



A Research of Four-dimension Variational Data Assimilation with ATOVS Clear Data

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Introduction



- More and more deduced atmospheric parameters from satellite data are used in numerical weather forecasting
- How to introduce radiances from satellite into numerical model
- 3D-VAR, 4D-VAR



- Definition of cost function

$$J(x) = J_b + J_s$$

$$J_b = \frac{1}{2} (X - X_b)^T B^{-1} (X - X_b)$$

$$J_s = \sum_i \sum_{ich} [F_i(x_{ich}) - y_{i,ich}^{obs}]^T (O + F)^{-1} [F_i(x_{ich}) - y_{i,ich}^{obs}]$$

- Where $X = (u, v, p', t, q)$ are all control variables



$F(x_{ich})$ is a fast transfer model to generate radiances

- In this test, RTTOV5 is used

$$L^{Clr}(\nu, \theta) = \tau_s(\nu, \theta) \varepsilon_s(\nu, \theta) B(\nu, T_s) + \int_{\tau_s}^1 B(\nu, T) d\tau + (1 - \varepsilon_s(\nu, \theta)) \tau_s^2(\nu, \theta) \int_{\tau_s}^1 \frac{B(\nu, T)}{\tau^2} d\tau$$

- Where $B(\nu, T)$ is Planck function
- $\tau_s(\nu, \theta)$ is transmittance from surface to space
- $\varepsilon_s(\nu, \theta)$ is the surface emissivity



the optical depth from each pressure level to space for each channel

$$d_{i,j} = d_{i,j-1} + Y_j \sum_{k=1}^K a_{i,j,k} X_{k,j}$$

- $a_{i,j,k}$ is regression coefficients
- Y_j and $X_{k,j}$ are prediction factors



Data

- Operational TOVS data of NSMC in East Asia every day from May to August 1998
- Climatic profiles are used to represent atmospheric state from 100 hPa to 0.1 hPa for temperature and from 300 hPa to 0.1 hPa for water vapor

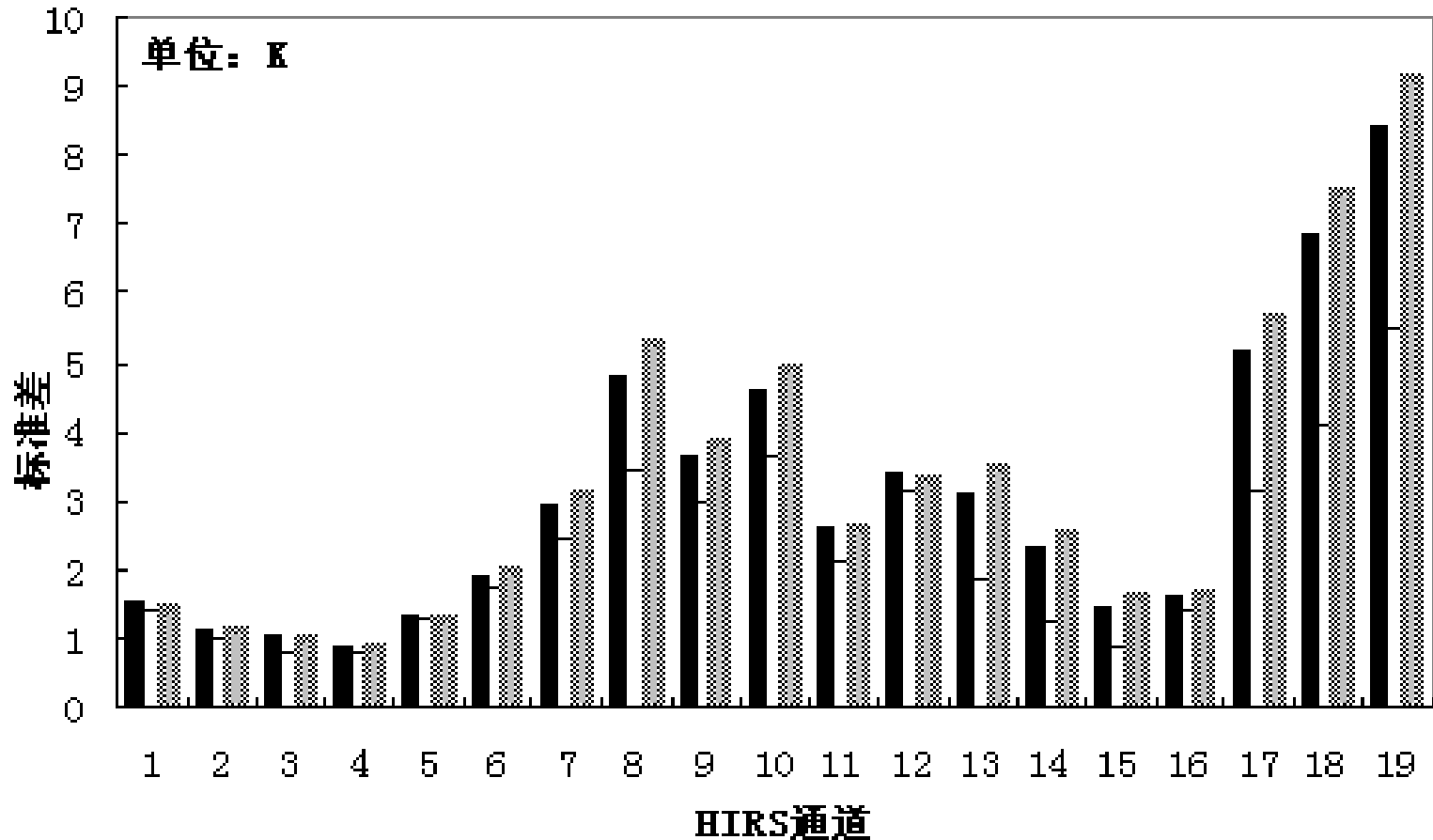


Confirmation of cloud-clear data

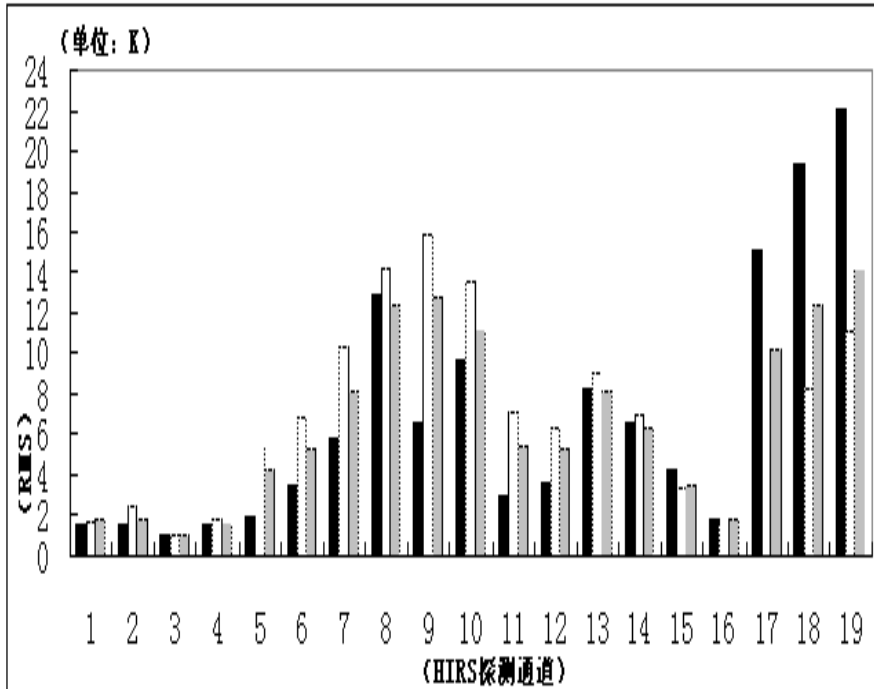
$$T_{skin} - Tb_{ch8} = \begin{cases} < 10 K & \textit{clear} \\ > 10 K & \textit{cloudy} \end{cases}$$

- T_{skin} is the surface skin temperature
- Tb_{ch10} is brightness temperature of HIRS channel 8

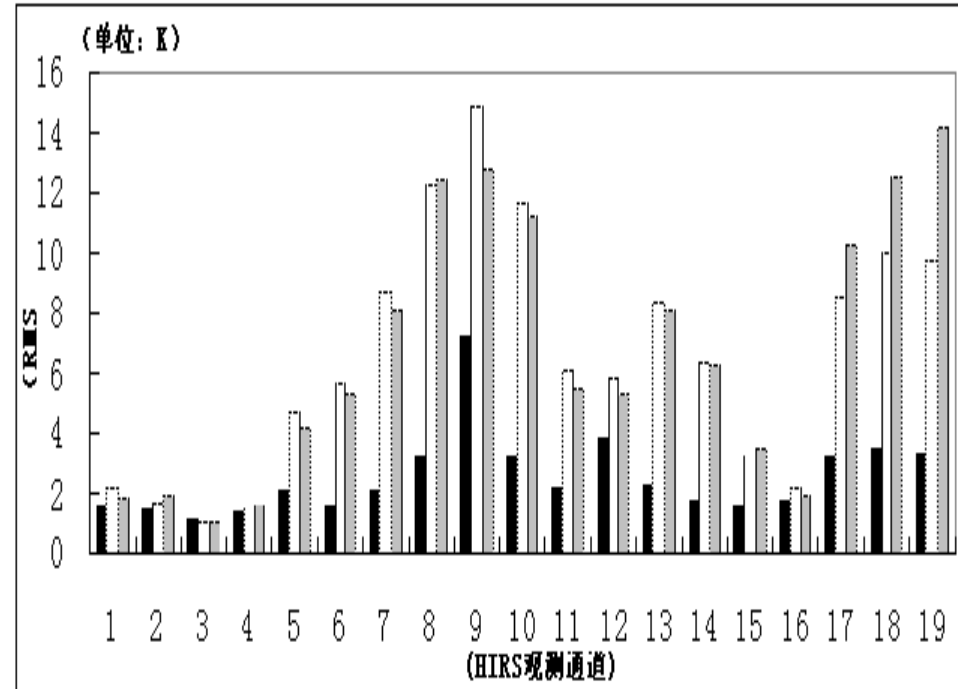
Biases from computed cloud-clear radiances to satellite observations



Biases from computed cloud-clear radiances to satellite observations

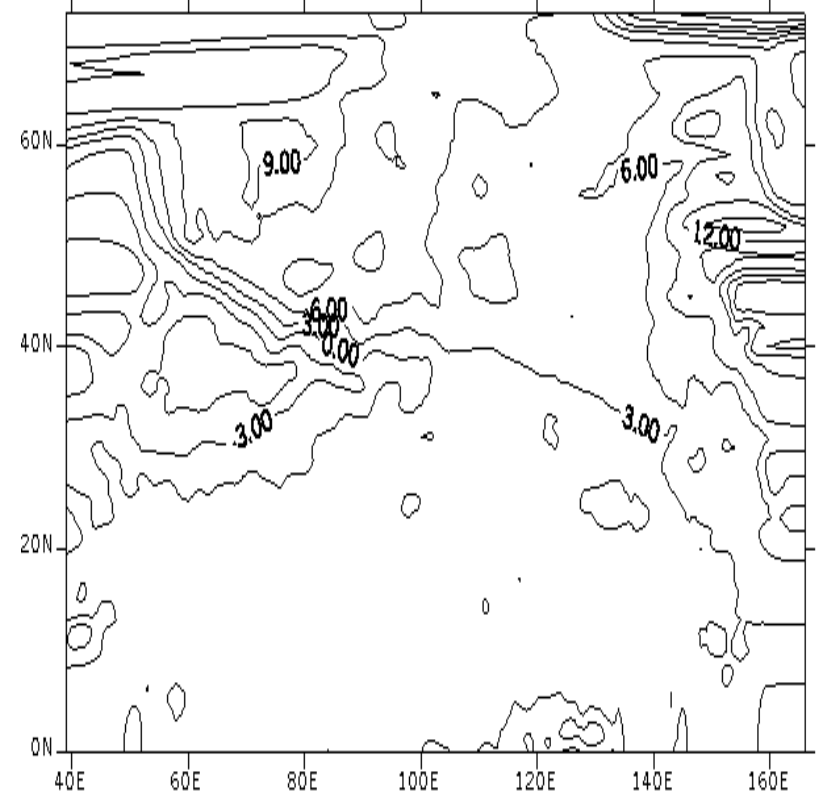
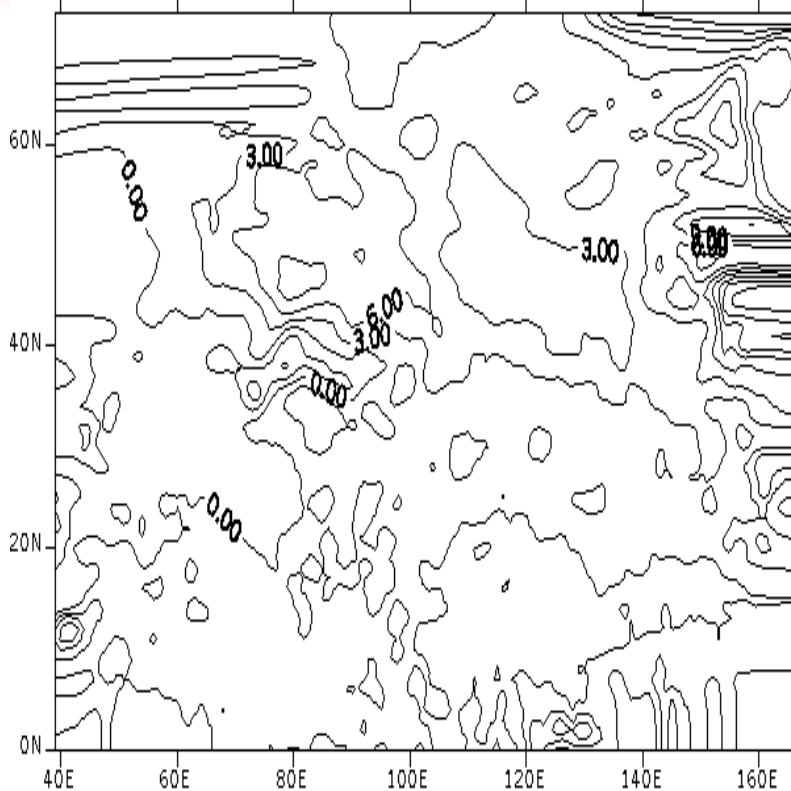


■ Daytime



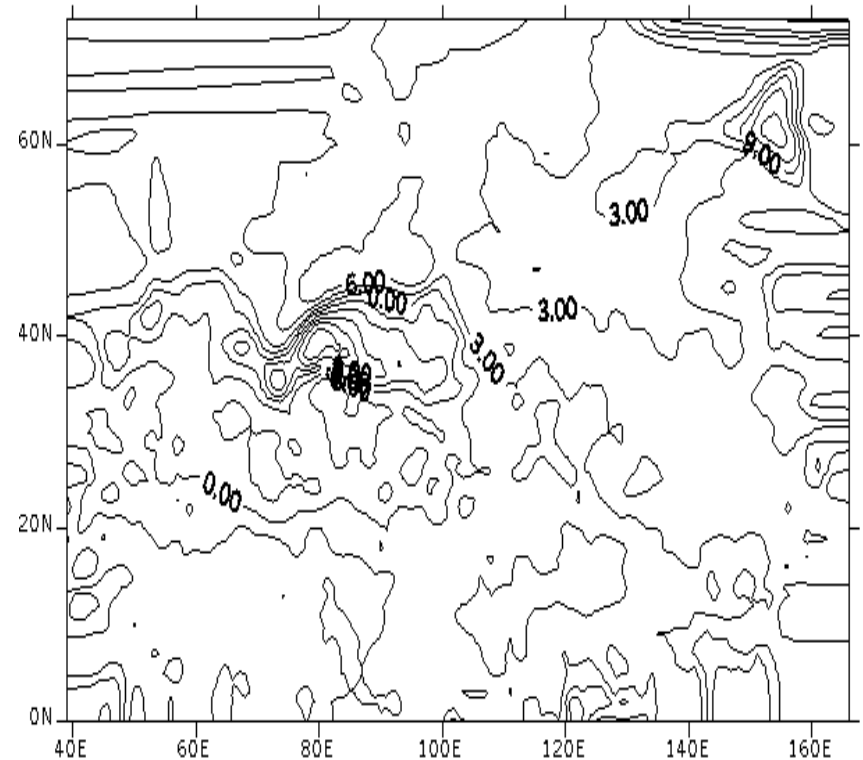
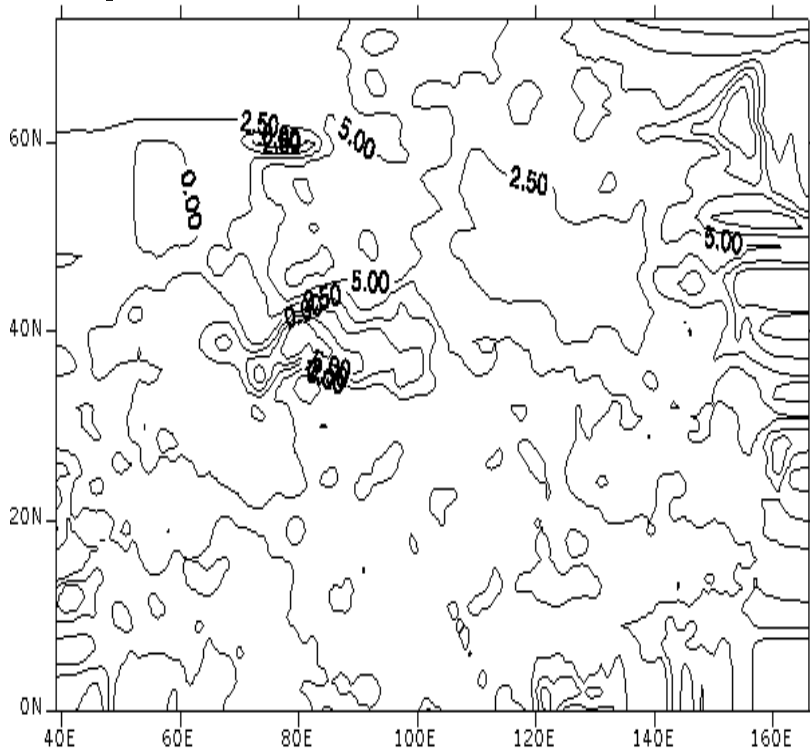
■ Night

Horizontal distribution of simulated bias (K)



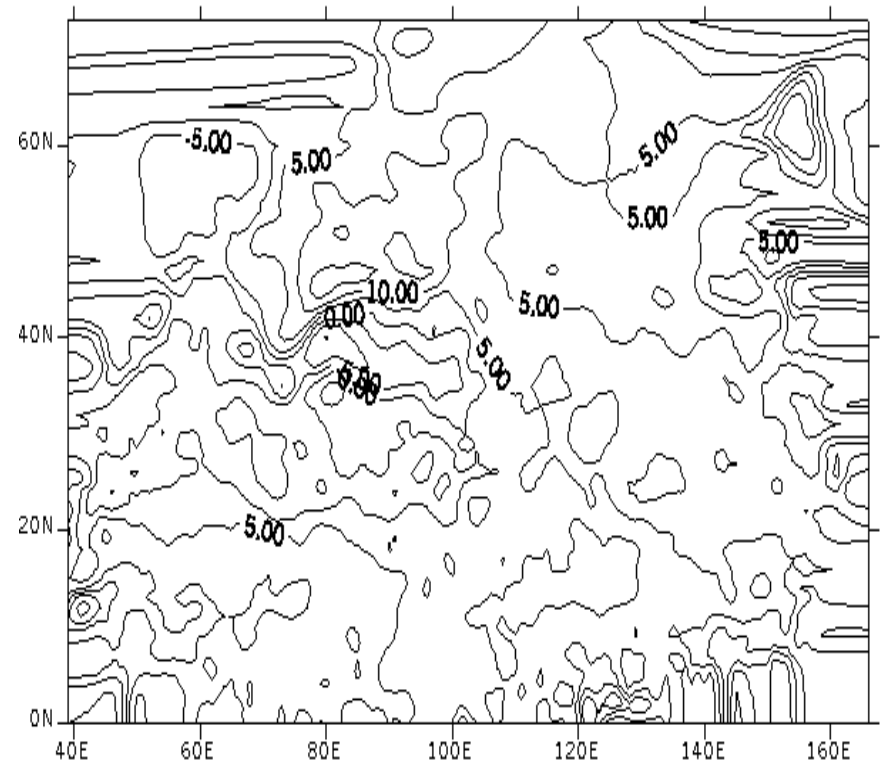
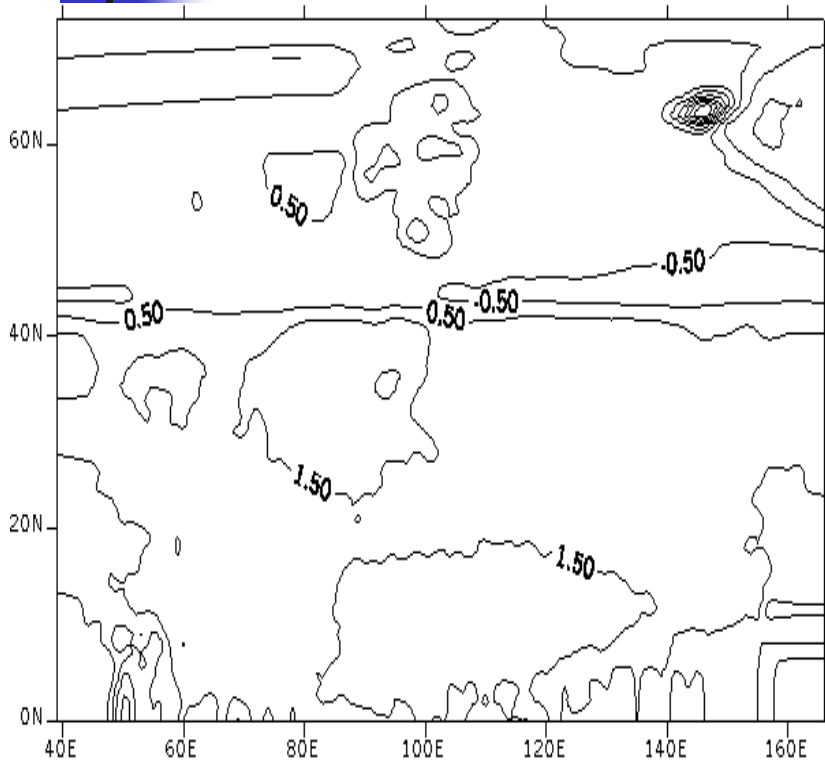
- simulated bias to channel 11 (left) and channel 12 (right)

Horizontal distribution of simulated bias (K)



- simulated bias to channel 6 (left) and channel 14 (right)

Horizontal distribution of simulated bias (K)



- simulated bias to channel 2 (left) and channel 17 (right)



Conclusion(1)

- 1 smaller RMS errors of simulated brightness temperature from real observations are available after cloud clearing
- 2 To channels for air temperature in upper atmosphere, the uniform simulated bias is available everywhere. To channels for water vapor and air temperature in lower atmosphere and channels for surface air temperature, similar uniform bias is obtained except in the area of Tibetan Plateau, where peak of simulated bias is demonstrated.



Conclusion(1)

- 3 STD at daytime and at night show a large bias in channels of 17-19 because of the impact of sun

Using 1D-VAR to get air temperature profile

- Forcing from satellite data

$$J = \sum_{r=0}^n [T - T_B]^T B^{-1} [T - T_B] + [X_r(T, q) - y_r^{obs}]^T W_r [X_r(T, q) - y_r^{obs}]$$

- Variation of the forcing

$$\delta J = 2 \sum_{r=0}^n F'_r(T, q) W_r [X_r(T, q) - y_r^{obs}] \cdot \delta Q$$

Weighting coefficients of assimilation variables



- Errors of HIRS data and the fast forward model are estimated
-
- Weighting coefficients are defined as the inverse of the square of these errors

scaling factor

While multi-variables are assimilated at same time, scale of every variable must be considered

- For instance, order of air temperature is about 1×10^2 and the order of water vapor is about 1×10^{-3}
- Scaling factor $S_j = X_j^{\max} - X_j^{\min}$

Gradients test(TL)

The check to tangent linear model

$$\Phi(\alpha) \equiv \frac{\|Q_r(z + \alpha h) - Q_r(z)\|}{\|\alpha P_r h\|} = 1 + O(\alpha)$$

$$\frac{\alpha \cdot (\text{Forward}(x + \frac{\delta x}{\alpha}) - \text{Forward}(x))}{TL(\delta x)} = 1$$



Gradients test(TL)

TL= -0.2828500366E+02

BRUTE FORCE: -0.2928326416E+02	0.1035292983E+01	1
BRUTE FORCE: -0.2839401245E+02	0.1003853917E+01	2
BRUTE FORCE: -0.2830657959E+02	0.1000762820E+01	3
BRUTE FORCE: -0.2836608887E+02	0.1002866745E+01	4
BRUTE FORCE: -0.2777099609E+02	0.9818275571E+00	5
BRUTE FORCE: -0.1525878906E+02	0.5394656658E+00	6
BRUTE FORCE: -0.3051757813E+02	0.1078931332E+01	7

Gradients test(AD)

The check to the adjoint model

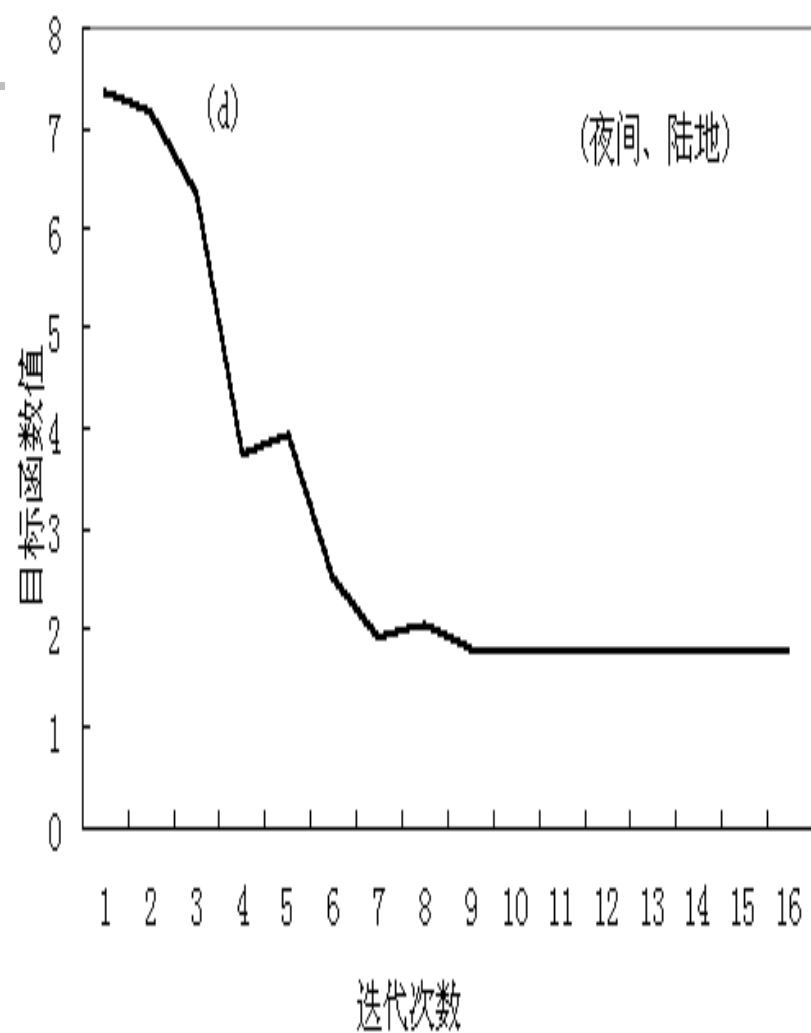
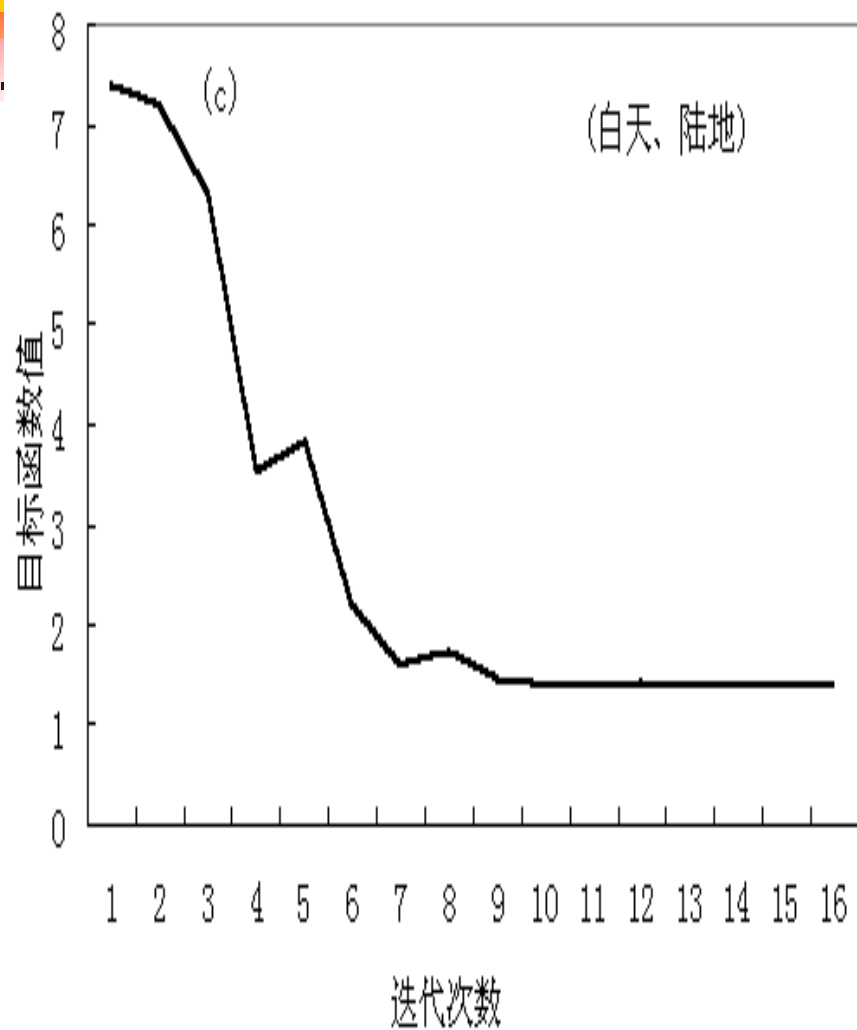
$$SumR = SumP$$

$$SumR = \sum DBT \cdot \delta y_{random} \quad SumP = \sum DPr of \cdot \delta x_{random}$$

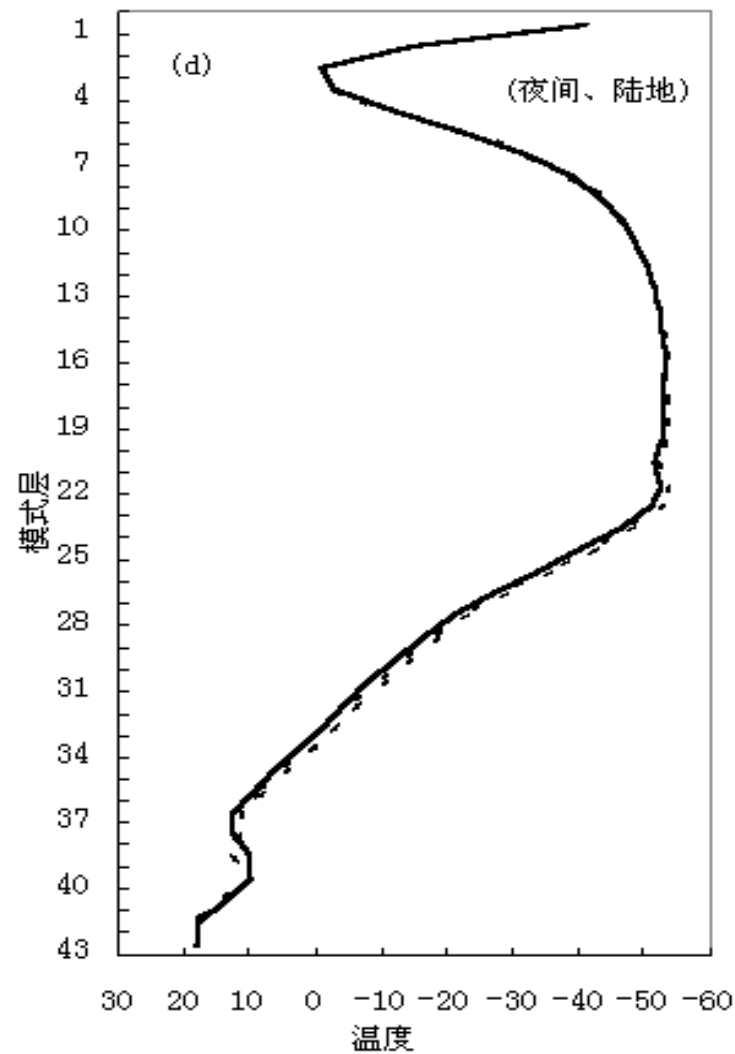
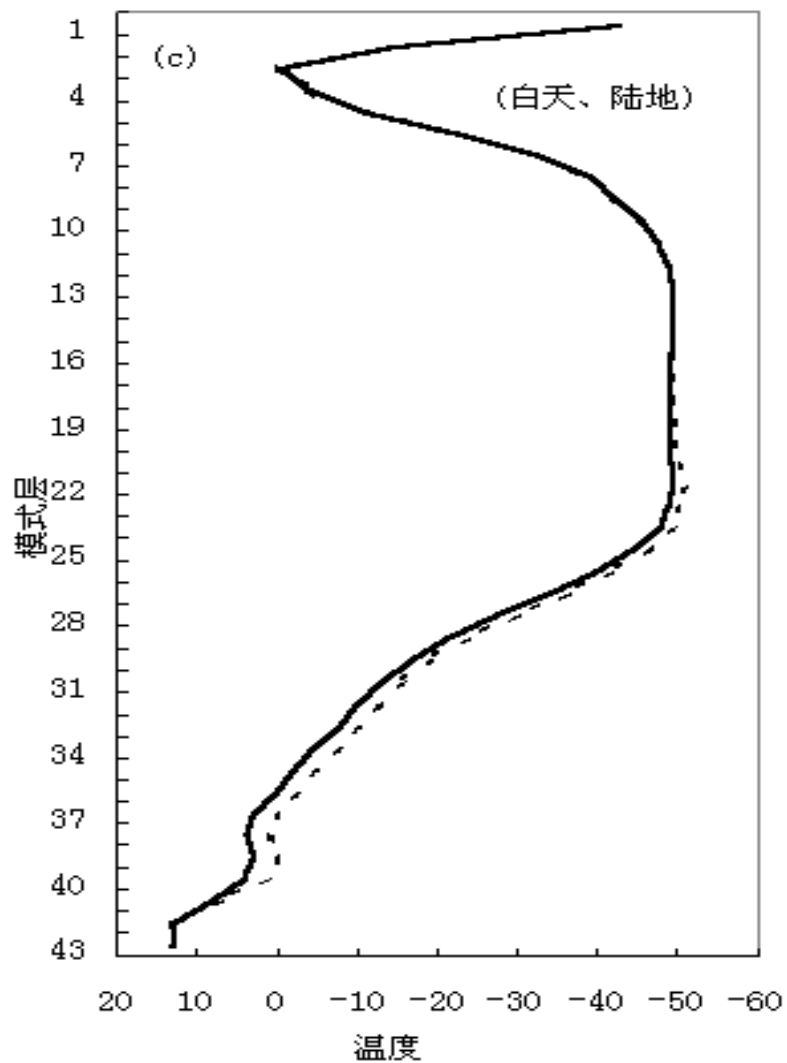
$$\delta x_{random} \xrightarrow{TL} DBT \quad \delta y_{random} \xrightarrow{AD} DPr of$$

- SUMRAD = -0.1463224602E+02
- SUMPROF = -0.1463224697E+02

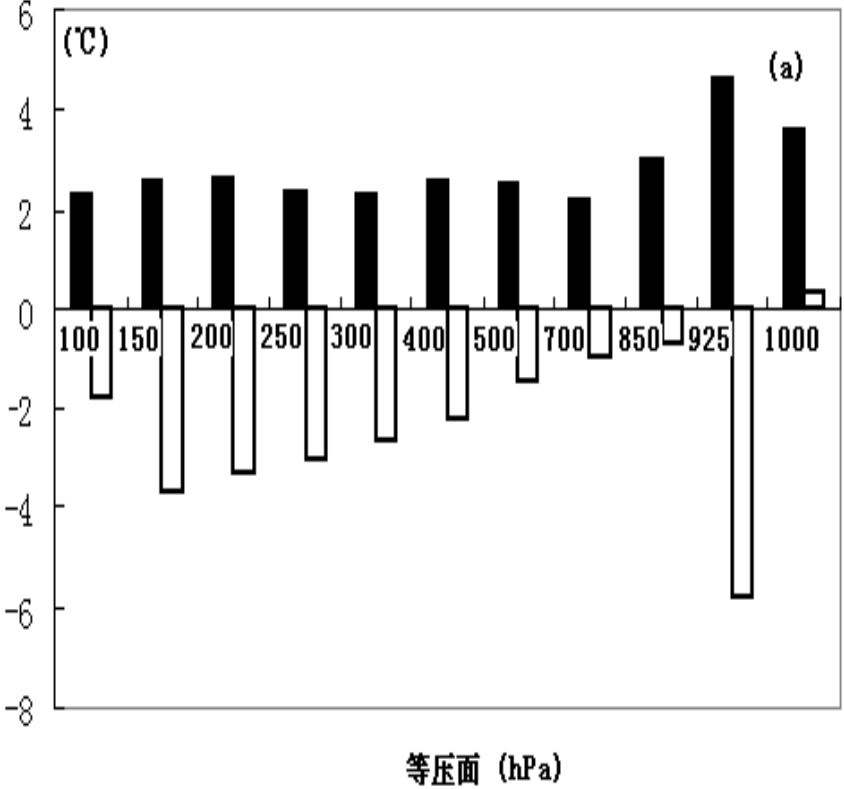
Convergence of cost function



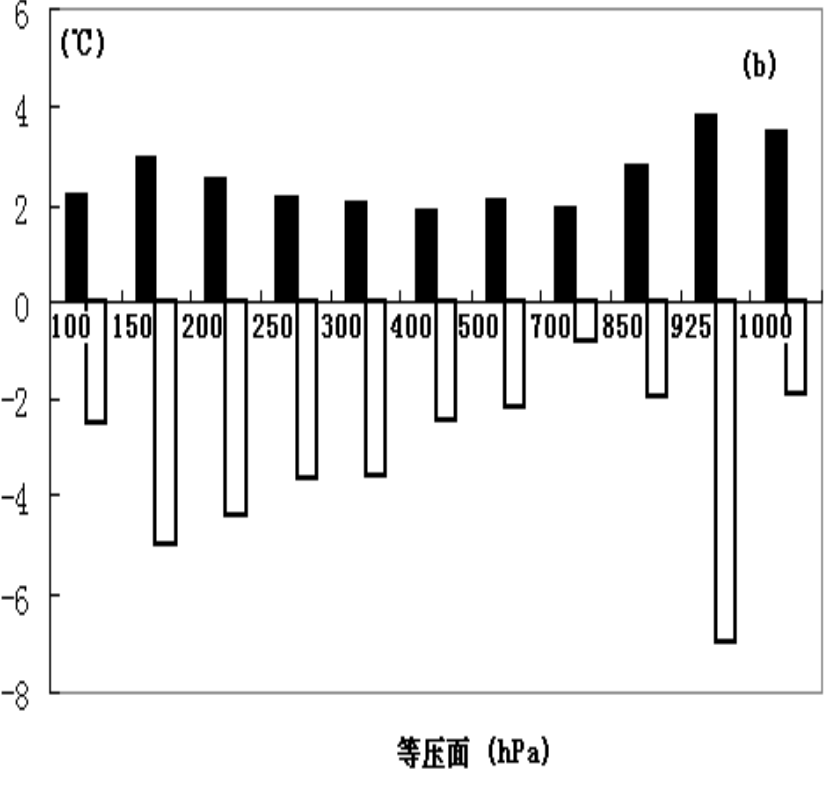
Assimilated air temperature



Biases from profiles of 1D-VAR to radio sounding data



■ 00UTC



■ 12UTC



4D-VAR

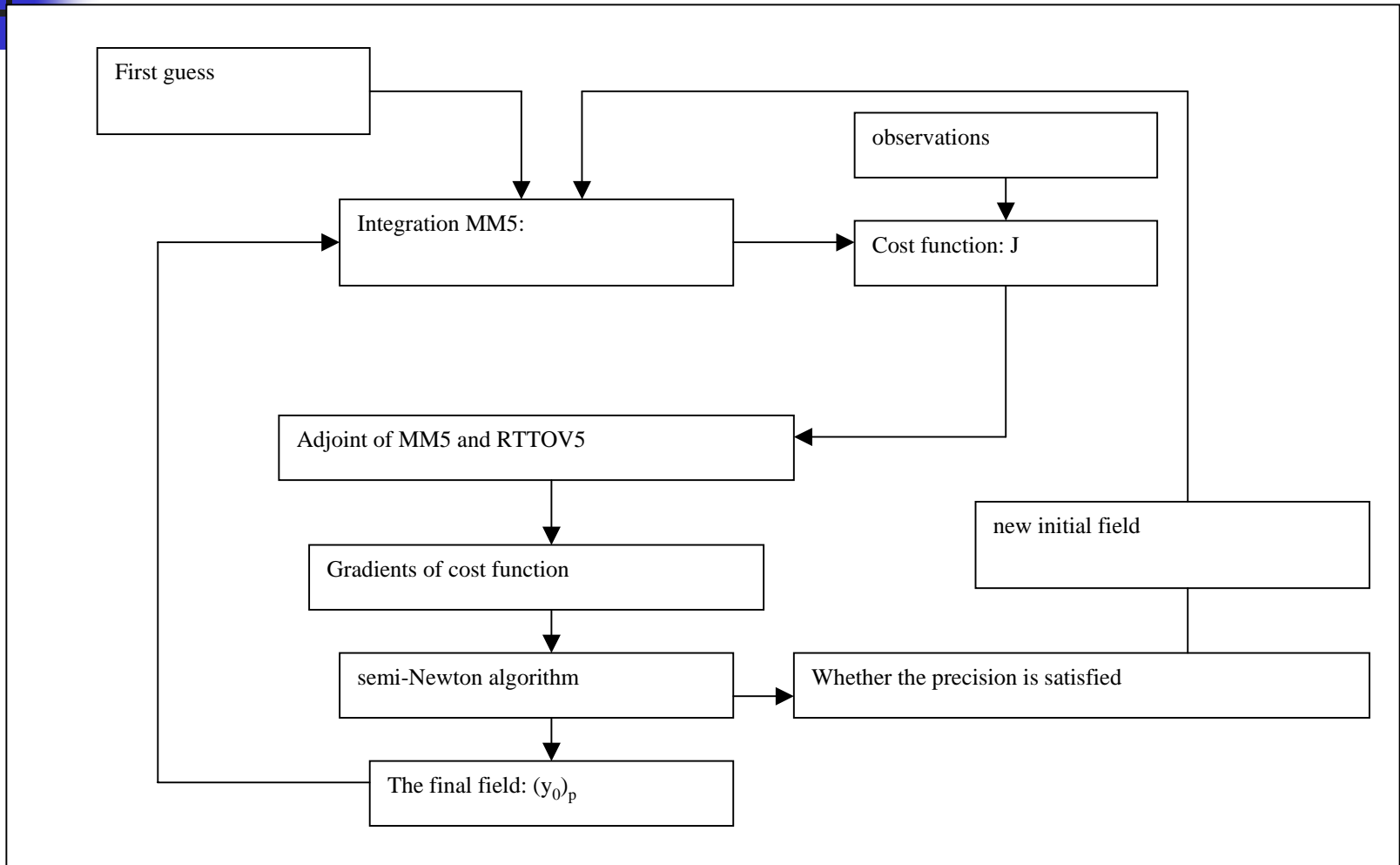
- Numerical model: MM5
- Central of the test domain : 25° N, 120° E
- Grid spacing : 45 km
- total grids : $61 * 61$
- Prediction period:
 - 12 UTC, July, 21, 2002 – 12UTC, July, 22, 2002
- Fast transfer model : RTTOV5



4D-VAR

- Assimilation window: 2 hours
- Background data: T106 analysis fields
- Satellite data: HIRS cloud-clearing radiances

The flow of the 4D-VAR

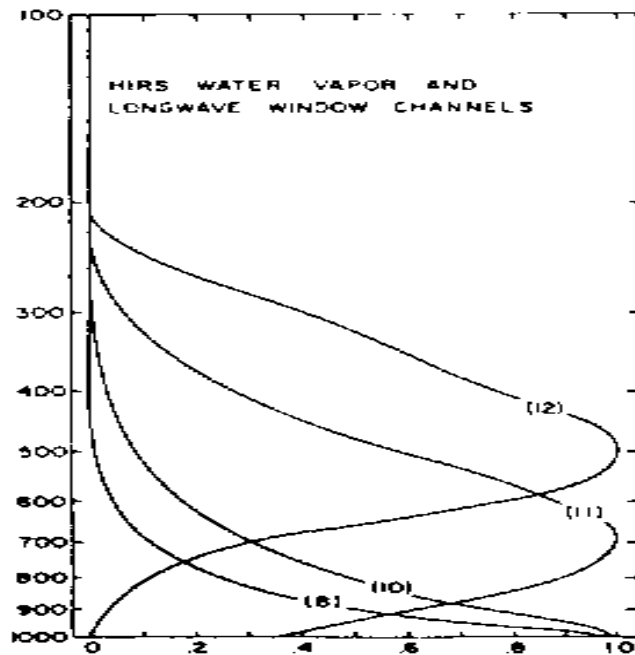
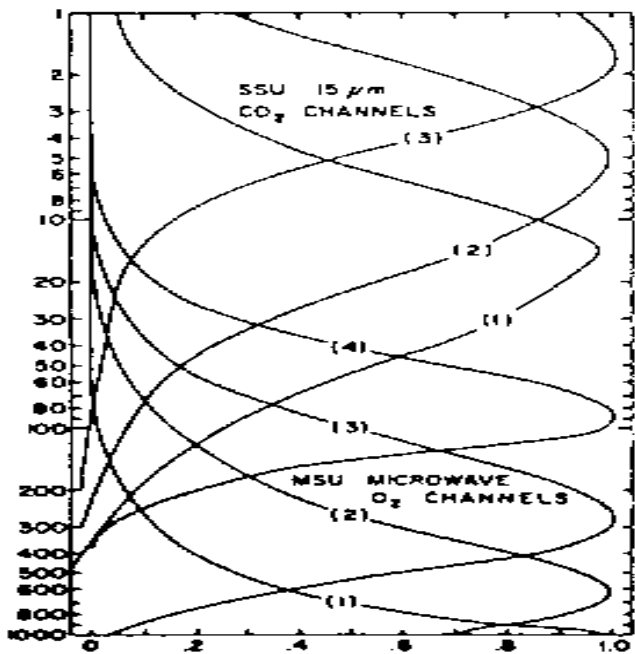
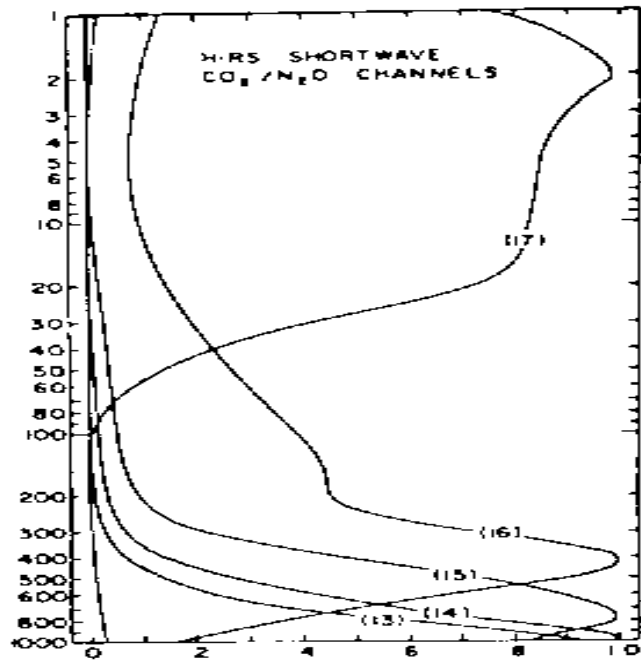
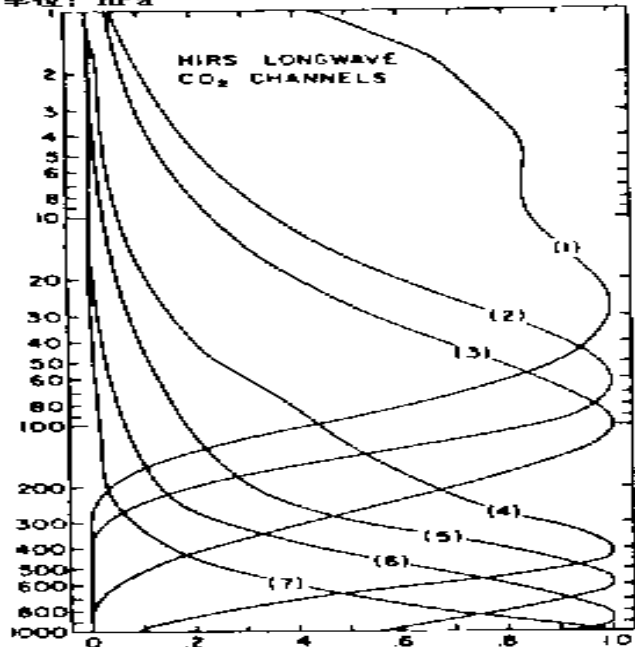


Channel selection



- Errors of these channels are independent for the pressure layers of the channels do not overlap
- Selected HIRS channels should be more than the number of vertical levels in MM5
- Errors of these channels should be less than a given value

单位: hPa



Channel selection



- O₃ channel and short wave window channels (channel 17-19) are ignored
- Data from 9 HIRS channels are used to be involved in the data assimilation
- Channel: 2, 4, 6, 7, 8, 10, 11, 12, 13, 14, 15

How to select the independent satellite data



- In general errors of sounding data should be independent to each other
- The density of cloud-clear HIRS data are much more than the radio soundings
- Every two or three HIRS observations are introduced into the data assimilation

Gradients test



- Similar to the check to tangent linear model in HIRS 1D-VAR:

$$\frac{TL (x + \alpha \cdot \delta x)}{AD (\alpha \cdot \delta x)} = 1$$



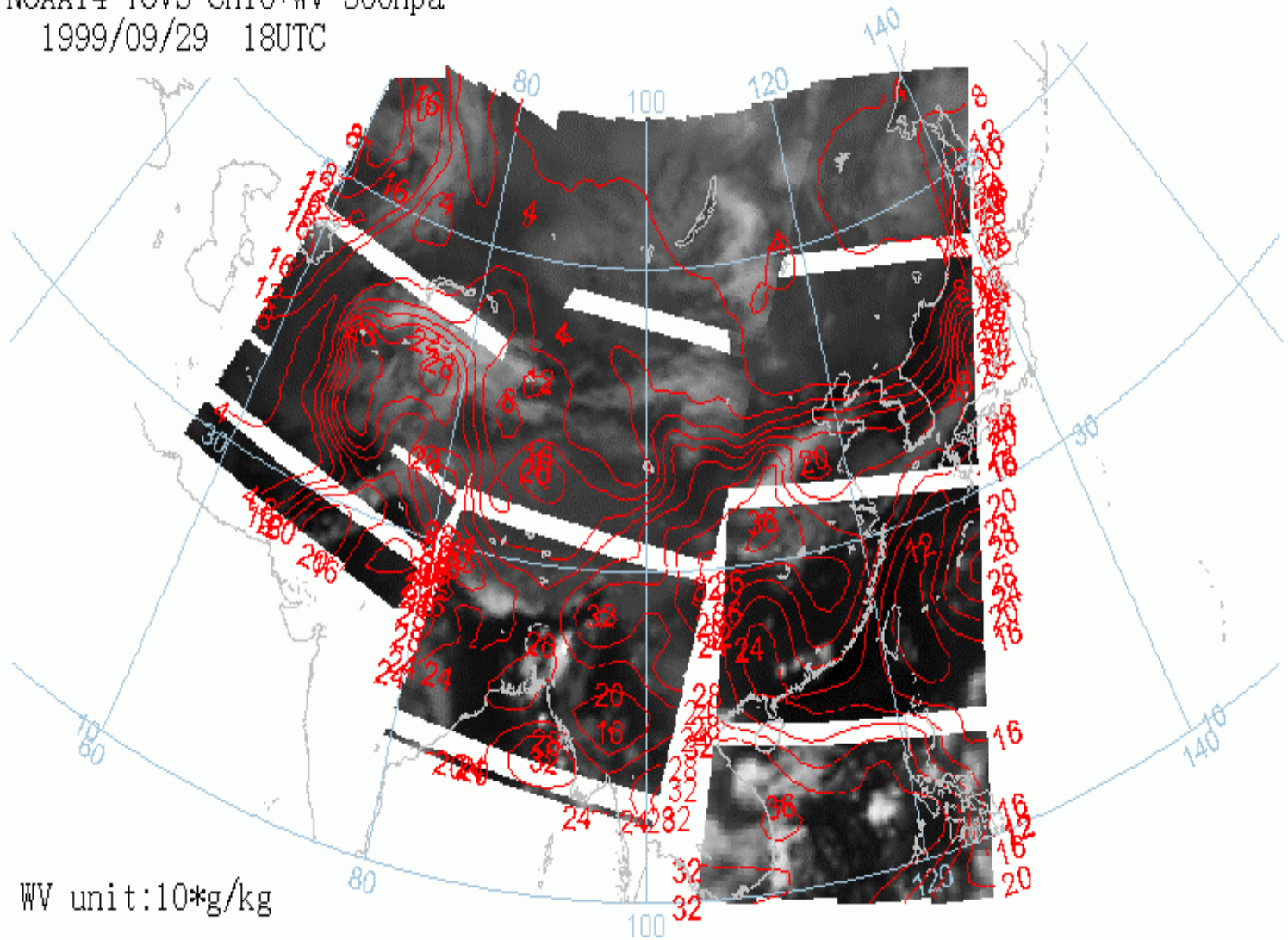
Gradients test

$$\alpha_0 = 0.00000000E+00$$

$$++ < \alpha_{-00} = 9.99999990E-05 > ++$$

l=1	$\alpha =$	0.10000E-04	F(A) =	0.1002861300E+01
l=2	$\alpha =$	0.10000E-05	F(A) =	0.9989470500E+00
l=3	$\alpha =$	0.10000E-06	F(A) =	0.1019896300E+01
l=4	$\alpha =$	0.10000E-07	F(A) =	0.9372019800E+00
l=5	$\alpha =$	0.10000E-08	F(A) =	0.5512952800E+00

NOAA14 TOVS CH10+WV 500hpa
1999/09/29 18UTC

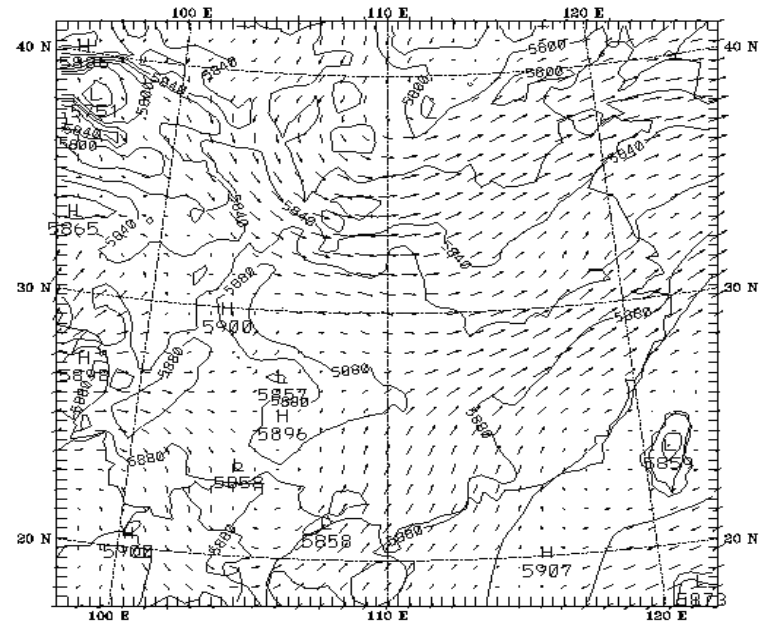
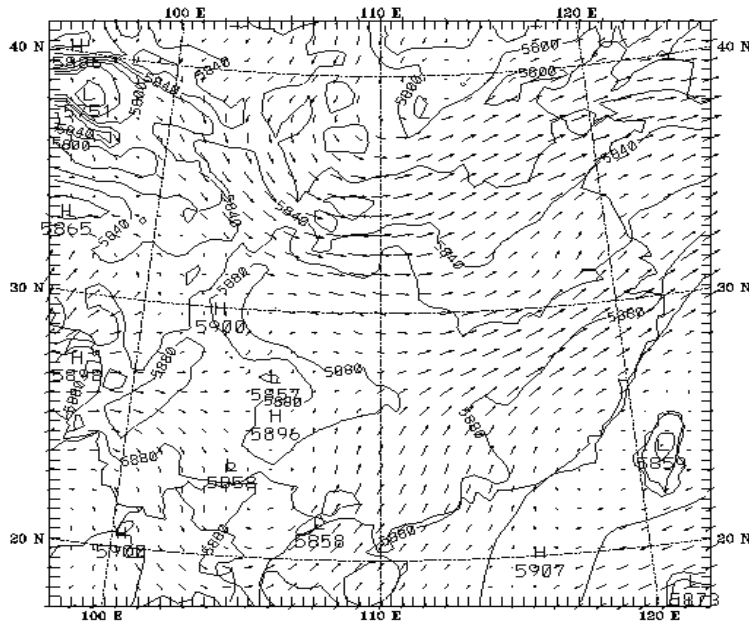


WV unit:10*g/kg

Geopotential height and streamline field on 500 hPa at 24th hour (the control test: left; the assimilation test: right)

PRESSURE=500. mb HEIGHT 1m/s | 2072113 = 2072113 : 0.00H SMOOTH= 0
 PRESSURE=500. mb BARB'Uv 1m/s | 2072113 = 2072113 : 0.00H SMOOTH= 0

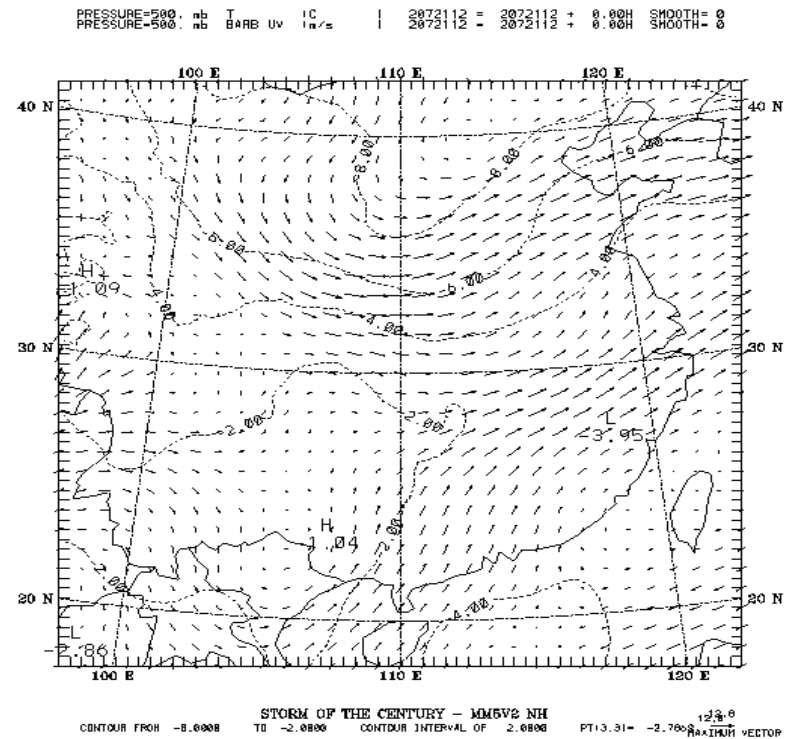
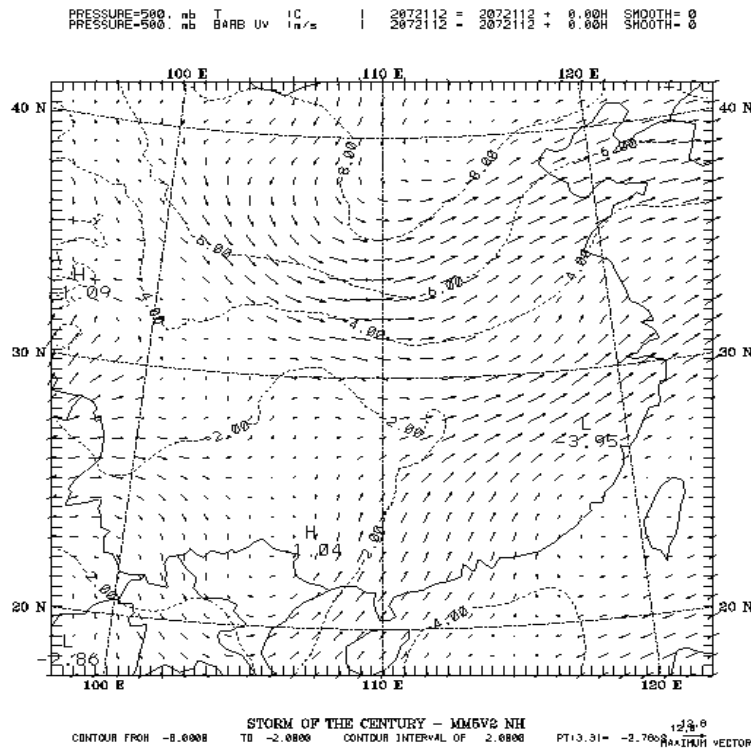
PRESSURE=500. mb HEIGHT 1m/s | 2072113 = 2072113 : 0.00H SMOOTH= 0
 PRESSURE=500. mb BARB'Uv 1m/s | 2072113 = 2072113 : 0.00H SMOOTH= 0



STORM OF THE CENTURY - MM5V2 NH
 12.0
 CONTOUR FROM 5740.0 TO 5908.0 CONTOUR INTERVAL OF 20.000 PT13.31- 5085.0 MAXIMUM VECTOR

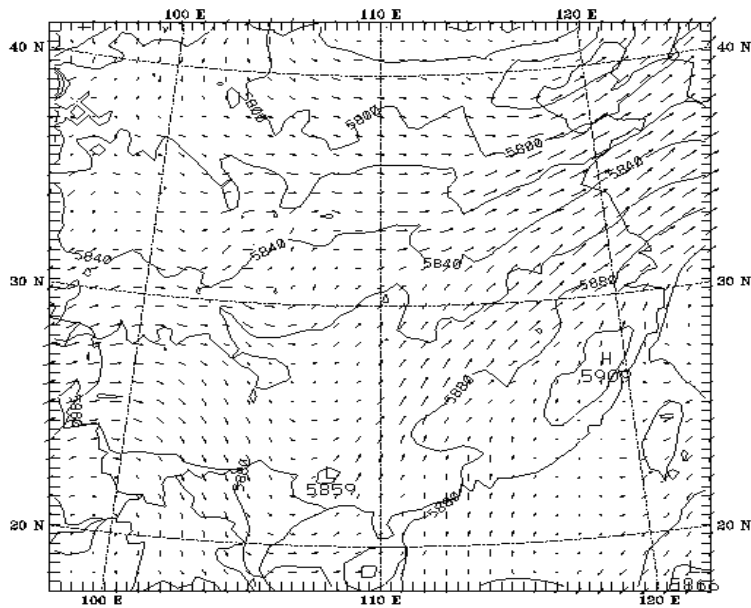
STORM OF THE CENTURY - MM5V2 NH
 12.0
 CONTOUR FROM 5740.0 TO 5908.0 CONTOUR INTERVAL OF 20.000 PT13.31- 5085.0 MAXIMUM VECTOR

Air temperature and streamline field on 500 hPa at 24th-hour (the control test: left: the assimilation test: right)



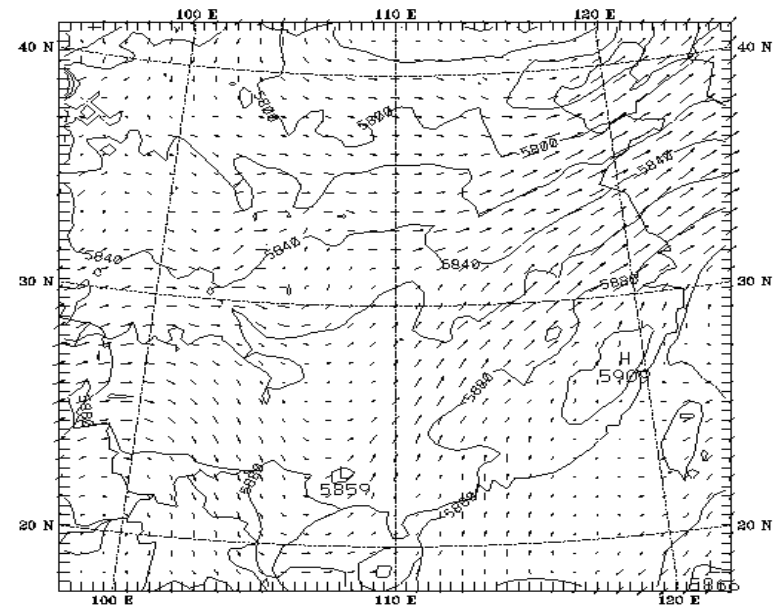
Geopotential height and streamline field on 500 hPa at 24th-hour (the control test: left; the assimilation test: right)

PRESSURE=500. mb HEIGHT 1m | 2072212 = 2072112 + 24.00H SMOOTH= 0
 PRESSURE=500. mb BARB Uv 1m/s | 2072212 = 2072112 + 24.00H SMOOTH= 0



STORM OF THE CENTURY - MM5V2 NH
 CONTOUR FROM 5740.0 TO 5900.0 CONTOUR INTERVAL OF 20.000 PT13.31- 5003.10
 MAXIMUM VECTOR

PRESSURE=500. mb HEIGHT 1m | 2072212 = 2072112 + 24.00H SMOOTH= 0
 PRESSURE=500. mb BARB Uv 1m/s | 2072212 = 2072112 + 24.00H SMOOTH= 0

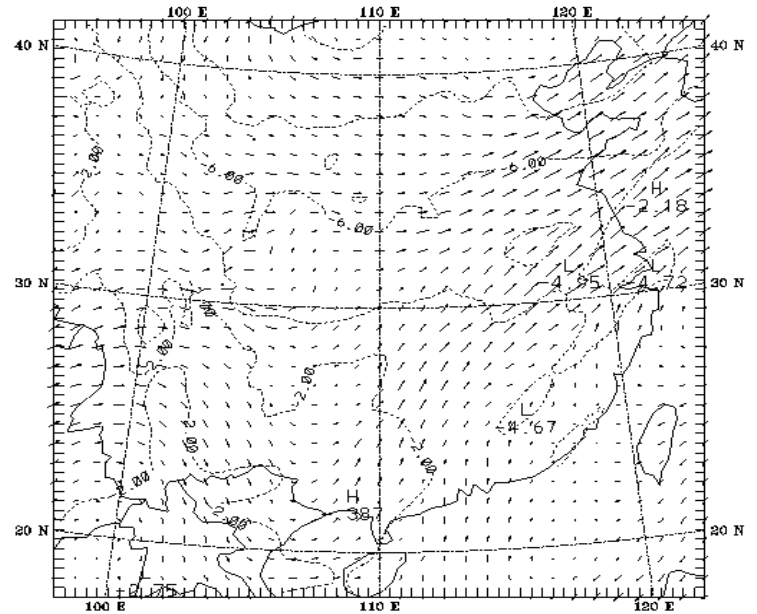
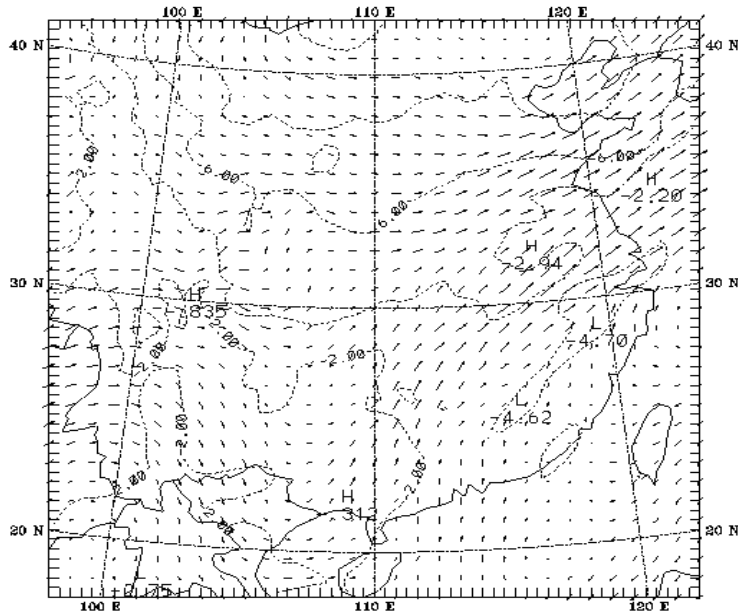


STORM OF THE CENTURY - MM5V2 NH
 CONTOUR FROM 5740.0 TO 5900.0 CONTOUR INTERVAL OF 20.000 PT13.31- 5003.10
 MAXIMUM VECTOR

Air temperature and streamline field on 500 hPa at 24th-hour (the control test: left; the assimilation test: right)

PRESSURE=500: mb T BARB UV IC | 2072212 = 2072112 + 24.00H SMOOTH= 0
 PRESSURE=500: mb T BARB UV IC | 2072212 = 2072112 + 24.00H SMOOTH= 0

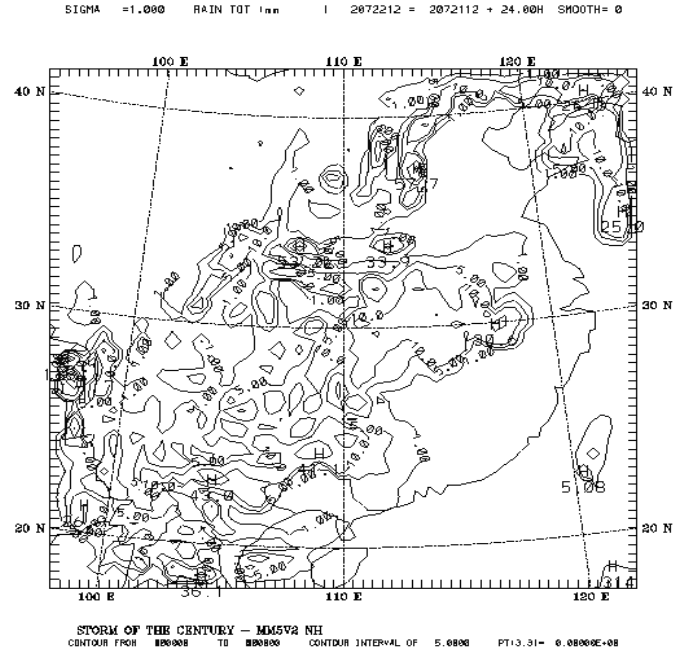
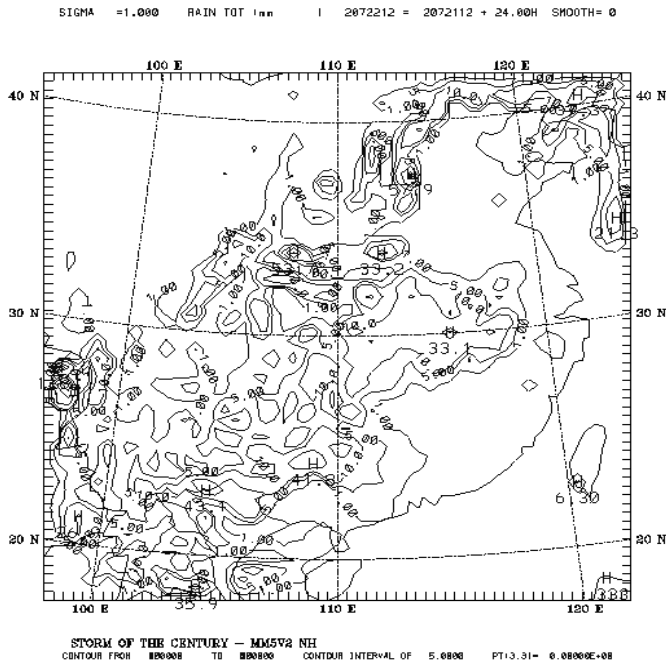
PRESSURE=500: mb T BARB UV IC | 2072212 = 2072112 + 24.00H SMOOTH= 0
 PRESSURE=500: mb T BARB UV IC | 2072212 = 2072112 + 24.00H SMOOTH= 0



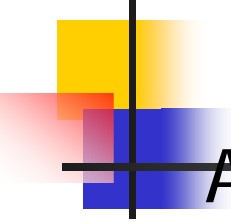
STORM OF THE CENTURY - MM5V2 NH
 CONTOUR FROM -10.000 TO -2.0000 CONTOUR INTERVAL OF 2.0000 PT13.31= -2.4930
 10/3/2
 MAXIMUM VECTOR

STORM OF THE CENTURY - MM5V2 NH
 CONTOUR FROM -10.000 TO -2.0000 CONTOUR INTERVAL OF 2.0000 PT13.31= -2.4650
 10/3/0
 MAXIMUM VECTOR

total 24-hours' precipitation in control test (right) and assimilation test with satellite data (left)



Conclusion



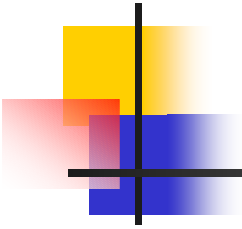
Assimilation HIRS data from satellite could get certain improvement to numerical model prediction

- The forecasted quantity of precipitation and the location of precipitation center are quite different from the results in control experiment
- Due to the impact of cloud, meso-scale forecasting in cloud area could not get any change

Work in the future



1. Introducing AMSU data into the 4D-VAR
2. Prolonging the length of assimilation window
3. Adding other satellite products, such as estimated precipitation, into the 4D-VAR
4. Testing other examples



Thank you!