



Analysis of Arctic clouds by means of hyper-spectral satellite

F. Romano*, D. Cimini[°], E. Di Tomaso*, E. Ricciardelli*, V. Cuomo*

***Institute of Methodologies for Environmental Analysis (IMAA/CNR)**

[°]CETEMPS, University of L'Aquila

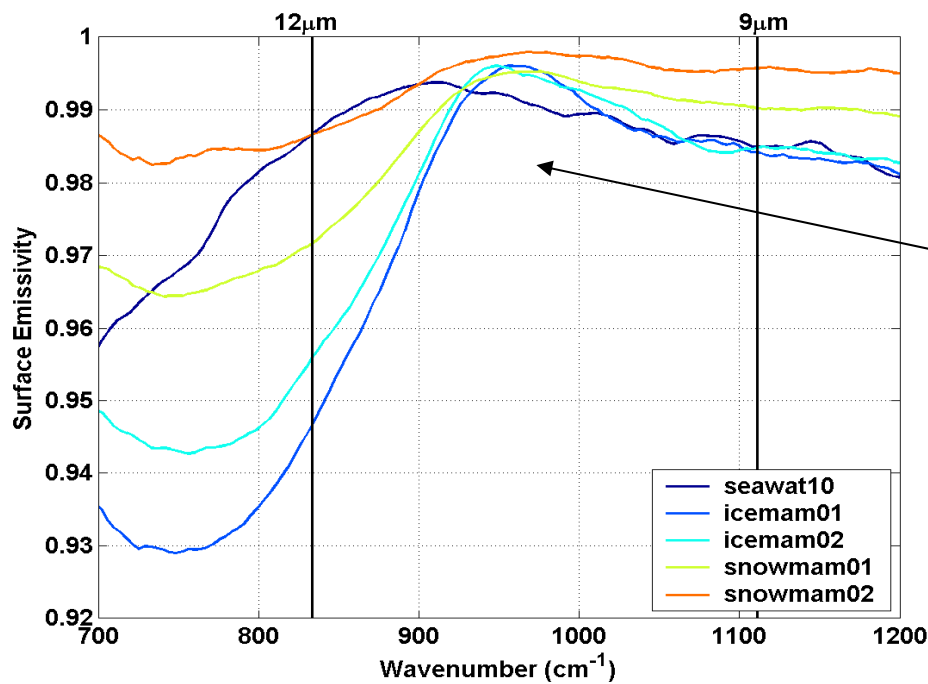


Satellite cloud detection in polar regions is difficult because of:

- **cold surface temperatures**
- **little infrared and visible contrast with snow/ice surface**
- **strong temperature inversions**
- **usually low and thin clouds**

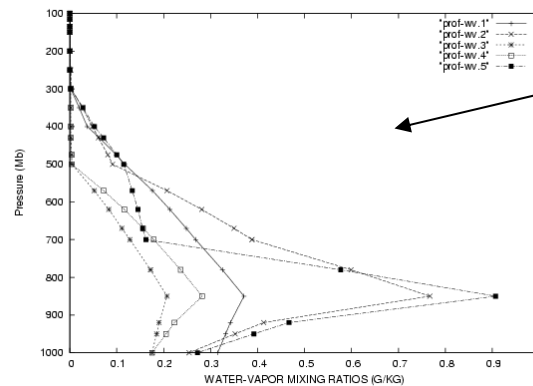
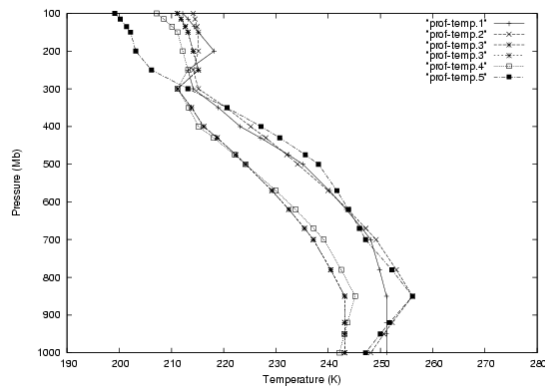
In addition during polar nights there is not possible apply reflectance tests (e.g. 1.6 μm test).

Infrared spectral emissivity differ significantly and it can play a crucial role in cloud detection due to the spectral features in the 700-1200 cm^{-1} range.



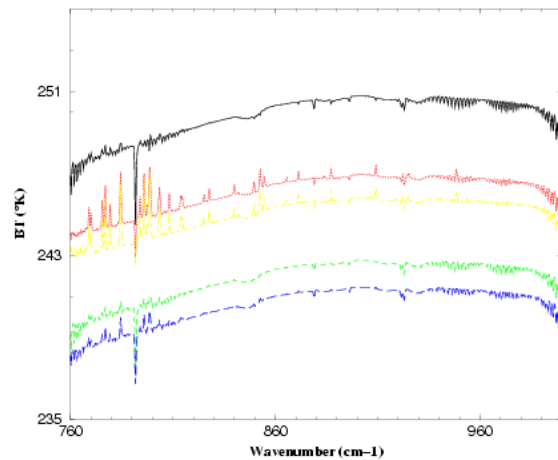
**from MODIS UCSB
Emissivity Library (PI:
Zhengming Wan)**

IASI Tb spectra are simulated in clear sky using LBLRTM for 305 arctic profiles

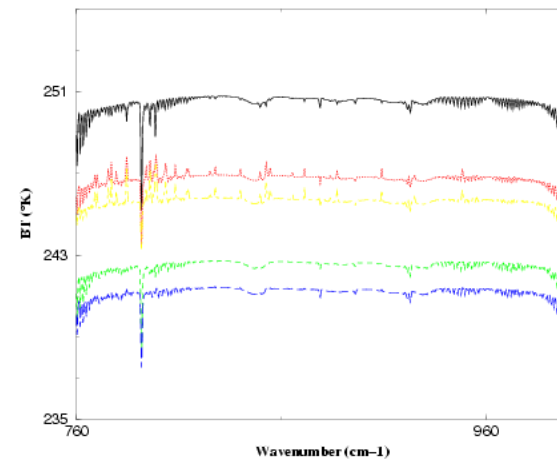


Water and temperature profiles

Ice Emissivity



Water emissivity



IASI Tb spectra are simulated in cloudy sky using RTX

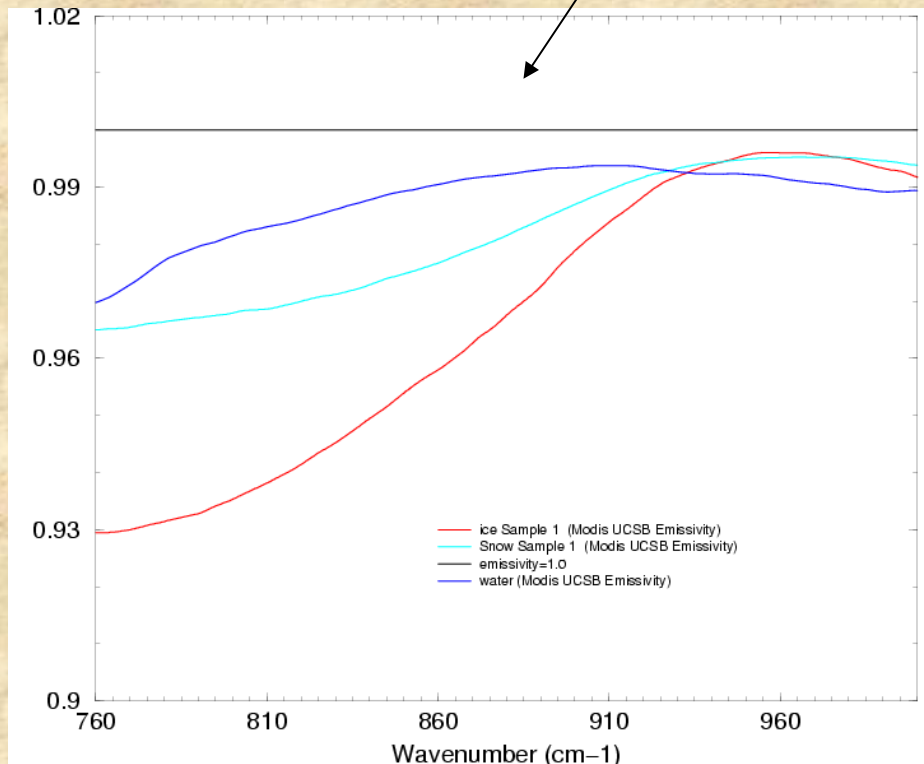
The numerical model RTX [Rizzi et al., 2001; Amorati and Rizzi, 2002, Maestri et al., 2005], solves the radiative transfer equation with the adding and doubling method taking into account the multiple scattering by randomly-oriented particles with a plane of symmetry. Polarized radiation is considered in term of Stokes parameters under the hypothesis of a plane-parallel and vertically inhomogeneous atmosphere including both thermal and solar sources.

Spectral properties of atmospheric gases are computed with the LBLRTM model while the extinction and scattering coefficients, the single scattering albedo and the Lagrange coefficients are computed for a gamma-modified size distribution of spherical cloud particles (water and ice) using a Mie code (Wiscombe 1979).

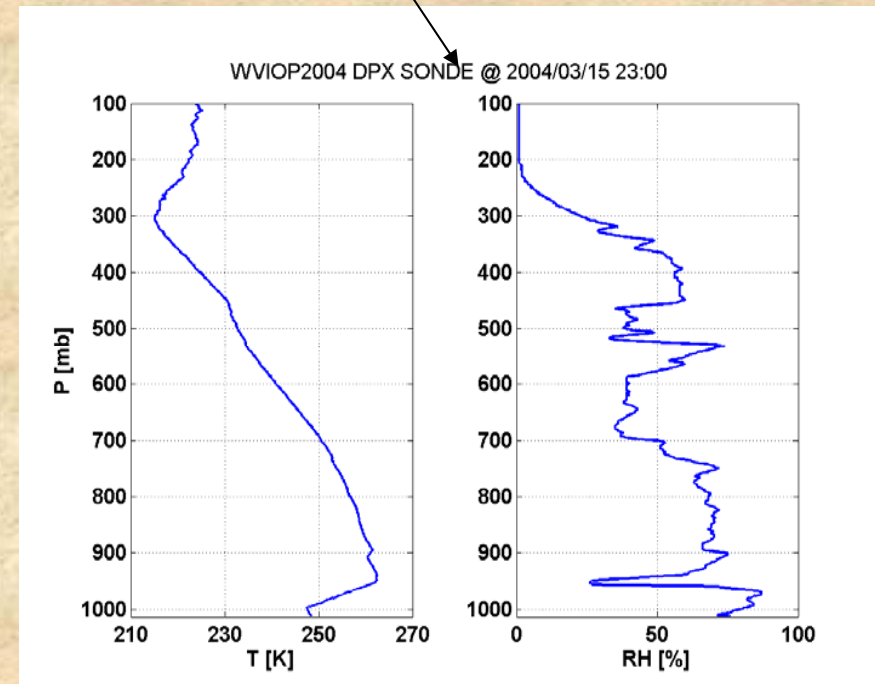
Cloud parameters for cloud spectra simulation

- **T and RH profiles**
- **cloud top 2.2, 2.7, 3.2, 3.7, 5.1, 5.5 km**
- **Cloud phase: liquid, ice, mixed phase**
- **Effective radius: 5, 15, 25, 30, 50, 70 μm**
- **Ice or water liquid content : 0.001 0.005 0.01 0.03
0.05 0.07**

Surface emissivity (ϵ): sea water, ice and snow

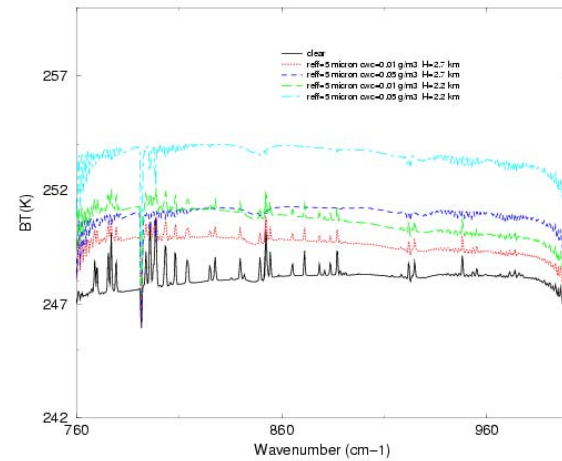
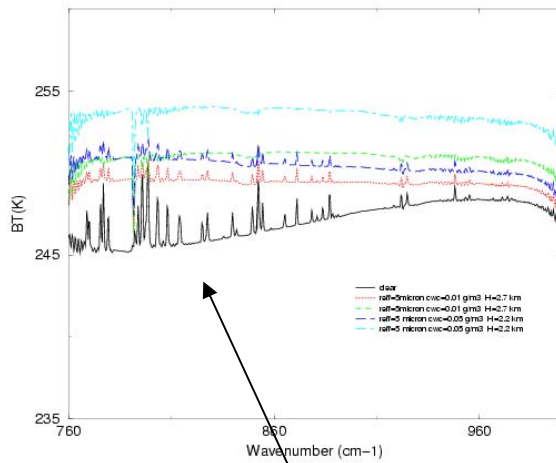


Temperature and water vapour Profile



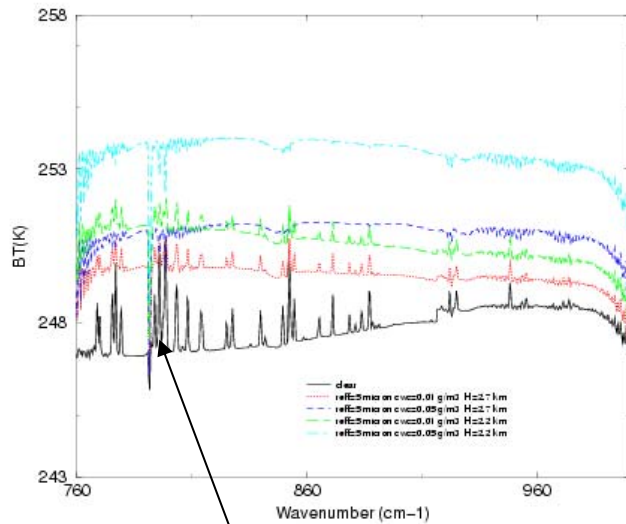
Clear and cloudy spectra using ice emissivity

Water clouds



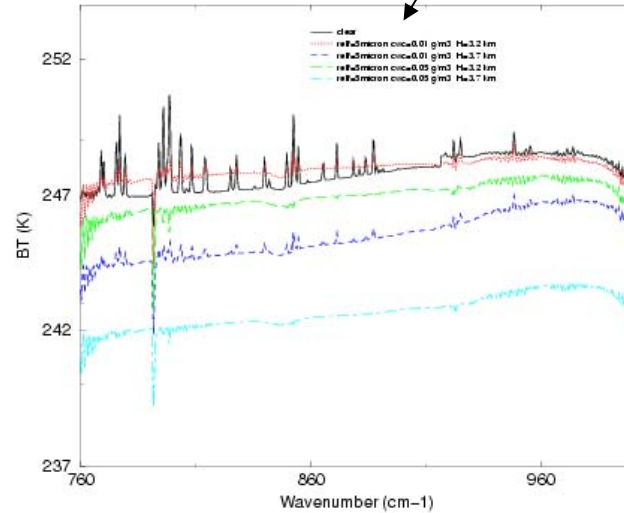
Ice clouds

Clear and cloudy spectra using snow emissivity

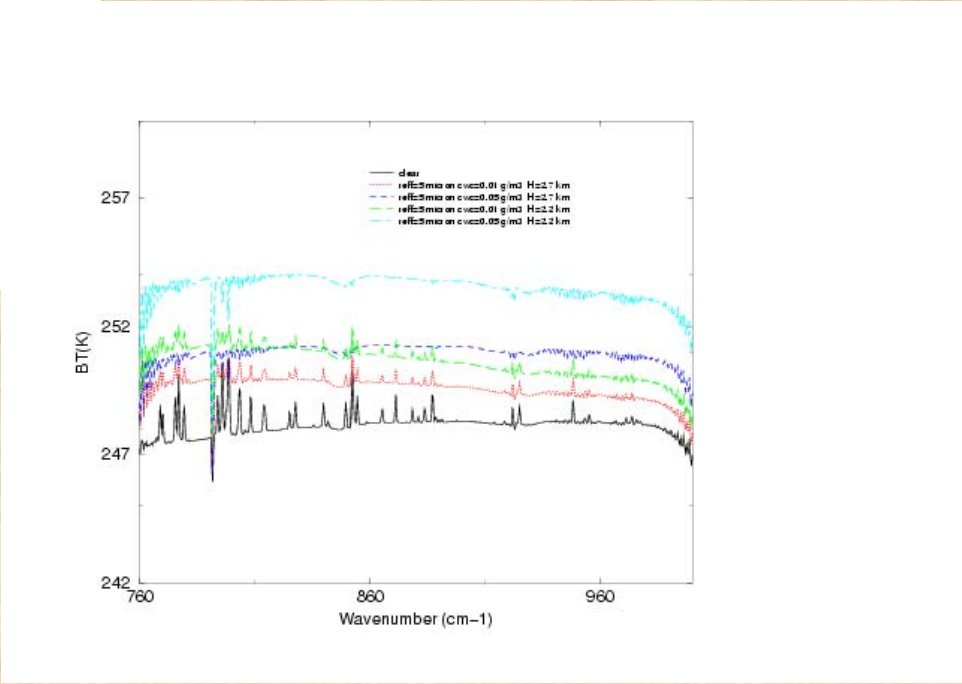
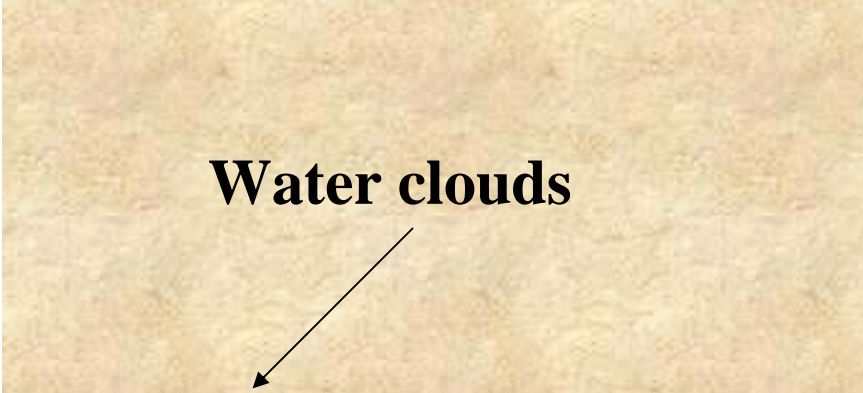
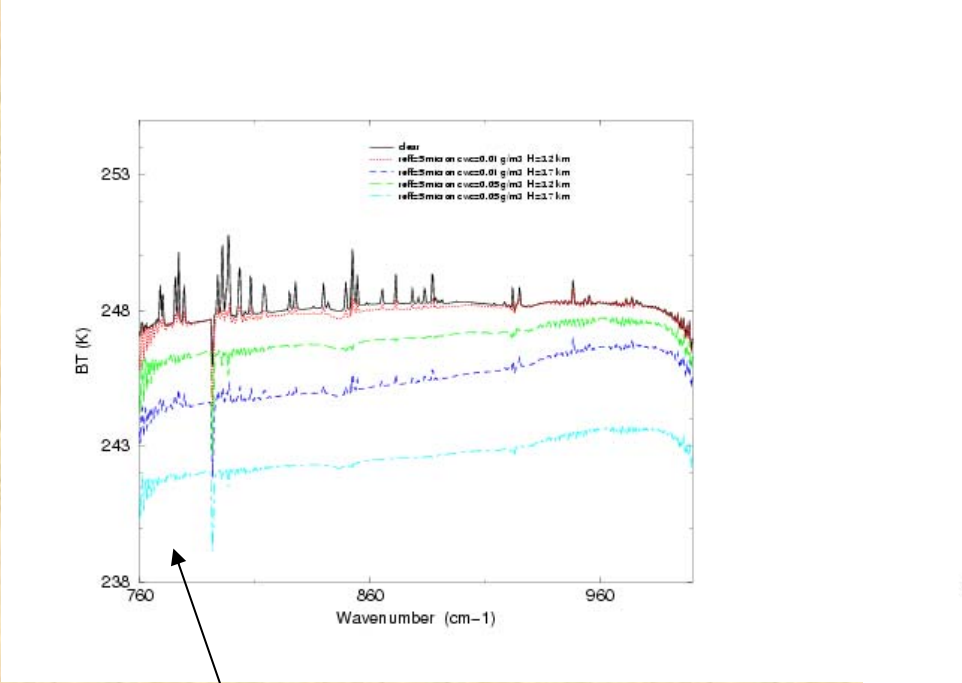


Ice clouds

Water clouds



Clear and cloudy spectra using water emissivity



Ice clouds

It is evident that spectral features caused by emissivity typical of polar surfaces are very similar to cloud spectral signatures.

It is evident that the clouds and the emissivity cause the slope of the atmospheric window to change significantly, but it seems difficult to detect on the basis of window spectral region clouds presence.

Ice clouds for instance increase the slope for a constant emissivity, but decrease the slop for ice surface emissivity.

Cloud detections

We have developed an identification algorithm that exploits the expected spectral signal in the window region for polar region.

Examining the slope in the 750-980 cm⁻¹ range for clear-sky spectra with different surface emissivity and different temperature and humidity profiles, it is evident that the slope values remains within a different range for each emissivity and the window shape it is the same for the same emissivity, the same satellite zenith angle and the same solar illumination.

Slope window test

The test is restricted to a suitable set of micro window differences.

Central wavenumbers of micro windows:

790.0 cm⁻¹

803.5 cm⁻¹

885.7 cm⁻¹

953.0 cm⁻¹

Thresholds

The thresholds have been derived using a measured and a simulated clear dataset. Simulated data also provide a useful estimates of variability around the mean value which should be exploited.

In this approach, the difference between predicted and measured channels must be within a range, whose medium value is calculated from measured data, while the boundaries (minimum and maximum) are estimated from simulated data.

The algorithm selects the thresholds according to satellite zenith angle, surface type, solar illumination and the brightness temperature ranges of the image.

Correlation test

When the probability to be clear or cloud is low, we apply the correlation test. We have built a database based on clear spectra on arctic region for different satellite zenith angles, solar illumination and surface type, identified using the Cloud Profiling Radar data and ground measurements.

The fit is restricted to a suitable set of 20 wavenumbers that give a fine representation of the atmospheric window shape.

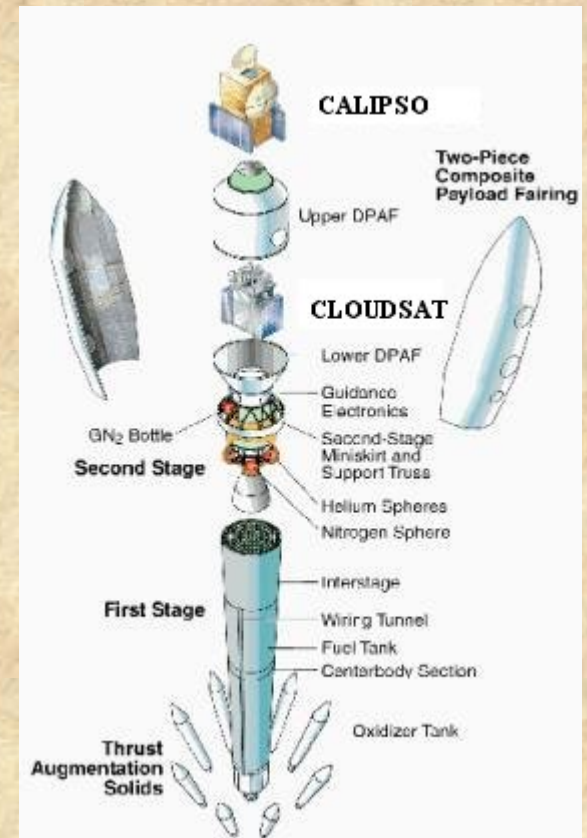
The correlation index is computed on a restricted set of spectra according to surface, solar zenith angle and solar illumination of the examined IASI spectrum :

$$r = \frac{\text{cov}(BT, BT_0)}{(\text{var}(BT) \text{var} BT_0)^{1/2}}$$

The Cloud Profiling Radar (CPR)

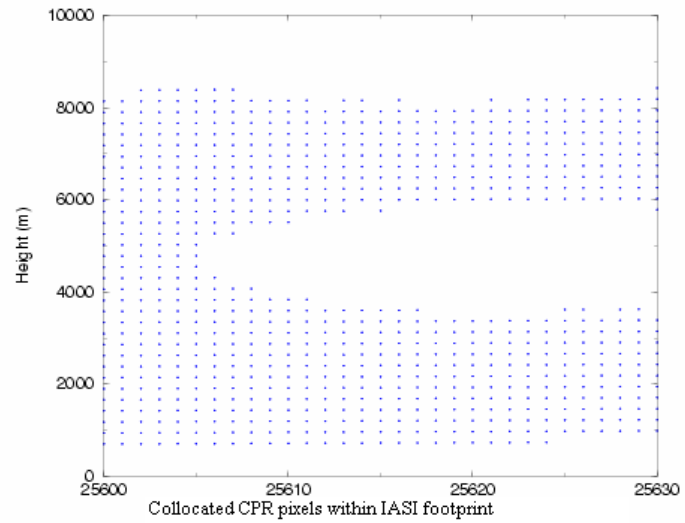
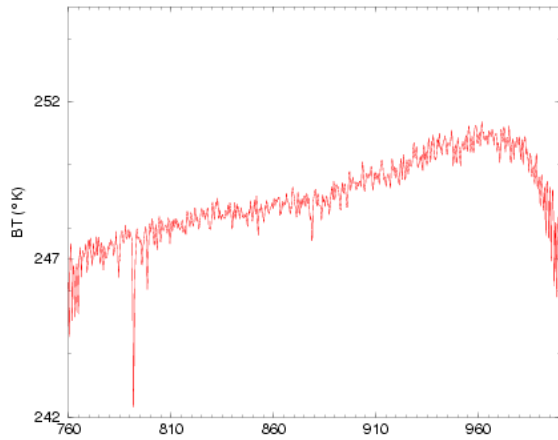
The Cloud Profiling Radar (CPR) is a 94-GHz nadir-looking radar which measures the power backscattered by clouds as a function of distance from the radar.

Nominal Frequency	94 GHz
Pulse Width	3.3 μ sec
PRF	4300 Hz
Minimum DetectableZ	-29 dBZ
Data Window	0-25 km
Antenna Size	1.85 m
Dynamic Range	70 dB
Integration Time	0.16 sec
Vertical Resolution	500 m
Cross-track Resolution	1.4 km
Along-track Resolution	1.7 km
Data Rate	15 kbps



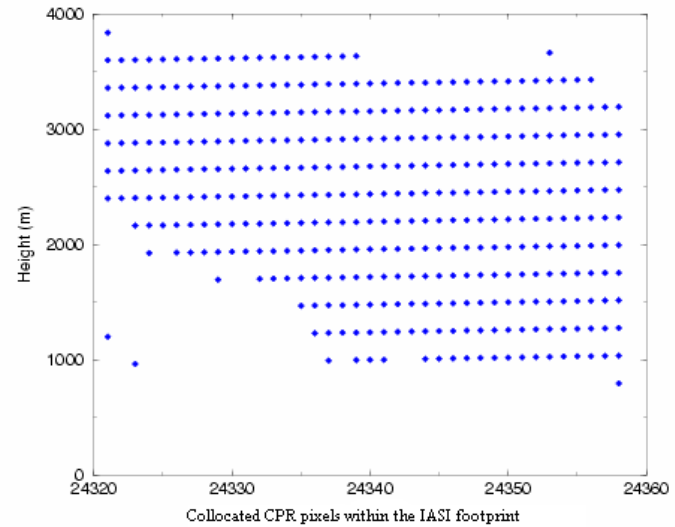
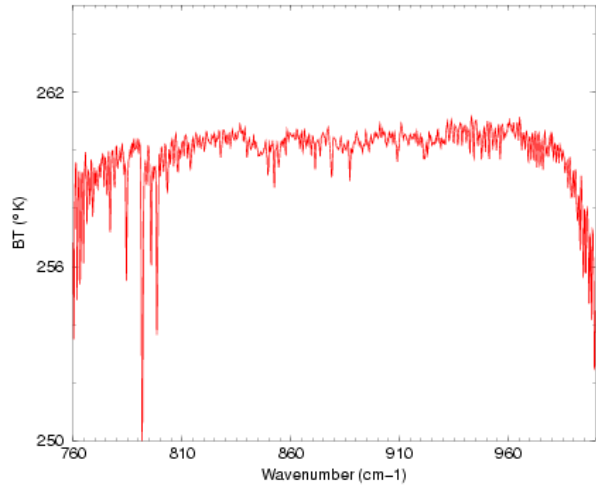
CPR System Characteristics

IASI spectra



Collocated CPR Cloud profiles

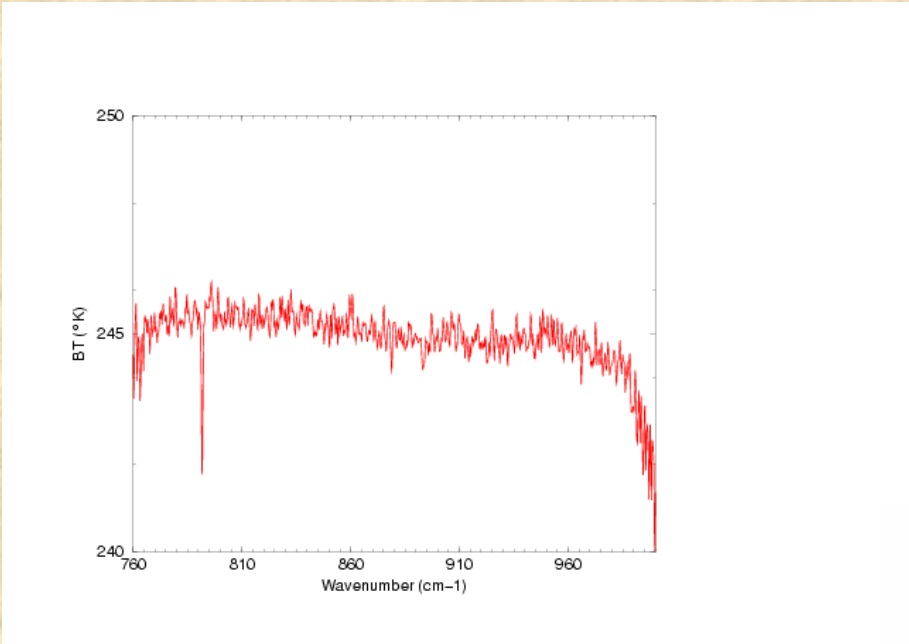
IASI spectrum



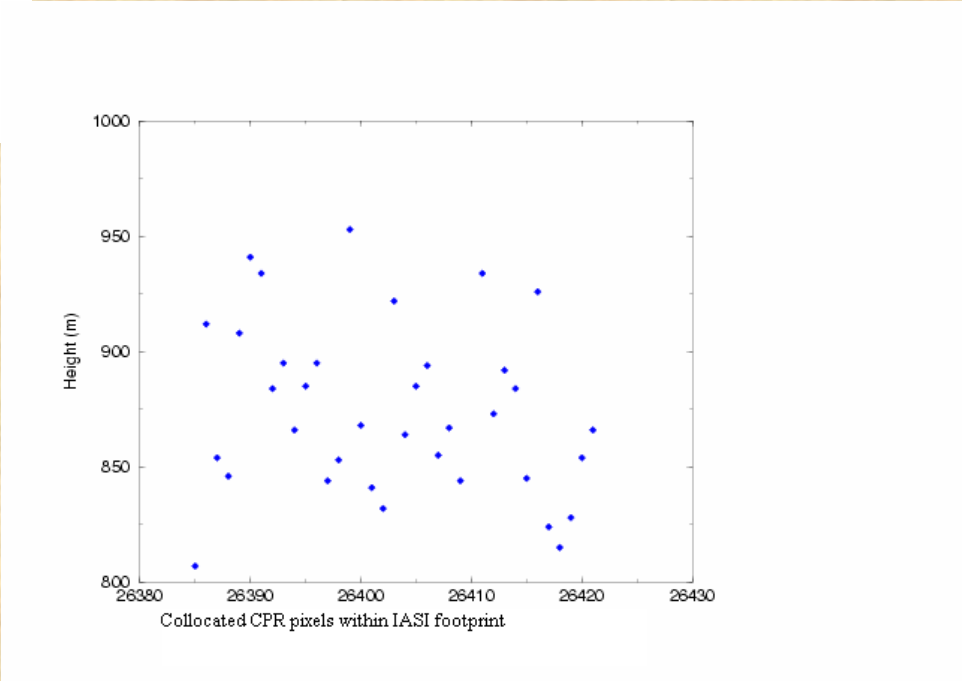
Collocated CPR Cloud profiles

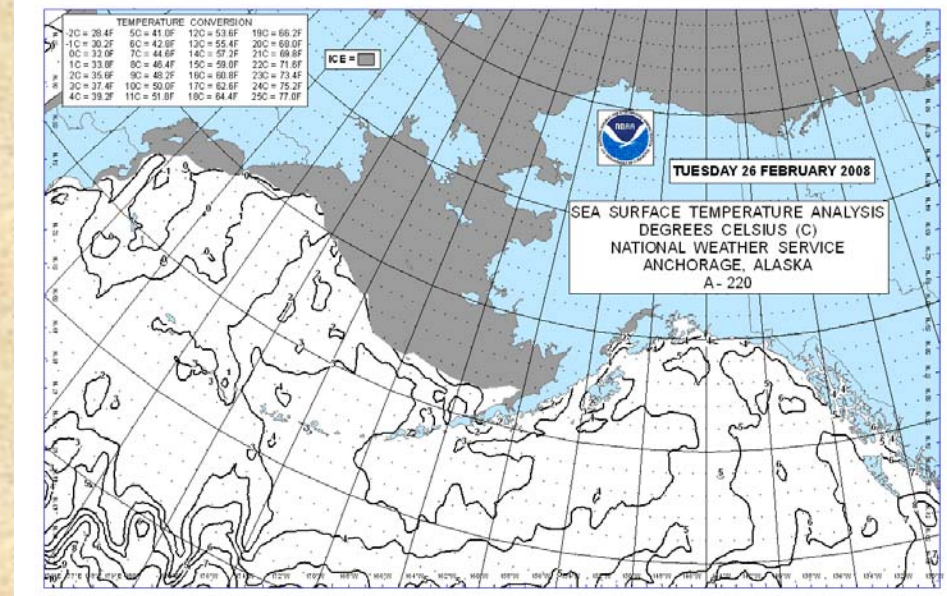
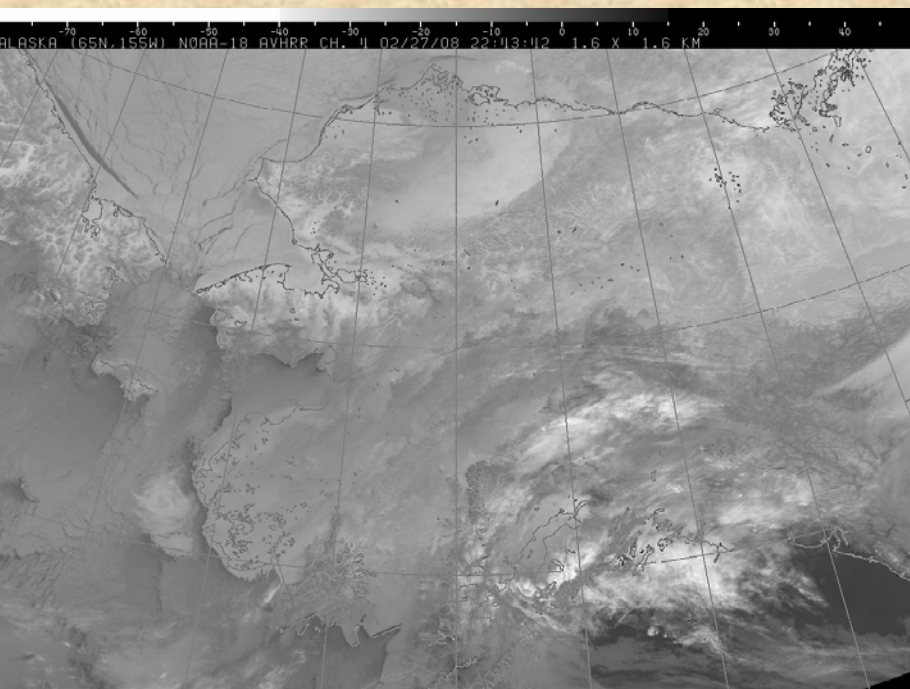


IASI spectrum

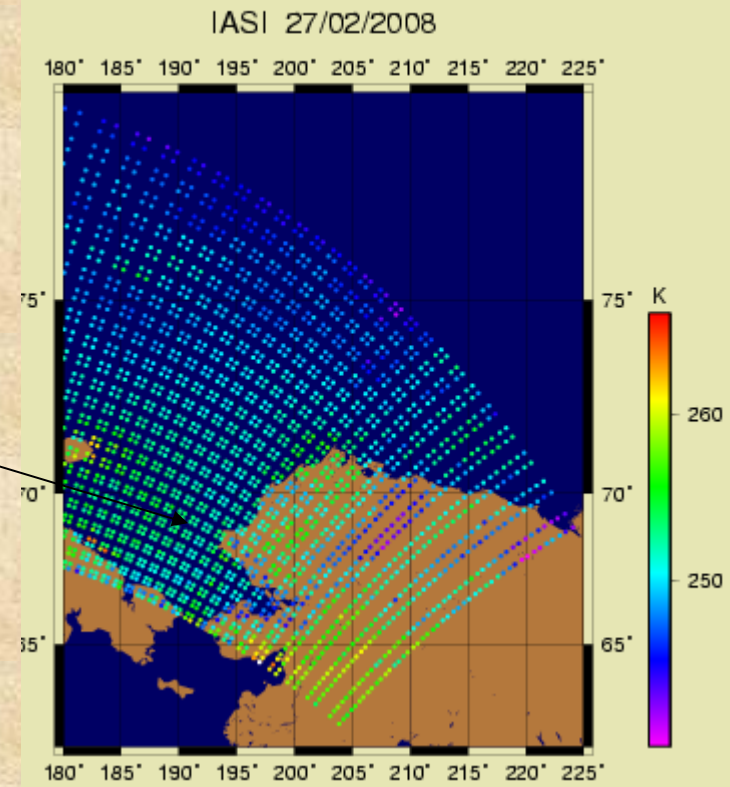
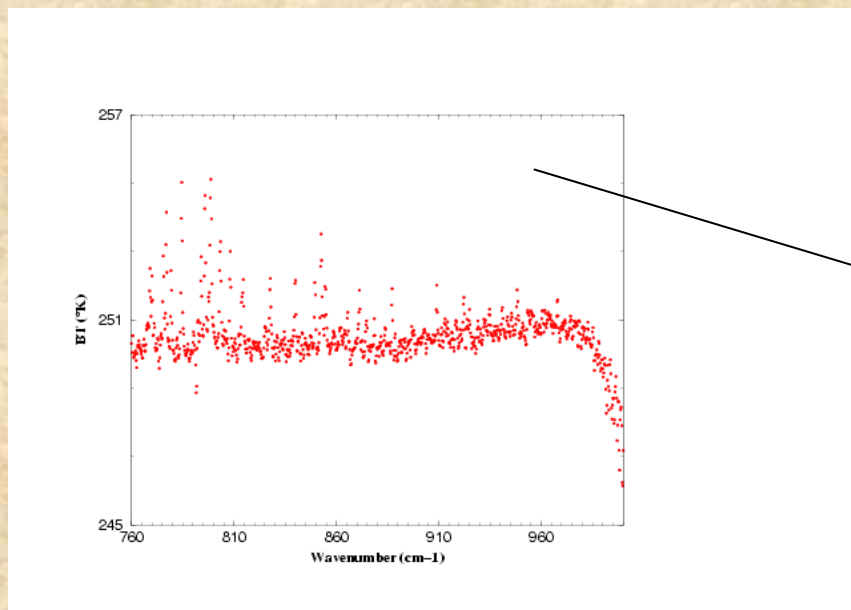


Collocated CPR Cloud profiles





Clear IASI FOV on ice surface



Cloud Test Results

Percentage of “clear– detected-clear”	96.9 %
Percentage of “cloud– detected-cloud”	98.7 %
Percentage of “clear– detected-cloud”	3.03 %
Percentage of “cloud– detected-clear”	3.70 %

1221 OVERCAST IASI FOVs.

**IASI FOV homogeneity => on the basis on collocated
AVHRR pixels within IASI footprint.**

Cloud Test Results

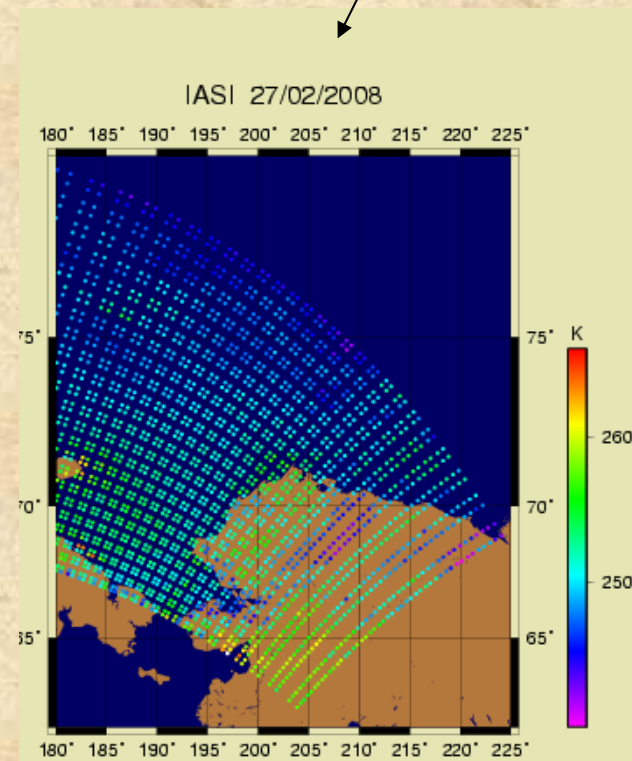
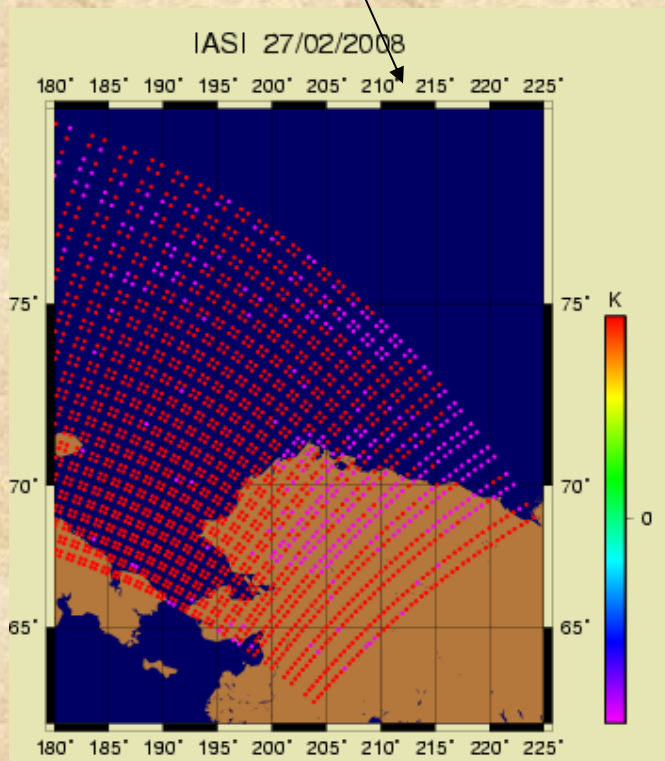
Percentage of “partially cloudy–detected–partially cloudy”	80.14 %
Percentage of “partially cloudy–detected–cloudy”	5.00 %
Percentage of “partially cloudy–detected–clear”	14.00 %

564 Partially cloudy IASI FOVS

IASI FOV homogeneity => on the basis on collocated AVHRR pixels within IASI footprint

IASI cloud mask

IASI (926. cm-1)



SUMMARY

The cloud detection scheme developed is able to detect cloud IASI FOVs for the different surfaces, also for ice surfaces.

For overcast IASI FOVs the scheme classifies as cloudy 98.7% of the pixels classified by CPR as cloudy and 96.9% as clear of the pixels classified by CPR as clear.

For partially cloudy IASI pixels the scheme detect 80.14% pixels correctly.