

Survey of cirrus and atmospheric properties from TOVS Path-B

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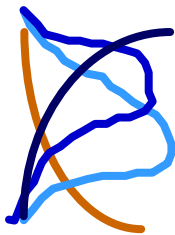
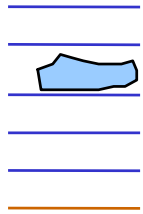


- ◆ **TOVS Path-B Dataset
& average cloud properties**
- ◆ **Variability of cirrus properties**
- ◆ **Reanalysis**
- ◆ **Upper trop. humidity + evolution of contrails**
(collaboration with U. Schumann, DLR)

TOVS Path-B climatology:..., 1987- 1995, ...

Scott et al., BAMS, 1999

MSU+HIRS $R_m(\lambda_i, \theta)$ along H_2O , CO_2 absorption bands



3I Inversion

(Chédin, Scott 1985)

- atmospheric temperature (9 layers, ≥ 10 hPa), water vapor (5 layers, ≥ 100 hPa)
- effective cloud amount (ECA), cloud top pressure (Stubenrauch et al. 1999)
- D_e , IWP of cirrus (CIRAMOS, poster Eddounia et al.)
- horizontal extent of high clouds (G. Rädcl)
- upper tropospheric relative humidity

3I based on: controlled use of a priori information: radiosondes - radiative transfer

TIGR dataset: $T(p_k)$, $H_2O(p_k)$, $T_s - R_{clr}(\lambda_i, \theta)$, $R_{cld}(\lambda_i, p_k, \theta)$

Thermodynamic Initial Guess Retrieval

Average cloud properties

8 years (1987-1995) TOVS Path-B / ISCCP

Cloud type amounts (%)	global		ocean		land	
all	73	65	74	71	69	58
Deep convection	2.4	2.8	1.9	2.8	3.5	2.7
Cirrus	27.3	19.1	26.9	18.0	27.8	21.7
Mid-level	12.1	18.5	10.3	18.4	16.6	18.5
Low-level	30.9	26.7	35.1	30.6	20.5	17.7

~ 70 % cloud amount: *more over ocean than over land*

~ 30% low clouds: *more over ocean than over land*

~30% high clouds: *same over ocean and land*

Vertical sounders more sensitive to Cirrus clouds (8% more than ISCCP)

Observed Global Climate, Chap. 'Clouds', June 2005, Springer

Average regional high cloud properties

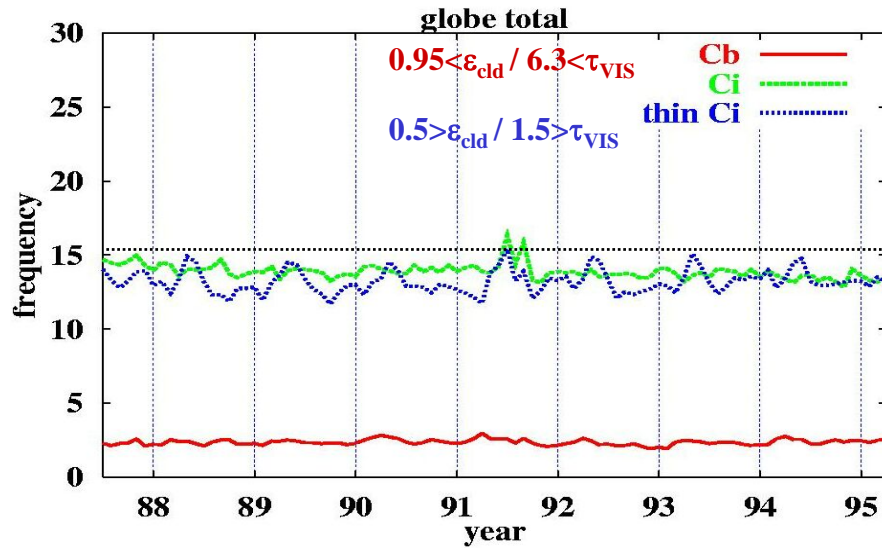
8 year (1987-1995) TOVS Path-B / ISCCP

Cloud type amounts (%)	NH midlat.		tropics		SH midlat.	
Deep convection	3.0	3.3	2.5	3.5	2.4	3.0
Cirrus	24.7	20.3	44.8	24.9	21.8	16.5

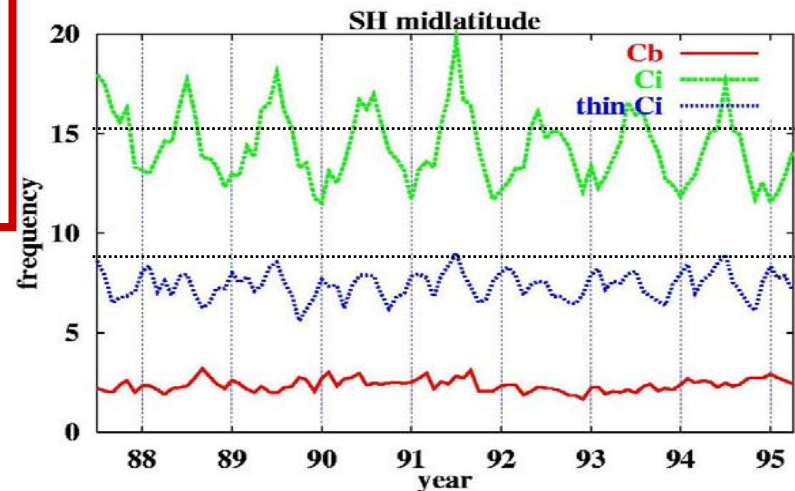
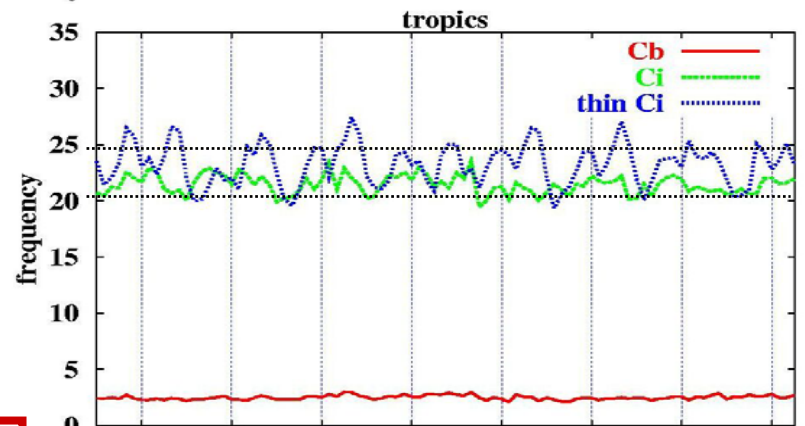
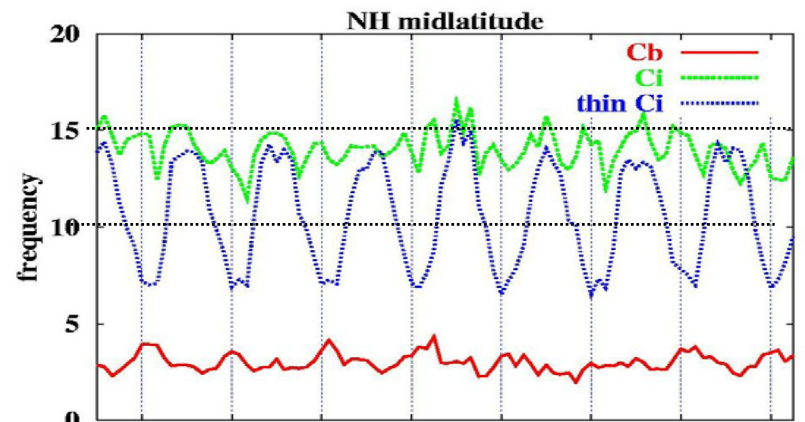
- only 3% convection
- IR vertical sounders:
identify Ci day + night
more sensitive to Ci: *midlat.* +4%
tropics +20%

Time series of TOVS Path B high cloud frequencies

NOAA10/12 7h30 AM&PM

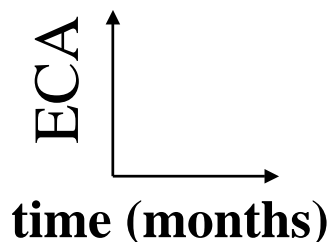
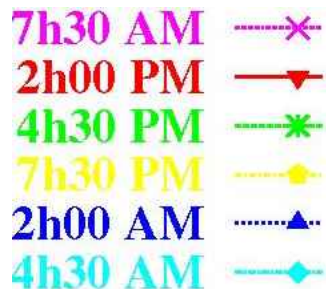


- stable over 8 years within 2%
- NH mid: strong seasonal cycle of thin Ci
- SH mid: seasonal cycle of Ci

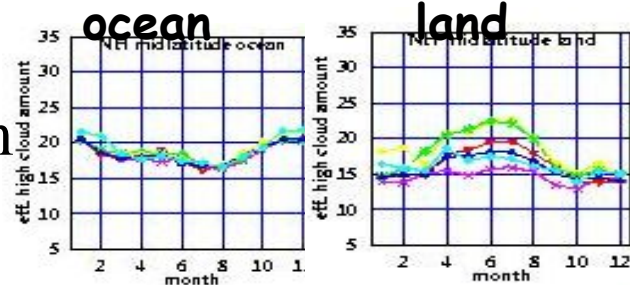


Seasonal & diurnal variations of effective high cloud amount

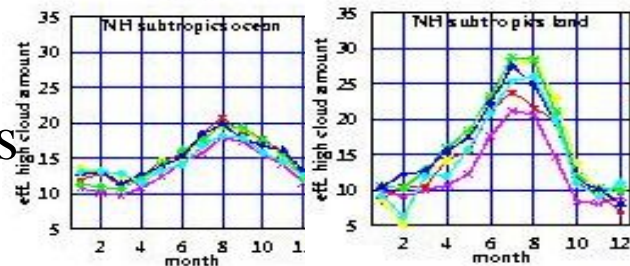
NOAA10/12 7h30 AM&PM,
 NOAA11 2h00 AM&PM(1989-90)
 NOAA11 4h30 AM&PM(1994-95)



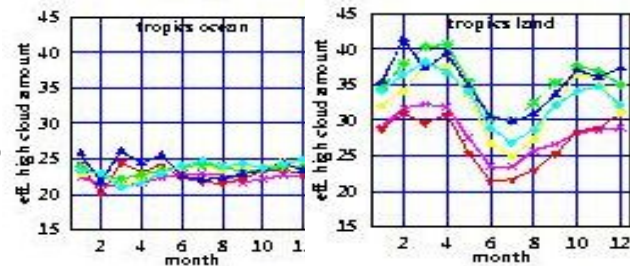
NHm



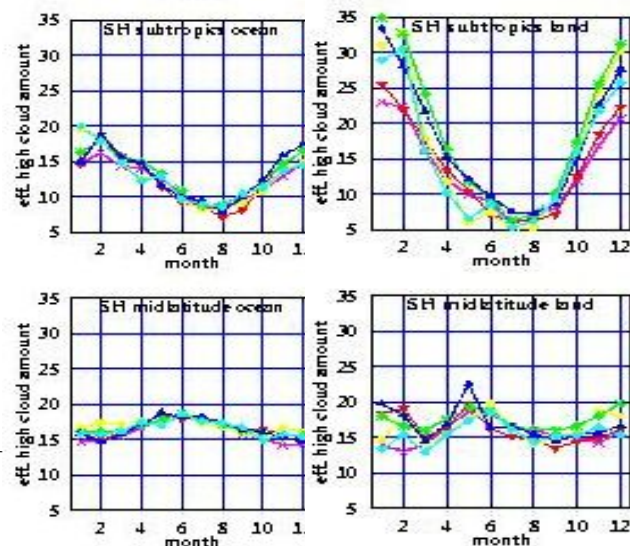
NHs



trp



SHm

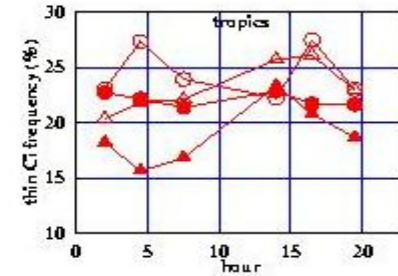
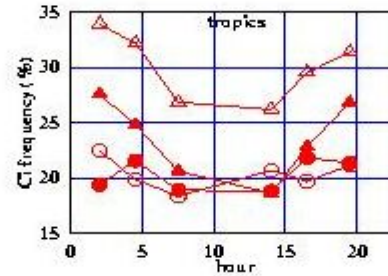
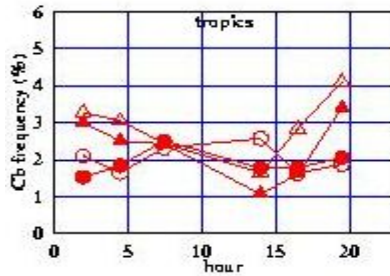


- cycles stronger over land than over ocean
- seasonal cycle strongest in subtropics: *ITCZ shift*
- tropics: diurnal variability stronger than seasonal
- NH land: seasonal cycle strongest in afternoon

diurnal cycle of high cloud type frequencies

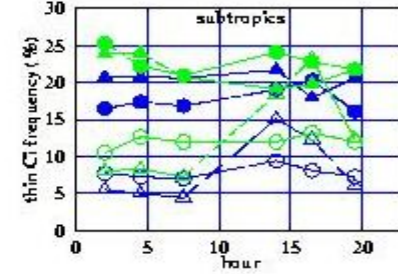
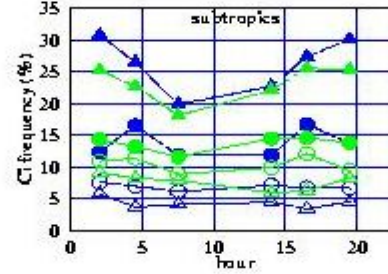
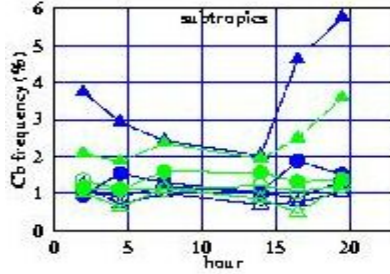
NOAA10/12 7h30 AM&PM, NOAA11 2h00 AM&PM(1989-90) NOAA11 4h30 AM&PM(1994-95)

tropics

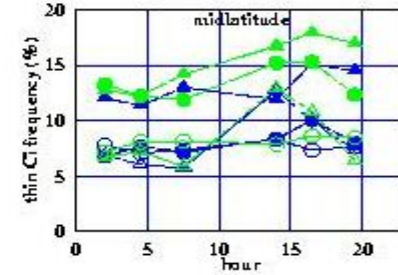
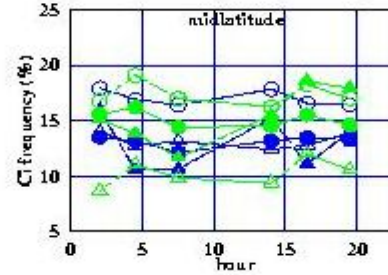
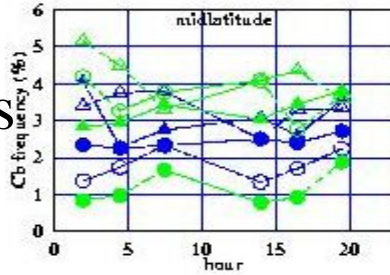


△ land
• Sea
sum
win

subtropics
NH, SH



midlatitudes
NH, SH



- strongest diurnal cycles over land in tropics and in summer
- convection strongest in evening
- more cirrus during night and more thin cirrus in afternoon

❖ Reanalysis of entire TOVS data at LMD:

Improved TIGR database

- extension from 1763 to 2311 atmospheric profiles (in tropics)
- new 4A model (spectroscopy, continuum) for T_B computation
- new surface emissivities (*FASTEM-2, S. English*)
- O_3 profiles from UGAMP climatology
- new extrapolation of T and H_2O towards stratosphere from ATMOS

3I Inversion

- scheme adapted to new TIGR
- new neural network inversion for $H_2O(p)$ and T_{surf}

Adjustment constants (« deltacs »)

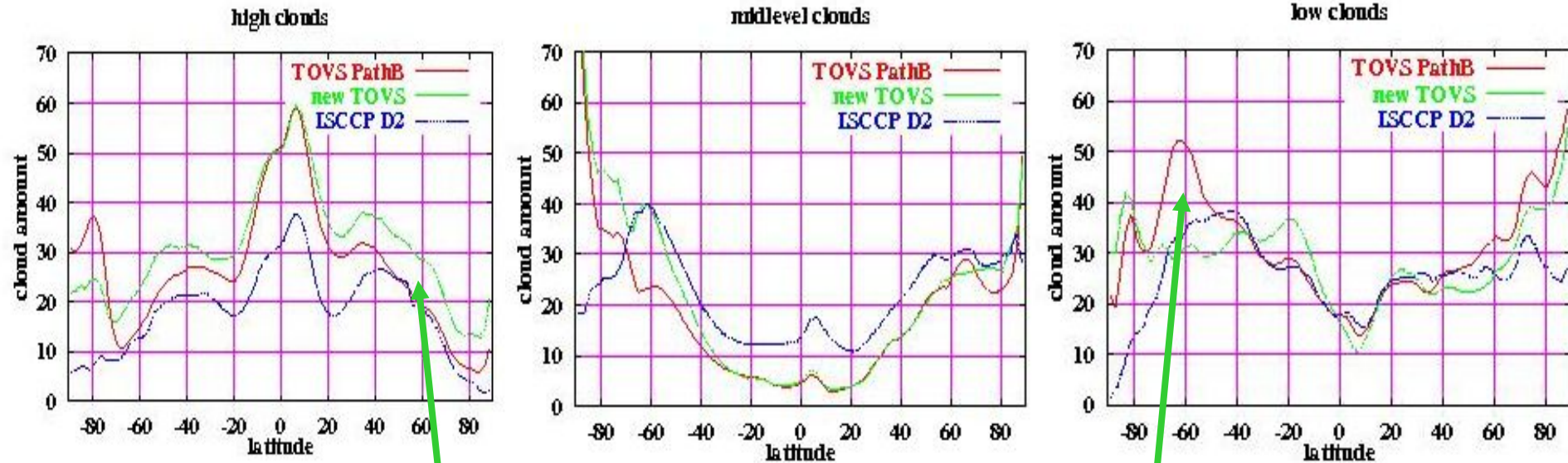
1987-1995: *DSD5* radiosonde-TOVS dataset from NOAA : clear /cloudy

1979-2004: *ERA-40* « cleaned » radiosonde-TOVS data collocated,
clear sky determination

(radiosonde temperatures during day not corrected in stratosphere)

TOVS Reanalysis: 1 year of cloud data (1990)

preliminary



tropics: nearly no change

subtropics- midlat.: slightly more high clouds

SH midlat. ocean: up to 20% less low clouds (closer to ISCCP)

Determination of TOVS relative humidity (per layer)

- ◆ TOVS Path-B precipitable water column: 300 - 100 hPa

$$W = \int_p^{p_0} q_s \frac{dp}{g\rho}$$

=> rel. humidity per layer: $RH^{\text{ice}}(\Delta p) = g\rho W / \int q_{\text{sat}}^{\text{ice}}(p) dp$

- 1) 3I retrieved atmospheric T profile (30 levels) -> $e_{\text{sat}}^{\text{ice}}(T)$

$$\ln(e_{\text{sat}}) = \frac{a_1}{T} + a_2 + a_3 T + a_4 T^2 + a_5 \ln(T) \quad (\text{Sonntag, 1990})$$

at 86, 106, 131, 162, 200, 223, 248, 276 and 307 hPa

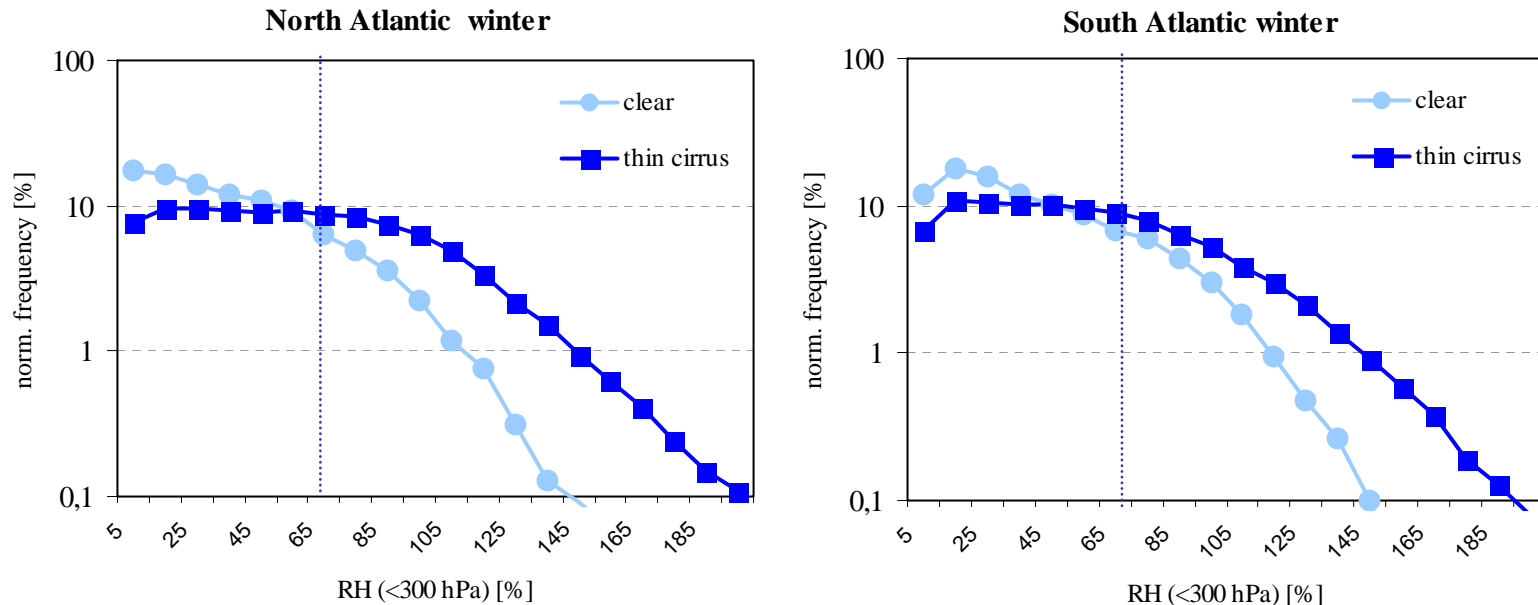
- 2) integrate $q_{\text{sat}}^{\text{ice}}$ over column (in steps of 1hPa) :

$$\int q_{\text{sat}}^{\text{ice}}(p) dp = \sum 0.622 \frac{e_{\text{sat}}^{\text{ice}}(T(p))}{p - (1 - 0.622)e_{\text{sat}}^{\text{ice}}(T(p))}$$

Relative humidity distributions

in case of clear sky and thin cirrus ($N\varepsilon < 0.5$)

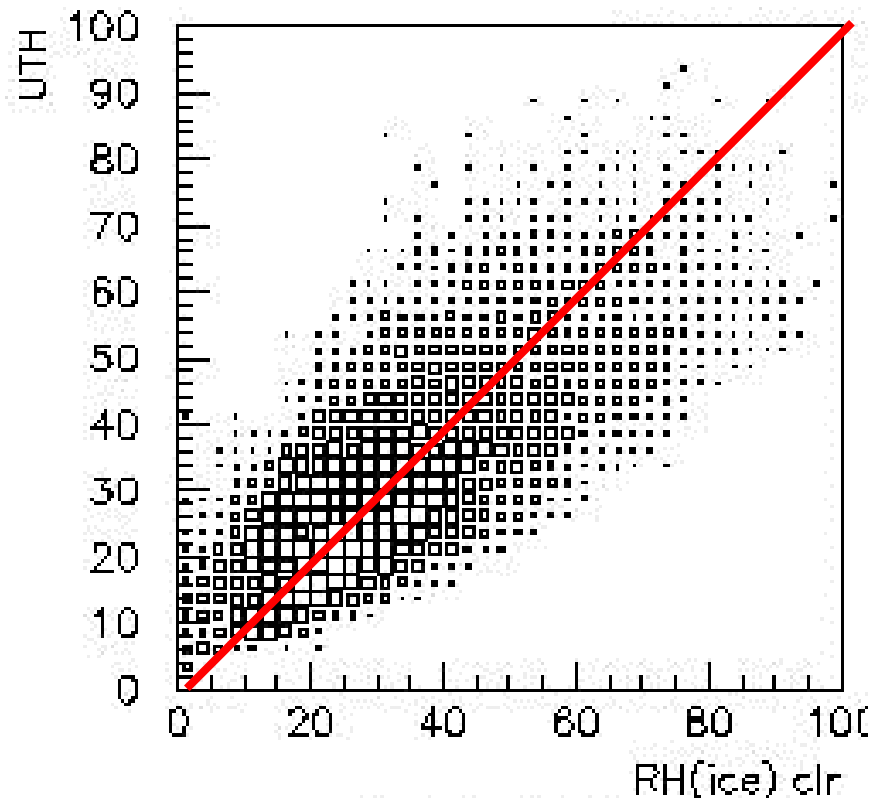
over 8 years



- ◆ Thin cirrus have broader RH distributions than clear sky
- ◆ However, clear sky can also be ice saturated (*in agreement with Gierens et al. 1999*)

3I relative humidity (300-500hPa) - UTH *in case of clear sky and thin cirrus*

$$UTH = \frac{\exp(a_1 + a_2 T_{HIRS12})}{a_3 + a_4 T_{HIRS6}} \quad (\text{Bates})$$





Evolution of contrails from TOVS



C. Stubenrauch, U. Schumann



◆ **contrails: cold and moist ambient air, $RH > U^*(T)$**

◆ **critical rel. humidity** TOVS: integrate over layer $\int [G \cdot (T - T_{lmax}) + e_{sat}^{liq}(T_{lmax})] dp$
 $T_{lmax} = 230.8K$ $e_{sat}^{liq}(T_{lmax}) = 20.6hPa$
 Kerosen: $G = 1.5$
 $U^*(\Delta p) = \frac{\int_{100-300hPa} [G \cdot (T - T_{lmax}) + e_{sat}^{liq}(T_{lmax})] dp}{\int_{100-300hPa} e_{sat}^{liq}(T) dp}$

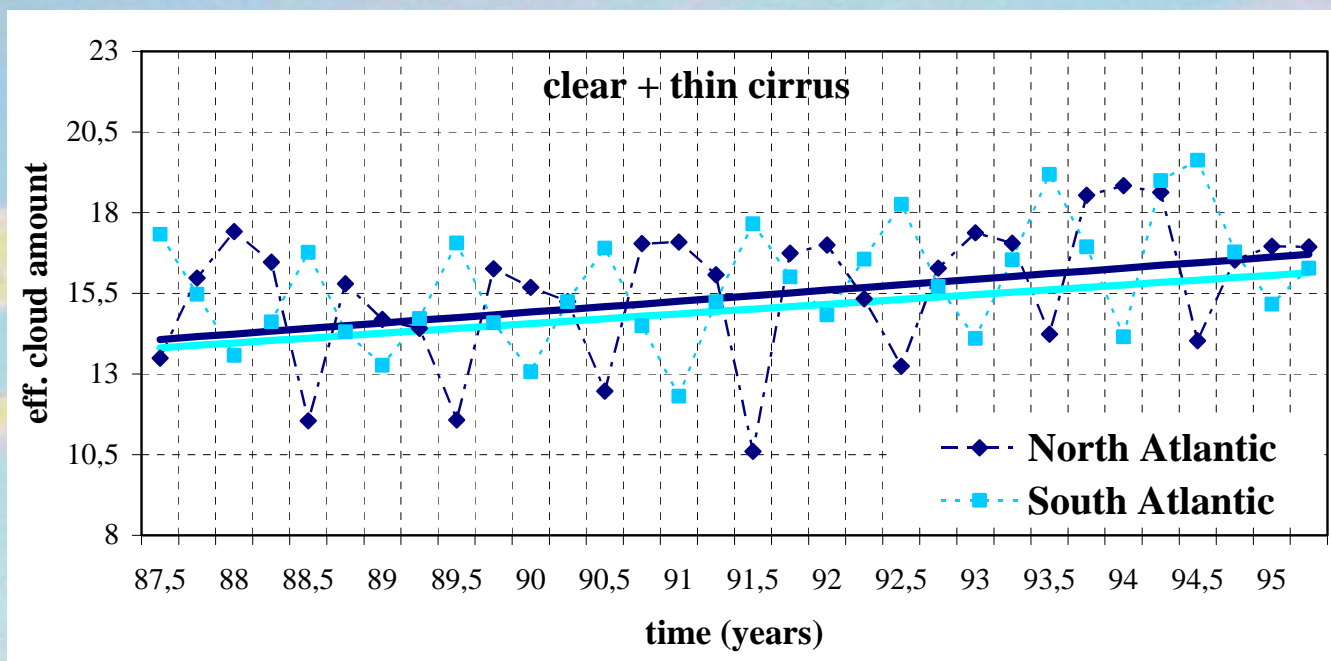
◆ *Sausen et al. 1997*: **potential contrail if $U_{ci} > RH > U^*U_{ci}$**

◆ **separate situations: $RH^{ice}(\Delta p) > 0.7$ cirrus**

$RH^{ice}(\Delta p) < 0.7$ & $RH^{liq}(\Delta p) > 0.4U^*(\Delta p)$ potent. contrail

$RH^{ice}(\Delta p) < 0.7$ & $RH^{liq}(\Delta p) < 0.4U^*(\Delta p)$ clear

◆ **Difference in trends of effective high cloud amount between situations of potential contrails - cirrus and situations of potential contrails - all**

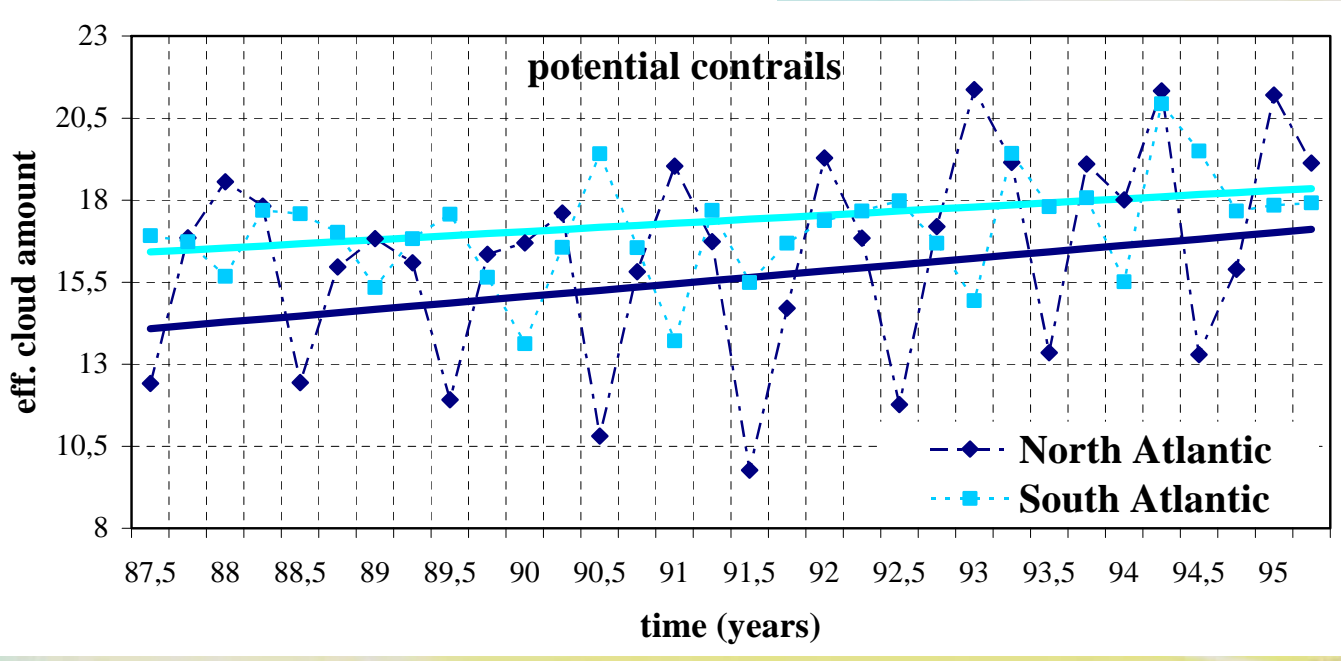


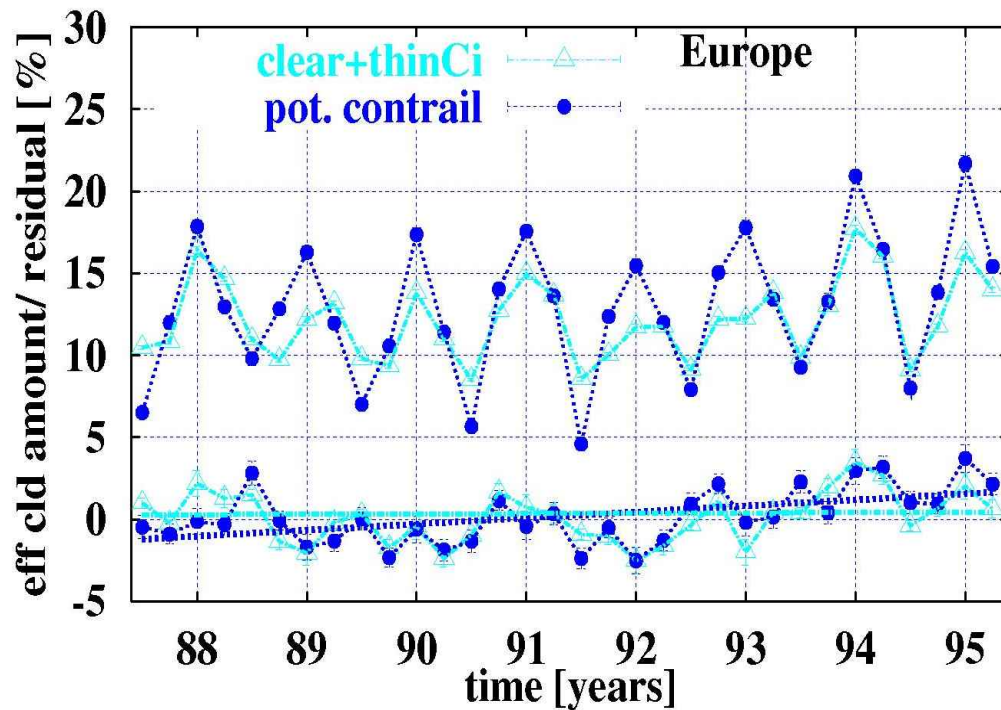
increase of thin Ci
in both hemispheres

stronger increase of Ci related to contrails in NH

Cirrus: $RH_{ice} > 70\%$

Pot. Contr.: $RH_{wat} > 0.4 U^*_{wat}$





ECA increase per decade for potential contrail situations

ECA trend difference (%/decade) between potential contrail and cirrus / all situations

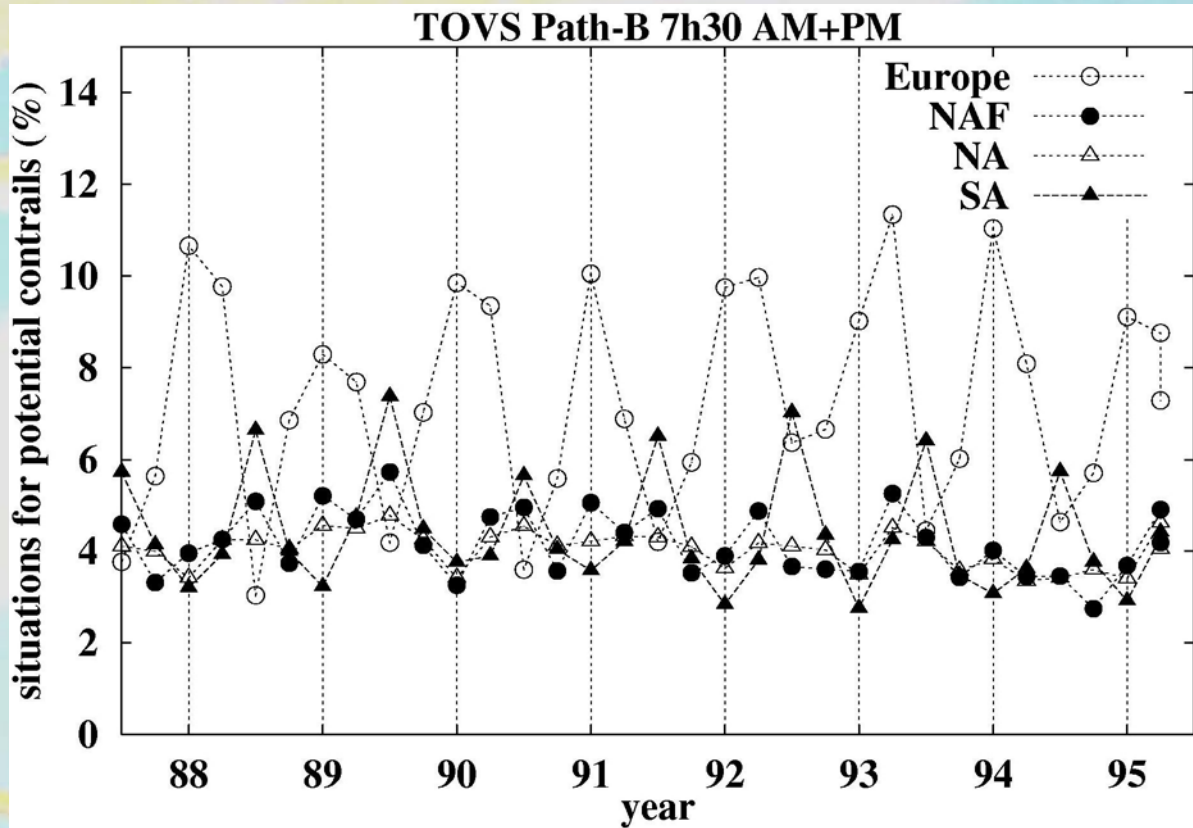
region/season	Europe	NAF	NA	SA
all pc-all/pc-ci	2.8 / 3.5	1.6 / 4.7	0.6/ -0.2	-1.6/ -0.9

uncertainty estimates 1.5%/decade (from threshold variations)

Stubenrauch + Schumann, GRL 2005 in revision

However:

Occurrence of pot. contrail situations is small: 5 - 10%



Overall effect: over Europe ~0.19% - 0.25% per decade
over NAF ~0.08% - 0.24% per decade

Satellite observations:

- ❖ unique possibility to study cloud properties over long period
30% high clouds, stable within 2% over globe
- ❖ seasonal and diurnal variabilities in high clouds:
 - strongest seasonal cycles over land in subtropics (*ITCZ shift*)
 - strongest diurnal cycles over land in tropics & summer
 - convection in evening, cirrus during night, thin cirrus in afternoon
- ❖ **TOVS reanalysis** : understand small changes in summer midlat. cloud properties
- ❖ **Contrail analysis**:
only by extracting situations of potential contrails
-> positive trend of εN in regions of high air traffic
in general small: $\sim 0.2\%$ per decade