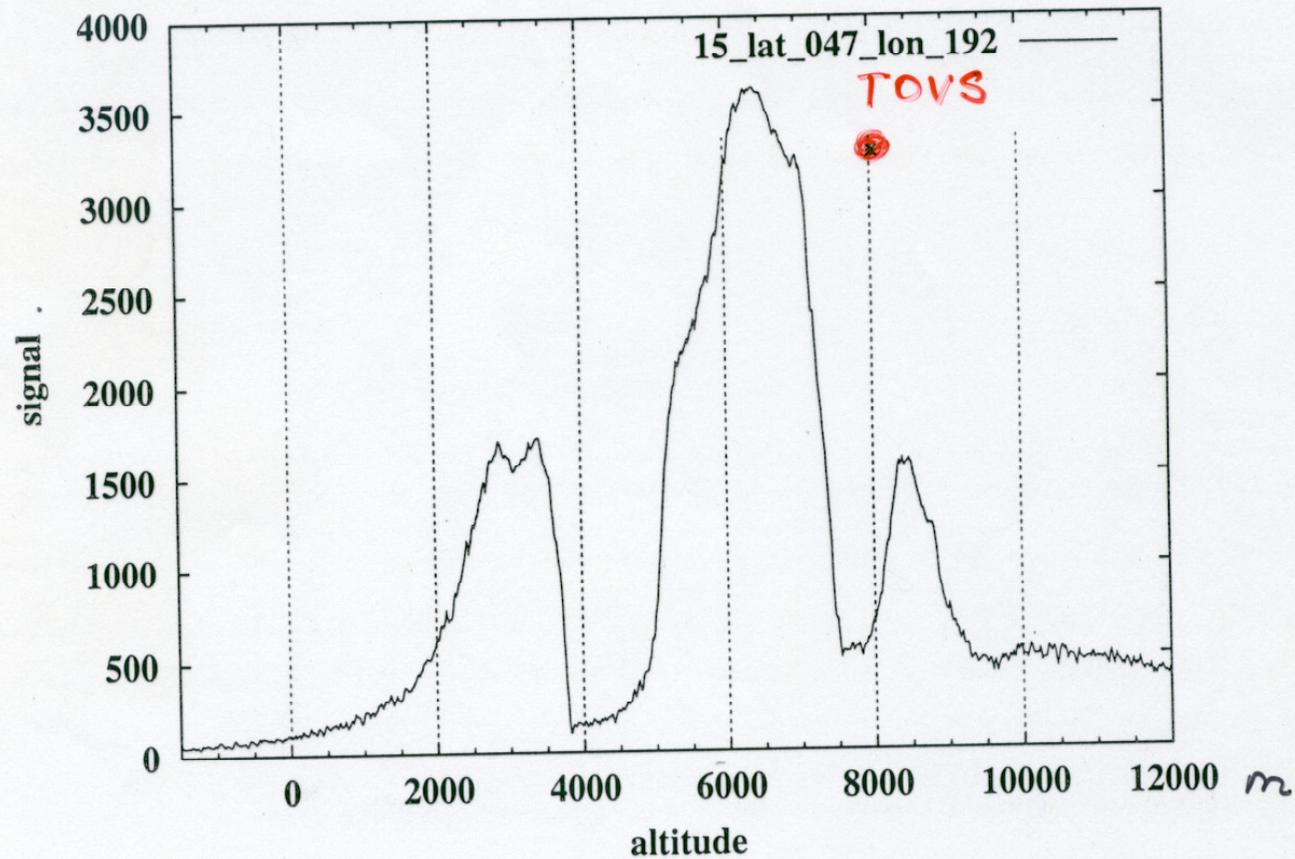
=> slight underestimation of ISCCP high CA
(4.5% in tropics)

TOVS high eff. CA stable (IR)

TOVS - LITE comparison 9/94

F. Eddounia

$\Delta t \sim \pm 3$ hours : ~ 2000 evts

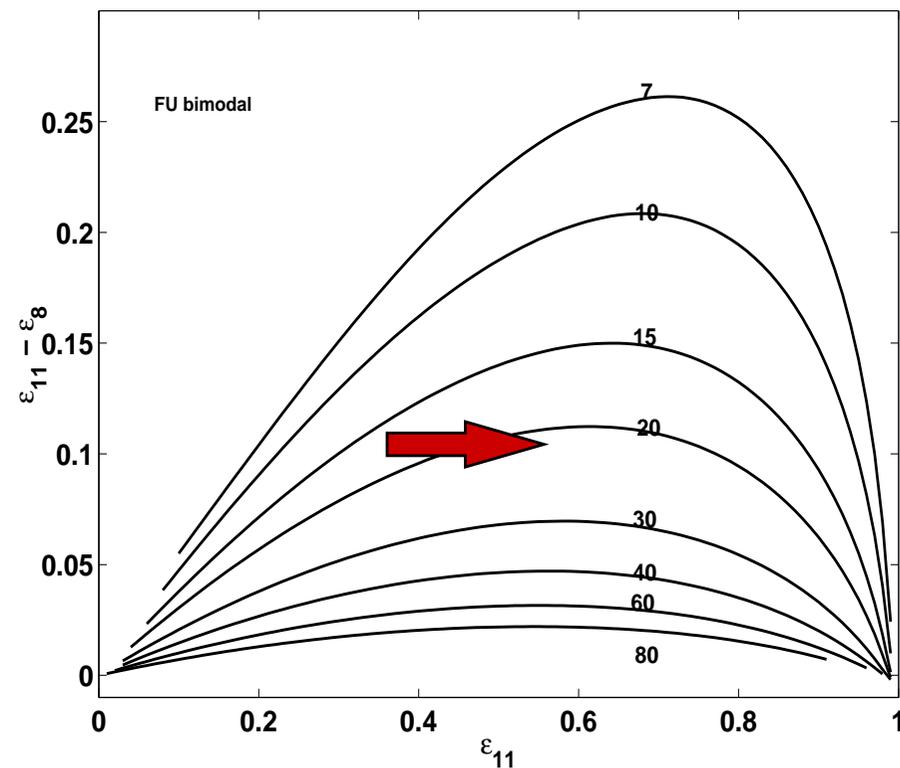
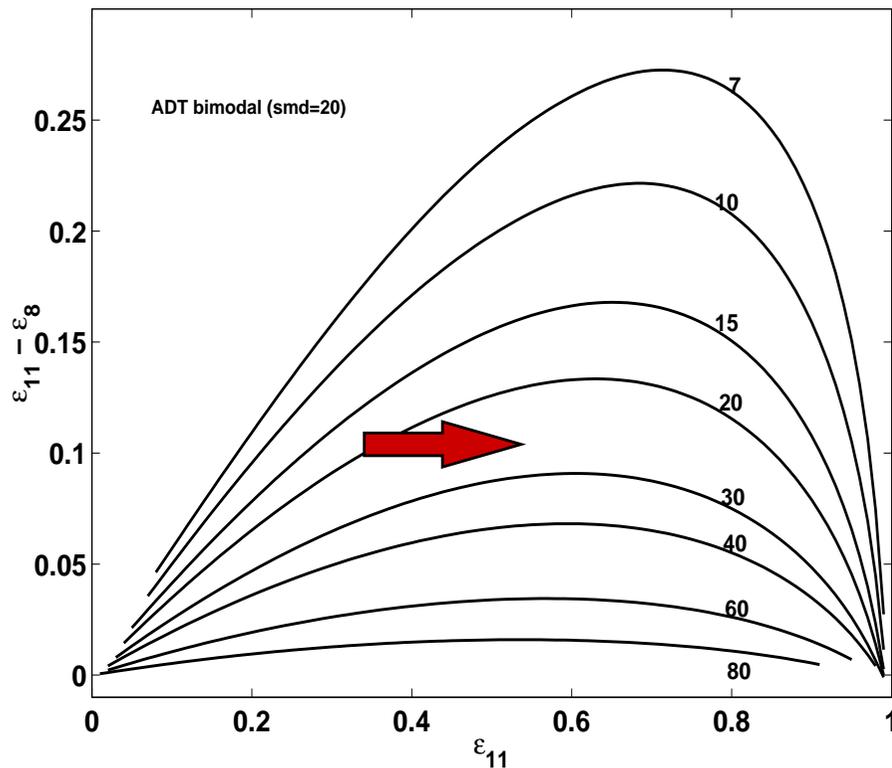


TOVS cirrus ice crystal size retrieval

based on spectral difference of cirrus emissivities

◆ ice crystal single scattering properties $Q_{\text{abs}}, \omega_0, g$
aggregates: modified ADA (*Mitchell 1996*) hexagonal columns: FDTD (*Fu 1998*)

radiative transfer based on discrete ordinate method (*Streamer, J. Key*)



Simulation of look-up tables

- ◆ ADA: $Q_{\text{abs}} = 1 - \exp\left(\frac{-4\pi n_i D_e}{\lambda}\right)$, ω_0 g from FDTD
- ◆ $D_e = 2r_e^{\text{VP}} = 2 \frac{\int \frac{3V}{4\pi} n(r) dr}{\int \frac{P}{\pi} n(r) dr} = \frac{3}{2} \frac{IWC}{\rho_i P}$
- ◆ Bimodal- Γ size distribution $N(D) = N_{0s} e^{-\lambda_s D} + N_{0l} D^{\nu_l} e^{-\lambda_l D}$
- ◆ $D_e(\epsilon_{11\mu\text{m}}, \epsilon_{8\mu\text{m}})$, for 8 surfaces and 4 θ_v 's
SARB surface emissivities
- ◆ for $0 < \epsilon_{11\mu\text{m}} - \epsilon_{8\mu\text{m}} < 0.01 \rightarrow D_e(\epsilon_i) = D_e(\epsilon_{i-1}) + 2.5\mu\text{m}$

Sensitivity study on ice crystal size retrieval

Homogeneous cloud, $z_{\text{cld}} = 10 \text{ km}$, $z = 1 \text{ km}$, $T_{\text{surf}} = 300 \text{ K}$,
 polycrystals, bimodal- Γ size distribution

		15 μm D_e 60 μm	
crystal shape	spheres	+ 15%	+30%
	hexagonal columns	+ 15%	- 10%
size distribution	tropical	- 20%	- 28%
	midlatitude	+ 5%	- 20%

		thin clouds		thick clouds	
		15 μm D_e 60 μm		15 μm D_e 60 μm	
z_{cld}	> 4km	- 4%	- 2%	- 6%	- 2%
T_{surf}	< 15 K	+ 3%	+ 2%	+ 5%	+ 2%
Δz	> 1 km	+ 4%	+ 6%	+ 8%	+ 10%
horizontal heterogeneity	90% cloud cover	+ 7%	+ 2%	+ 27%	+12%
vertical heterogeneity	IWC(2) = 1.5 IWC(1)	+ 0.5%	+ 1.0%	+1.2%	+ 0.7%
	$D_e(2) = 2 D_e(1)$	-15%	- 20%	-15%	- 20%
+ water cloud	$r_e=10\mu\text{m}$, LWC=0.03g/m ³	+ 10%	+ 15%	+ 9%	+ 9%
	$r_e=7\mu\text{m}$, LWC=0.20g/m ³	+ 25%	+ 25%	+ 15%	+ 15%

TOVS cirrus ice crystal size retrieval

(*Stubenrauch et al., JGR 1999, CIRAMOS*)

1°x1° overcast 3I high clouds ($p_{\text{cld}} < 440$ hPa)

$T_B^{\text{meas}}(8\mu\text{m}), T_B^{\text{meas}}(11\mu\text{m}), T_{\text{cld}}, T_{\text{surf}}$
closest TIGR atmosph. H₂O/T profile, SARB surf. emissivities

↓ **3R radiative transfer**

cirrus effective emissivities

$$\varepsilon(\lambda, \theta_v) = \frac{B(T_B^m(\lambda, \theta_v)) - B(T_{\text{surf}}(\lambda, \theta_v))}{B(T_{\text{cld}}(\lambda, \theta_v)) - B(T_{\text{surf}}(\lambda, \theta_v))}$$

$\varepsilon(11\mu\text{m}) - \varepsilon(8\mu\text{m})$ < - >

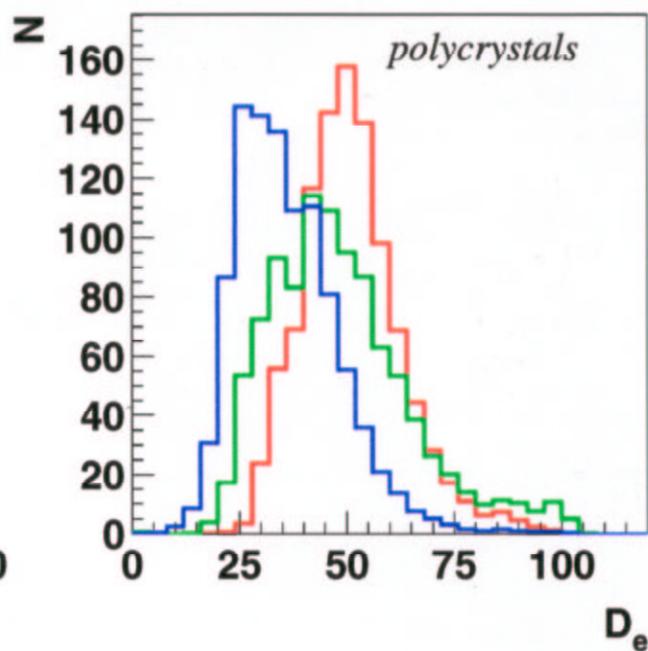
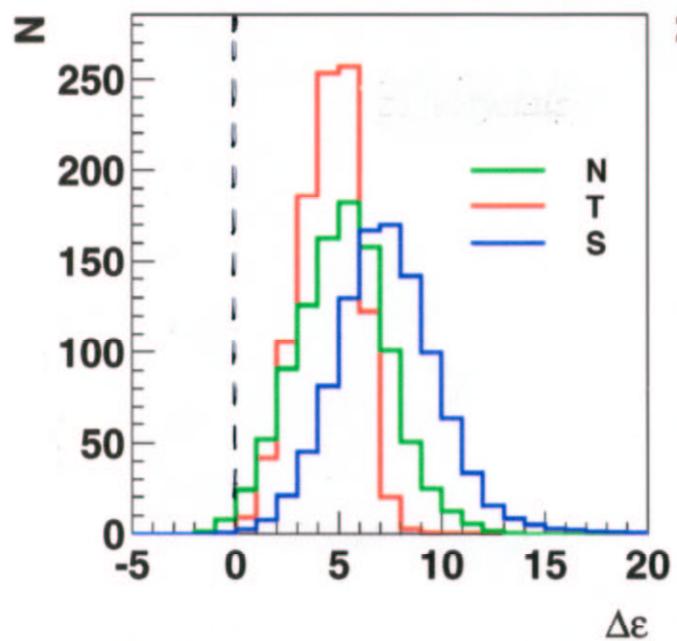
↓
D_{eff}

simulated
look-up tables

TOVS cirrus ice crystal size retrieval

- ◆ NOAA10 satellite observations: 1/1987 - 9/1991
(NOAA12: 8 μm channel problem at low temperatures!)
- ◆ $0.3 < \varepsilon(11\mu\text{m}) < 0.85$, sensitivity: $D_e \leq 90\mu\text{m}$
 $0.7 < \tau_{VIS} < 3.2$
- ◆ $\theta_v < 25^\circ$
- ◆ D_e uncertainties increase with D_e :
 - calculate median values of distributions

NH summer / ocean

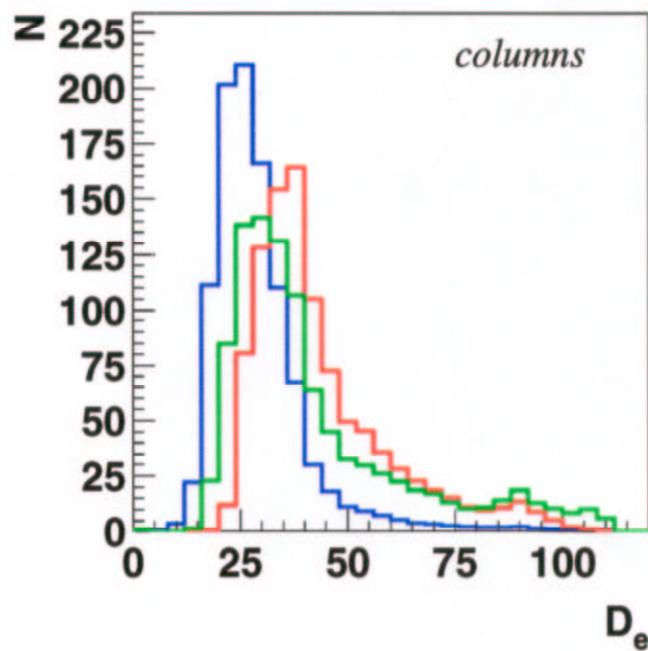
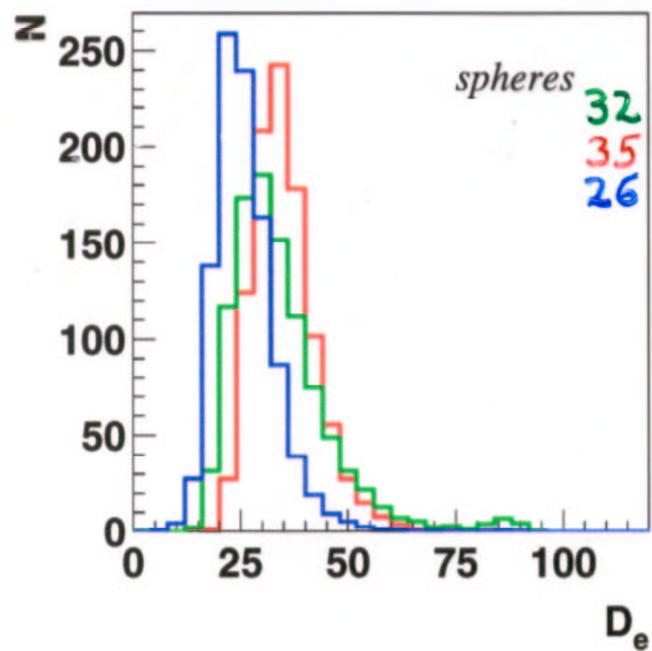


median
 D_e (μm):

47

51

35

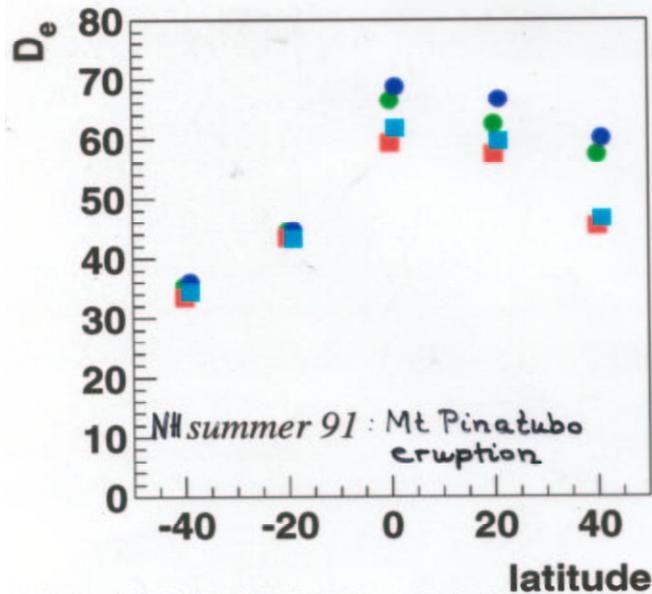
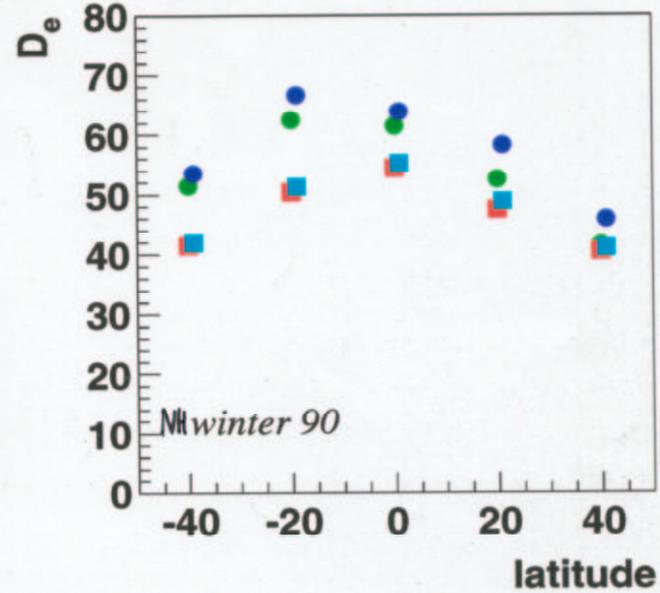
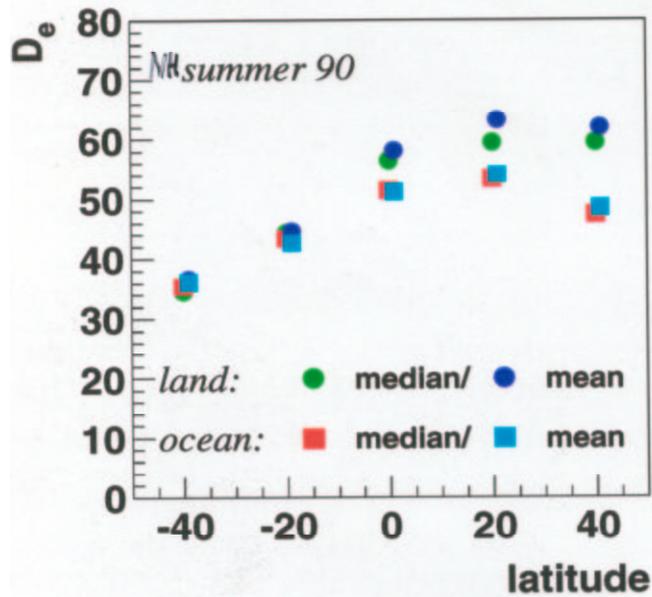


36

39

27

D_e as function of latitude



midlatitudes:

$$D_e(\text{sum}) > D_e(\text{win})$$

$$D_e(\text{tropics}) > D_e(\text{midlat.})$$

Conclusions and Outlook

- ◆ Retrieval of *Cirrus mean effective ice crystal size* feasible at a *global scale* using TOVS Path-B satellite data

- ◆ *aggregate assumption*: D_e between 35 and 65 μm

$$D_e(\text{tropics}) > D_e(\text{midlatitude}),$$

$$D_e(\text{summer}) > D_e(\text{winter})$$

$$D_e = f(\text{H}_2\text{O}, \varepsilon, \text{dynamics})$$

- ◆ *hex. column assumption*: D_e about 8 to 11 μm smaller

- ◆ recent NOAA satellites: no 8 μm channel \rightarrow 4 μm channel

- ◆ study other theoretical approaches (*Yang, Mishchenko, Baran*)

- ◆ check coherence with combined LW/SW fluxes (ScaRaB)

Cirrus microphysical properties and their effect on RAdiation:



CIRAMOSA



Kickoff: 9/2/2001

**survey and integration into climate MOdels
using combined SATellite observations**

Coordinator: Claudia Stubenrauch (C.N.R.S.-LMD)

Web-site: <http://www.lmd.polytechnique.fr/CIRAMOSA/Welcome.html>

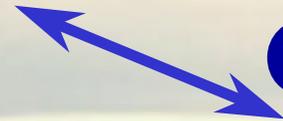
CIRAMOSA Objectives

- ◆ long-term survey of cirrus physical and microphysical properties (1987-1995)
- ◆ correlations between cirrus properties and the state of the atmosphere
- ◆ study their effect on the reflection and absorption of solar and thermal radiation
- ◆ improvement of cirrus radiative transfer in GCMs

Observations: TOVS, ATSR, POLDER, in-situ



state of atmosphere



Cirrus

macrophysics



microphysics

correlations



GCM radiative transfer model



ScaRaB observed - simulated

TOA LW, SW fluxes