

An aerial photograph of a snow-covered landscape, likely a ski resort, showing numerous tracks and runs in the snow. The snow is bright white, and the tracks are visible in various directions. The background is a soft, hazy blue, suggesting a clear sky or a light mist.

## On the benefit of using spectral albedo and light penetration depth in detailed snowpack simulations

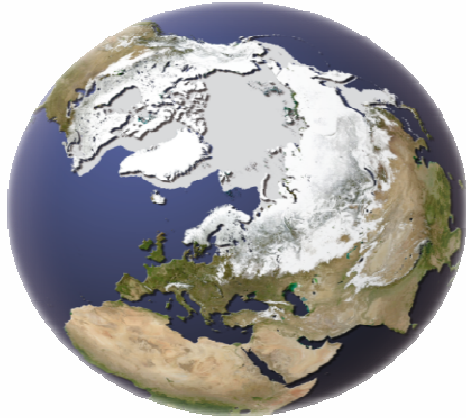
M. Dumont<sup>1</sup>, M. Lafaysse<sup>1</sup>, G. Picard<sup>2</sup>, L. Arnaud<sup>2</sup>, Q. Libois<sup>2,3</sup>  
and S. Morin<sup>1</sup>

1: Météo-France – CNRS, CNRM UMR 3589, Centre d'Études de la Neige, Grenoble, France

2: UGA / CNRS, Laboratoire de Glaciologie et Géophysique de l'Environnement (LGGE) UMR 5183, Grenoble, 38041, France

3: now at: ESCER Centre, Department of Earth and Atmospheric Sciences, Université du Québec à Montréal, 201 Av. du Président-Kennedy, Montréal, H3C3P8, Canada

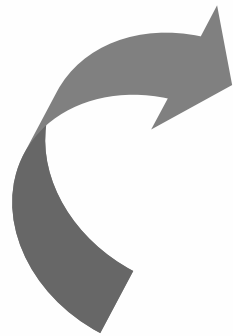
# White snow



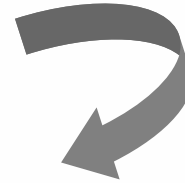
Snow is white which implies that it reflects much more solar light than other earth surfaces.

Impact on the Earth energy budget and on the global climate

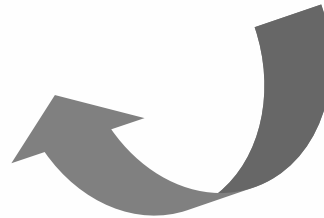
More snow



T decreases



Less energy absorbed



*[Imbrie & Imbrie, 1980]*

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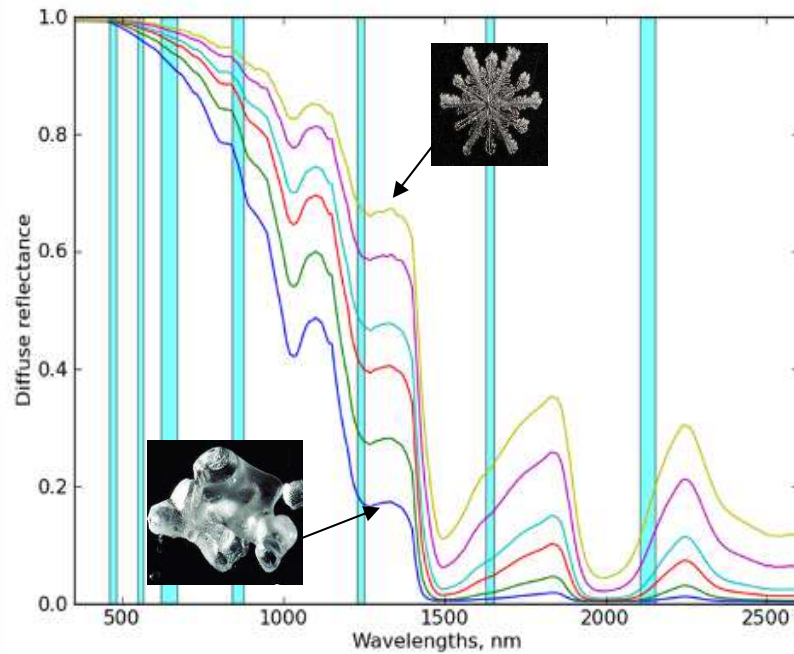
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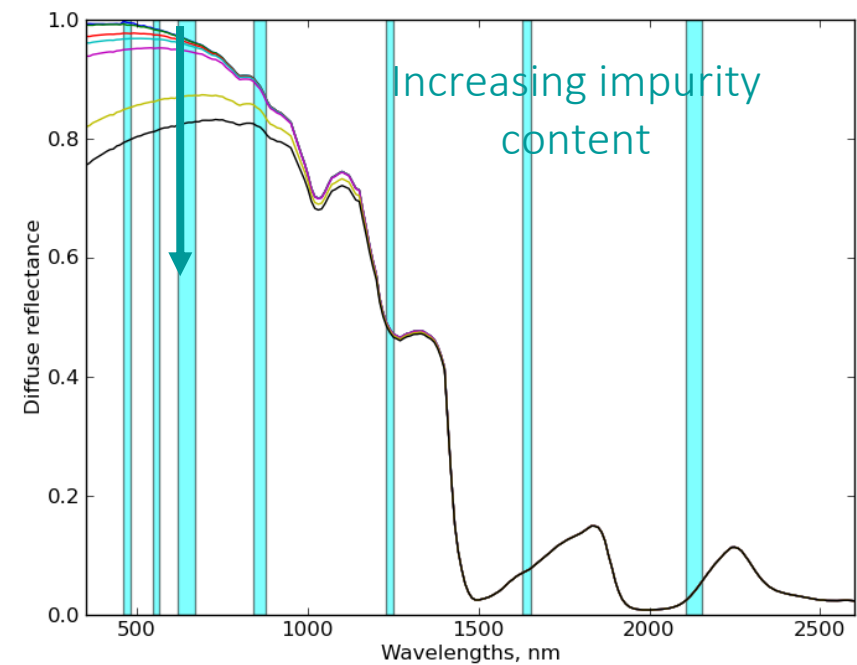
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# Snow optical properties : fifty shades of white

Snow microstructure



Snow impurity content



Snow optical properties are mainly determined by its microstructure and the light absorbing impurity content

This leads to several positive snow albedo feedbacks.

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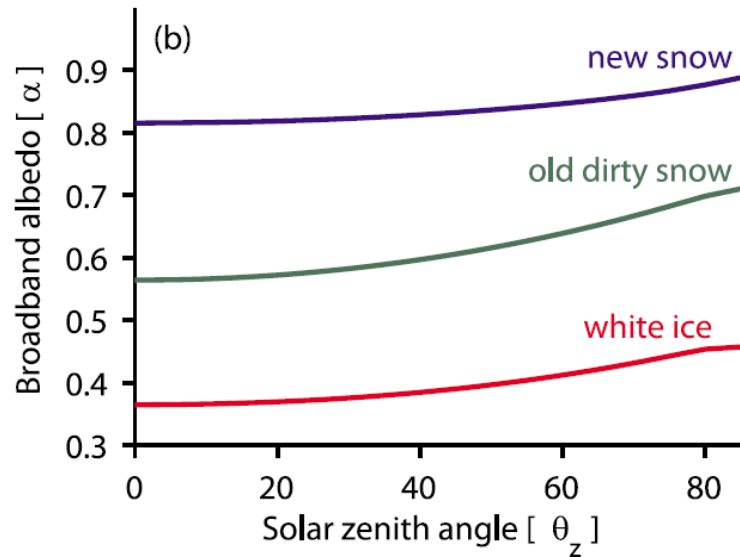
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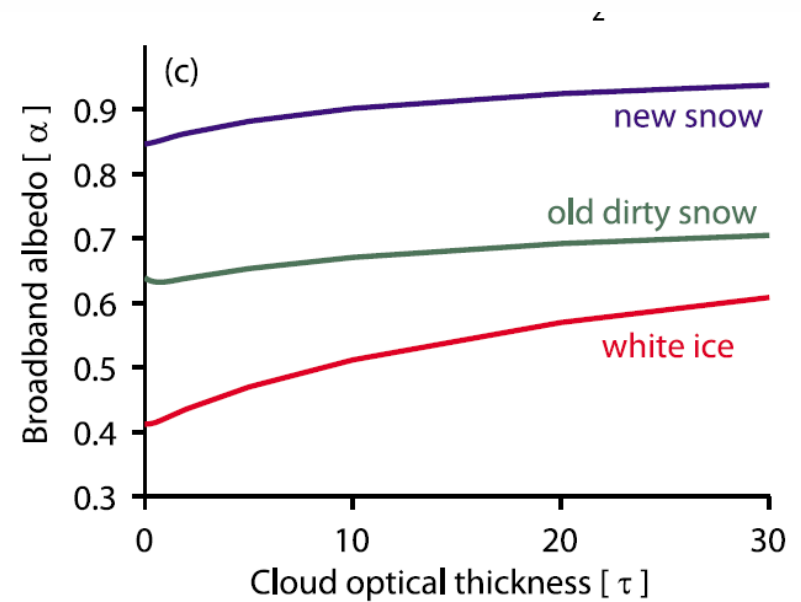
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# Snow optical properties and solar irradiance

Solar zenith angle



Cloud optical thickness



[Gardner and Sharp, 2010]

Snow albedo is also influenced by the angular and spectral characteristics of the solar incoming irradiance.

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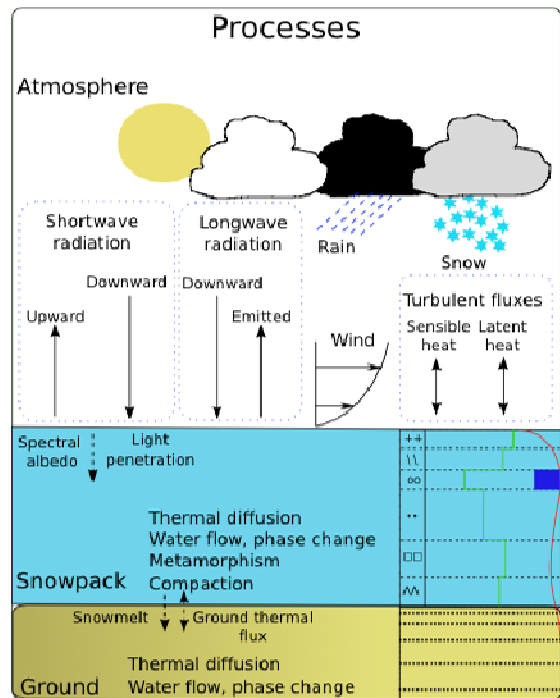


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# Detailed snowpack modelling



Detailed snowpack model Crocus [Brun et al., 1992 ; Vionnet et al., 2012] :

- Lagrangian representation of the snowpack
- Handles the major physical processes leading to snowpack evolution

Crocus is used for operational avalanches forecasting in Météo-France but also for a wide range of applications including hydrological and climatic studies.

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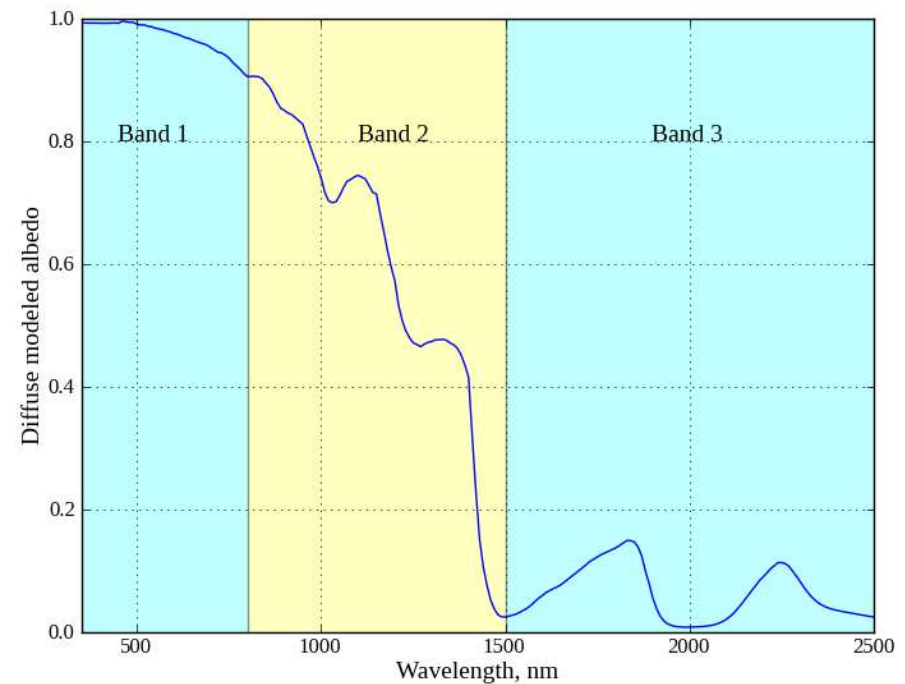
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# Crocus original radiative scheme

- 3 large spectral bands
- No dependence on solar zenith angle
- Account for the broadband diffuse to direct irradiance ratio only
- Parametrisations derived from *Wiscombe and Warren, 1980*
- Implicitly account for the impurity effect via an ageing coefficient
- Only one or two snow layers accounted for in the albedo calculation



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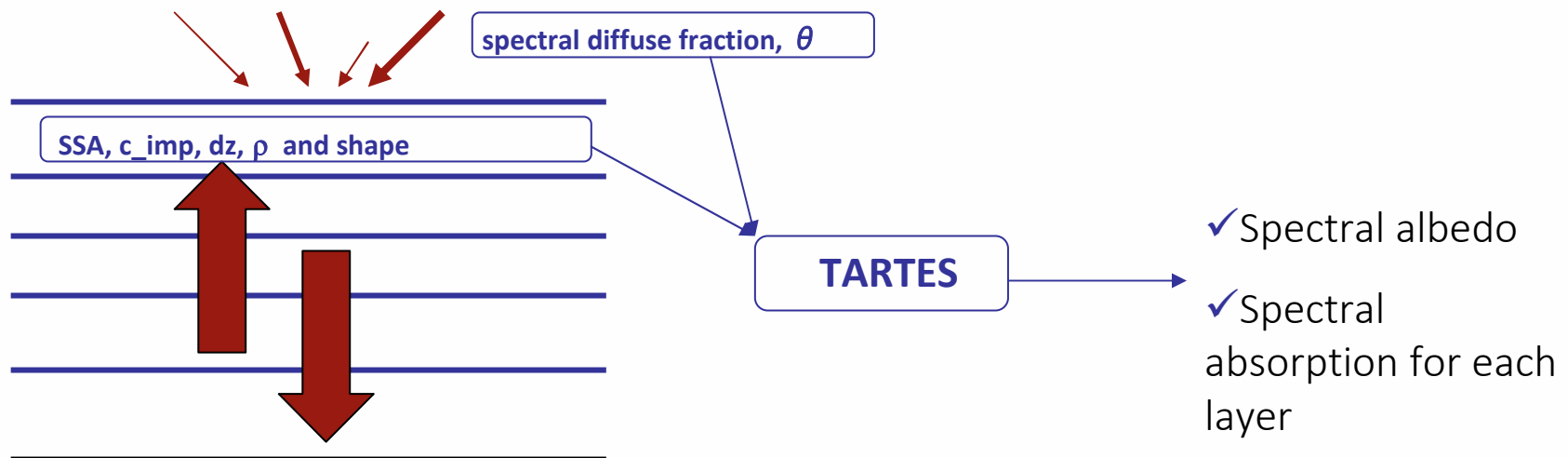
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# Two-stream Radiative Transfer in Snow, TARTES

- ✓ Two stream radiative transfer model based on *Kokhanovsky et Zege 2004* -> account for all the snow layers
- ✓ Explicit impurity content
- ✓ 10 nm spectral resolution
- ✓ Account for shape factor (snow grains differ from spheres)
- ✓ Account for spectral and angular characteristics of the incoming radiation



[Libois et al., 2013 ; 2014]

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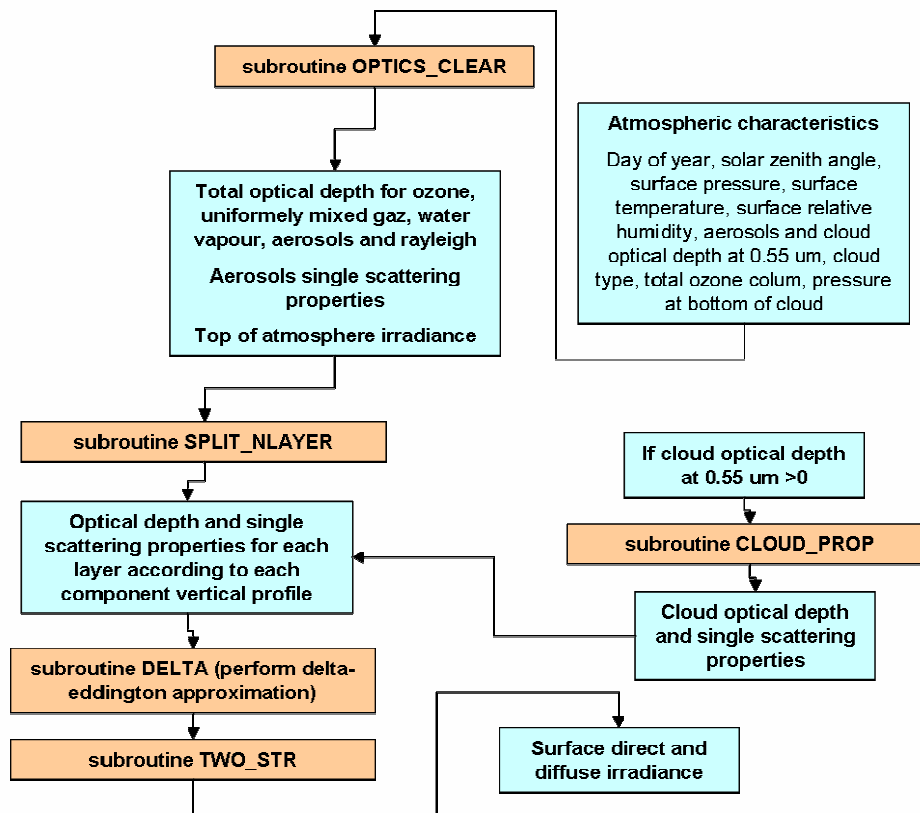
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# Spectral solar irradiance : ATMOTARTES

Two-stream model derived from SBDART [Richiazzi et al., 1998]



- Account for ozone, aerosols, mixed gaz, water vapour and clouds
- Evaluated vs SDBART ( $r^2 > 0.988$ )
- Used to compute the spectral diffuse to total ratio of the solar irradiance from either diffuse and direct or total broadband irradiance

[Dumont et al., 2015 ; AGU]

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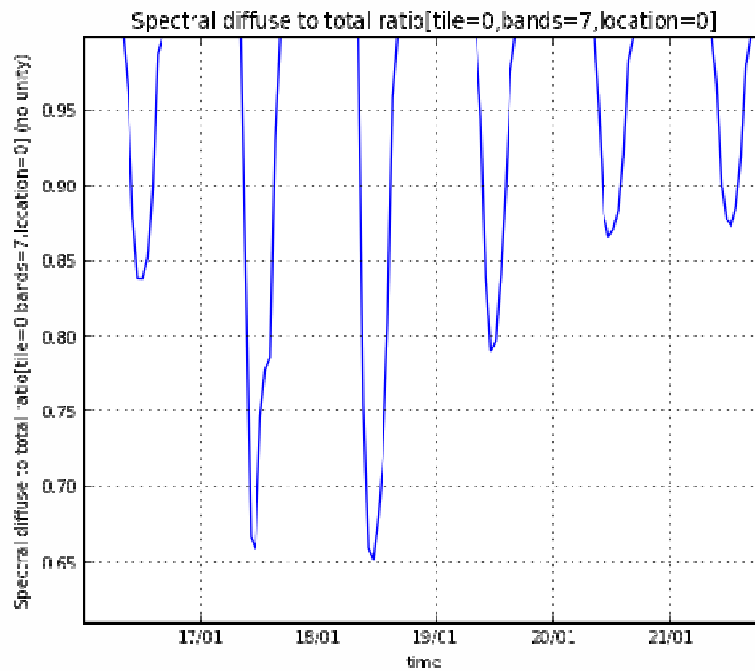
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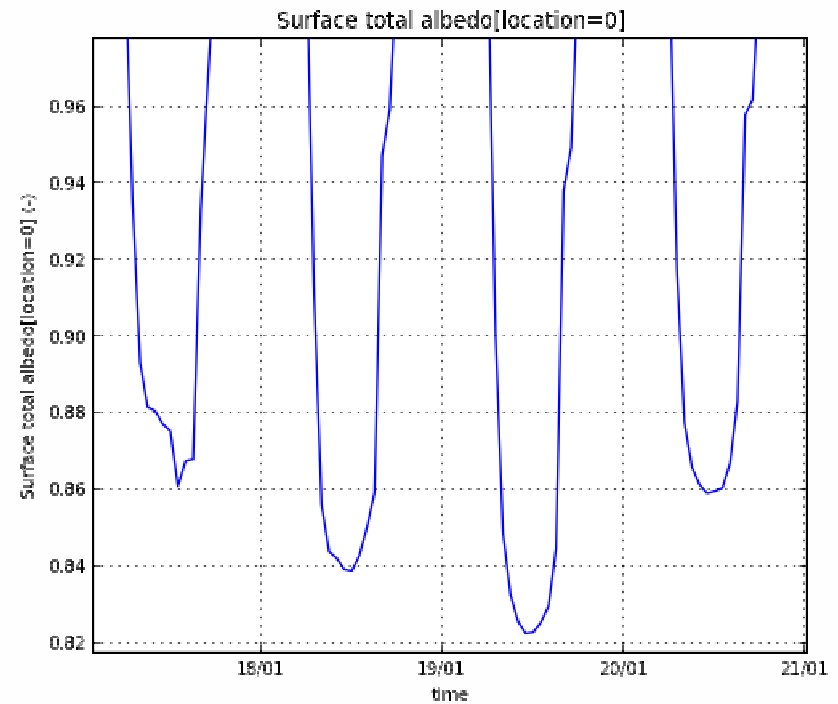
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# Crocus with the full radiative transfer schemes



At 420 nm



Diurnal cycle of broadband albedo

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Lab. OSUG 2020

RN VOR



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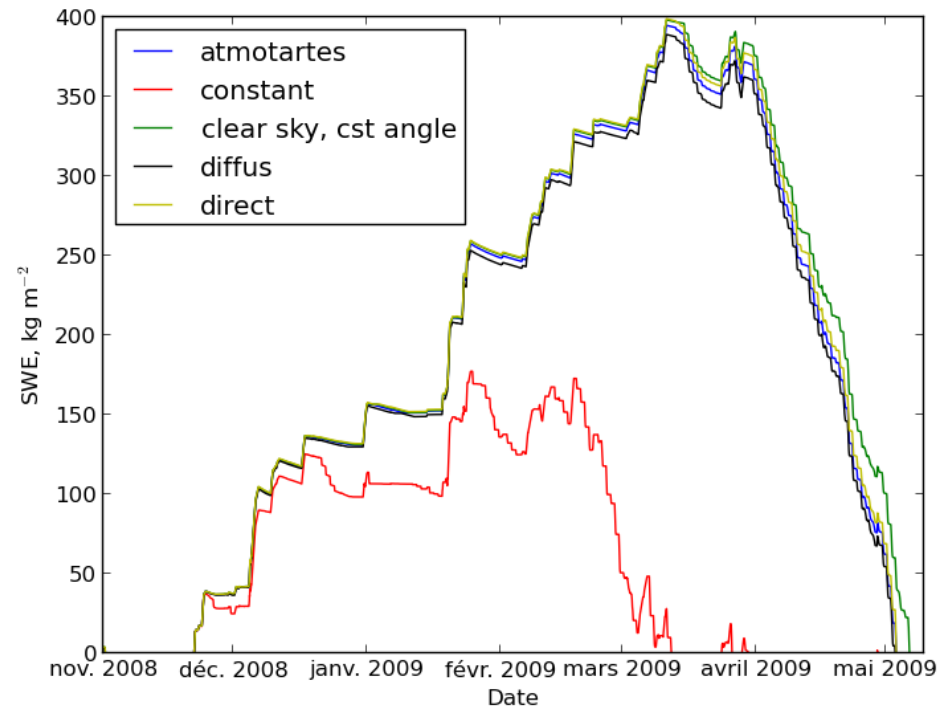


Pôle Grenoblois d'étude et de recherche pour la prévention des Risques Naturels



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# Impact on snow simulations



Limited impact of the direct/diffuse repartition as long as the spectral repartition is known, but depends on SZA.

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# Impurity content as a prognostic variable

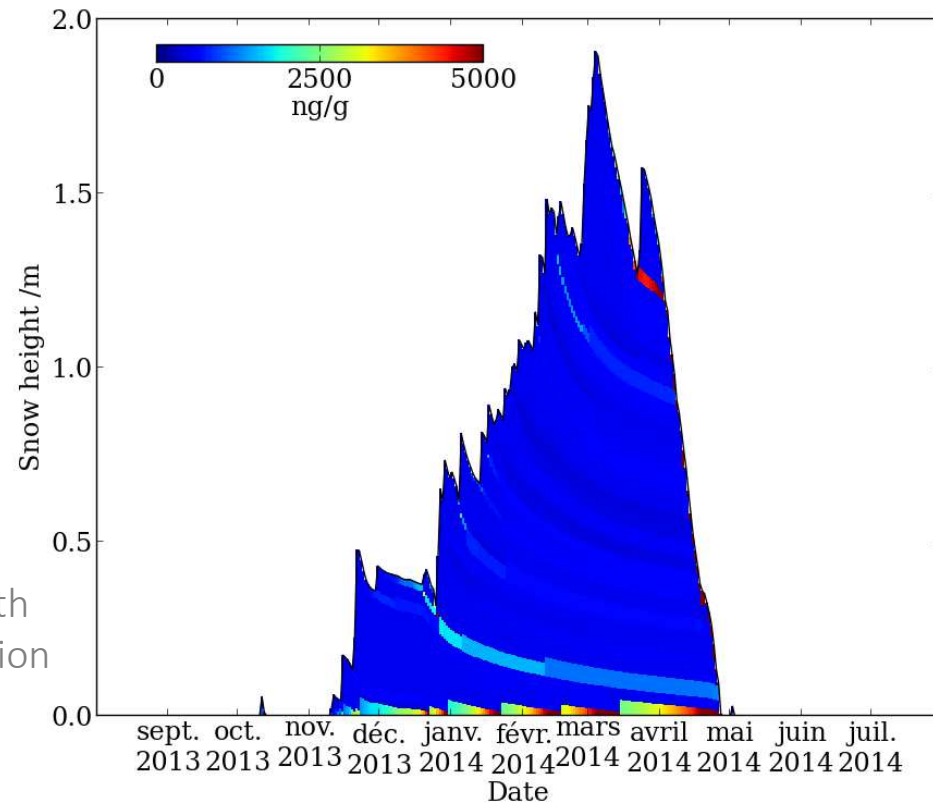
- As for now, only one type « BC equivalent », dry and wet deposition fluxes are set to constant values
- Total conservation of the impurity mass

$$c(t + \Delta t) = c(t) + \Delta t \tau_{\text{dry}} e^{-D/h_{\text{ref}}}$$

Dry deposition  
flux

e-folding depth  
for dry deposition

Depth



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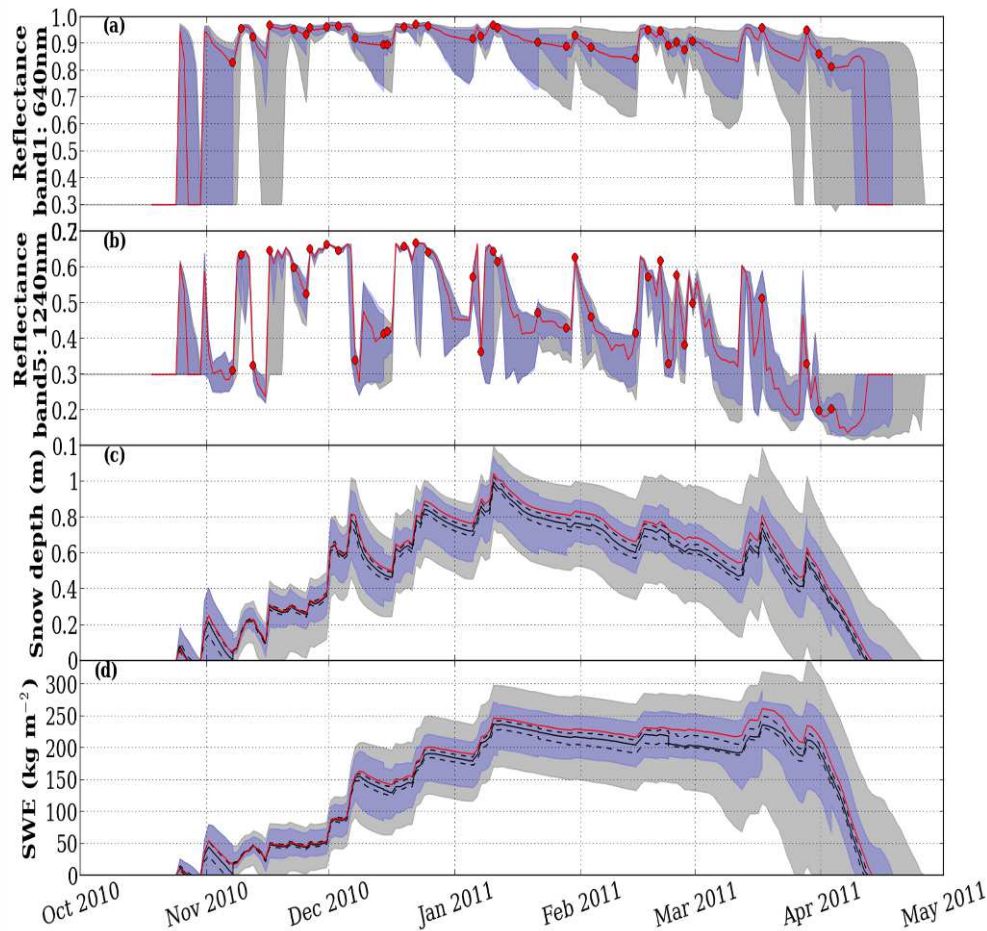


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# Application 1 : Assimilation



This new scheme allows the direct assimilation of satellite surface reflectance in the snow model.

ATMOTARTES + TARTES being the observation operator.

This is detailed in L. Charrois talk on Wednesday morning.

*[Charrois et al. ,2015]*

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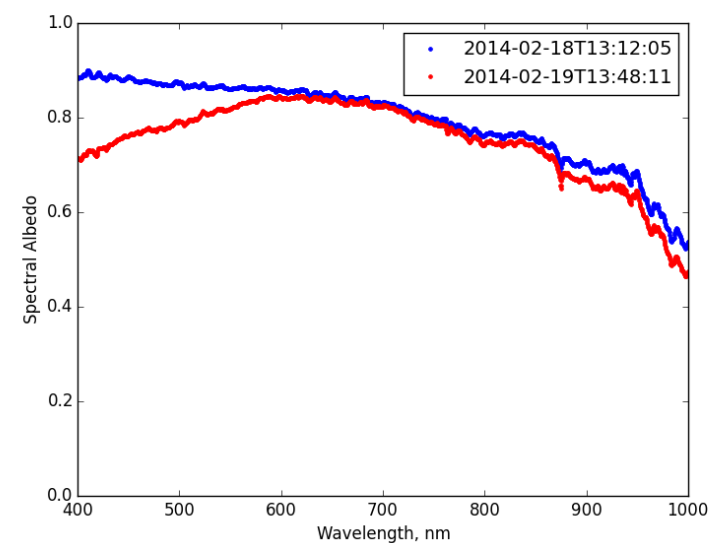


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## Application 2 : Spectral albedo measurements at Col de Porte, 1325 m a.s.l



- Operations : December 2013-May 2014
- Spectra collected every 12 min, spectral resolution  $>1$  nm, spectral range 300-1100 nm
- Raw spectra are corrected for : (i) dark current, (ii) calibration and (iii) angular response of the light collector using modelled diffuse to total irradiance ratio [Picard et al, 2015]

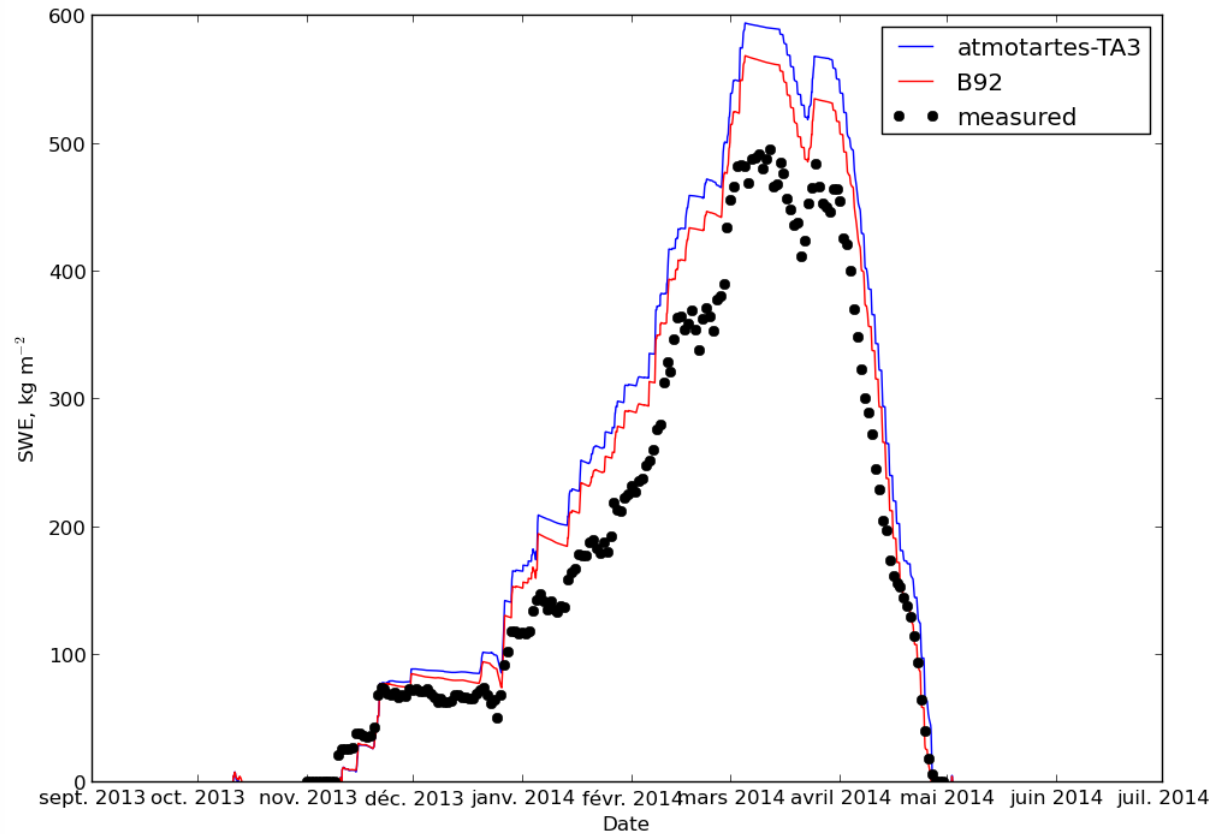


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# Crocus simulations



New doesn't necessarily mean better .... Probably less error compensation. Need to optimize wet and dry deposition fluxes of impurity and scavenging rates with melt water

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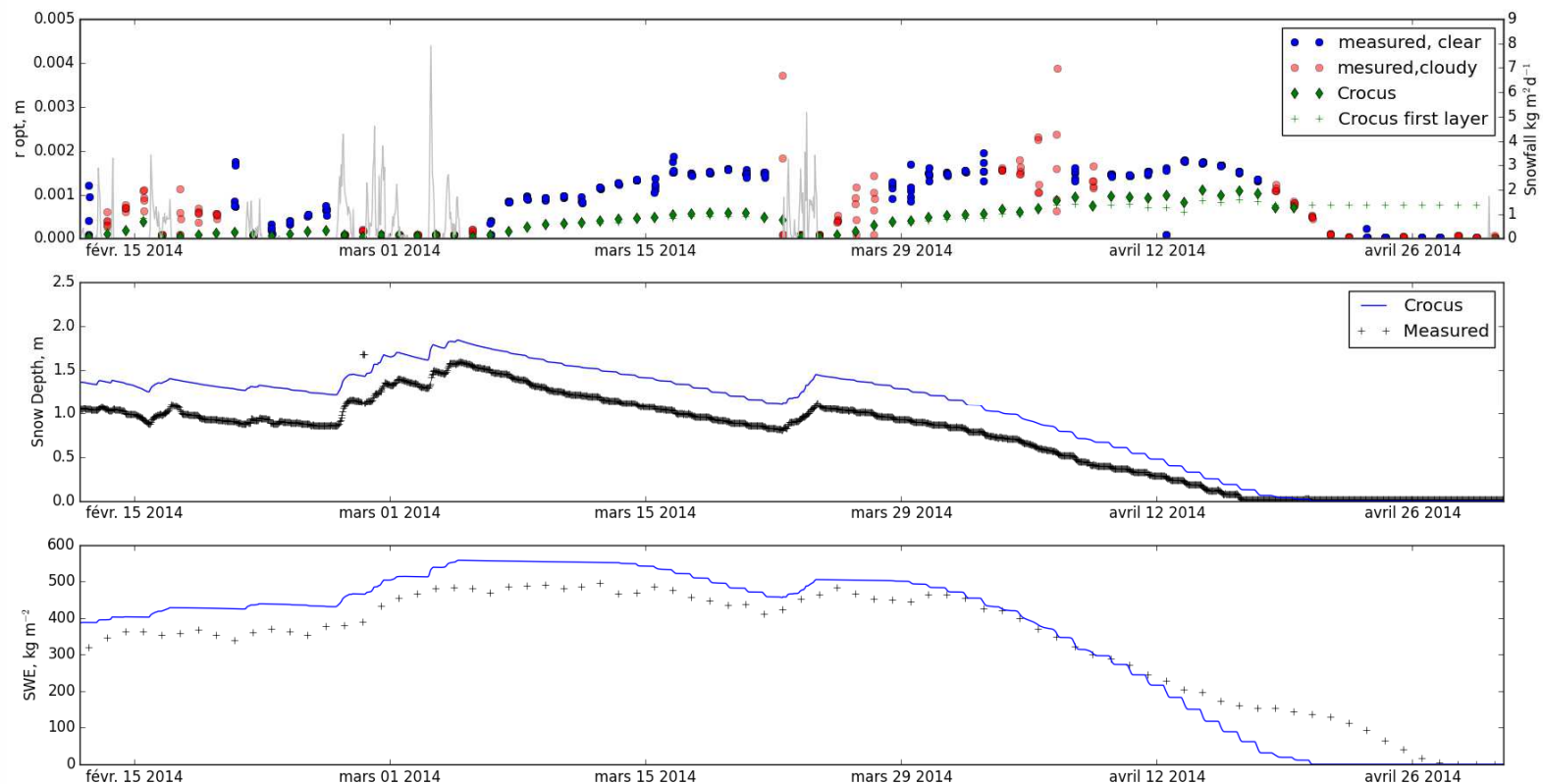


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# Surface snow optical radius



$R_{opt}$  is retrieved from measured and modelled spectral albedo using  
*Kokhanovsky and Zege, 2004.*

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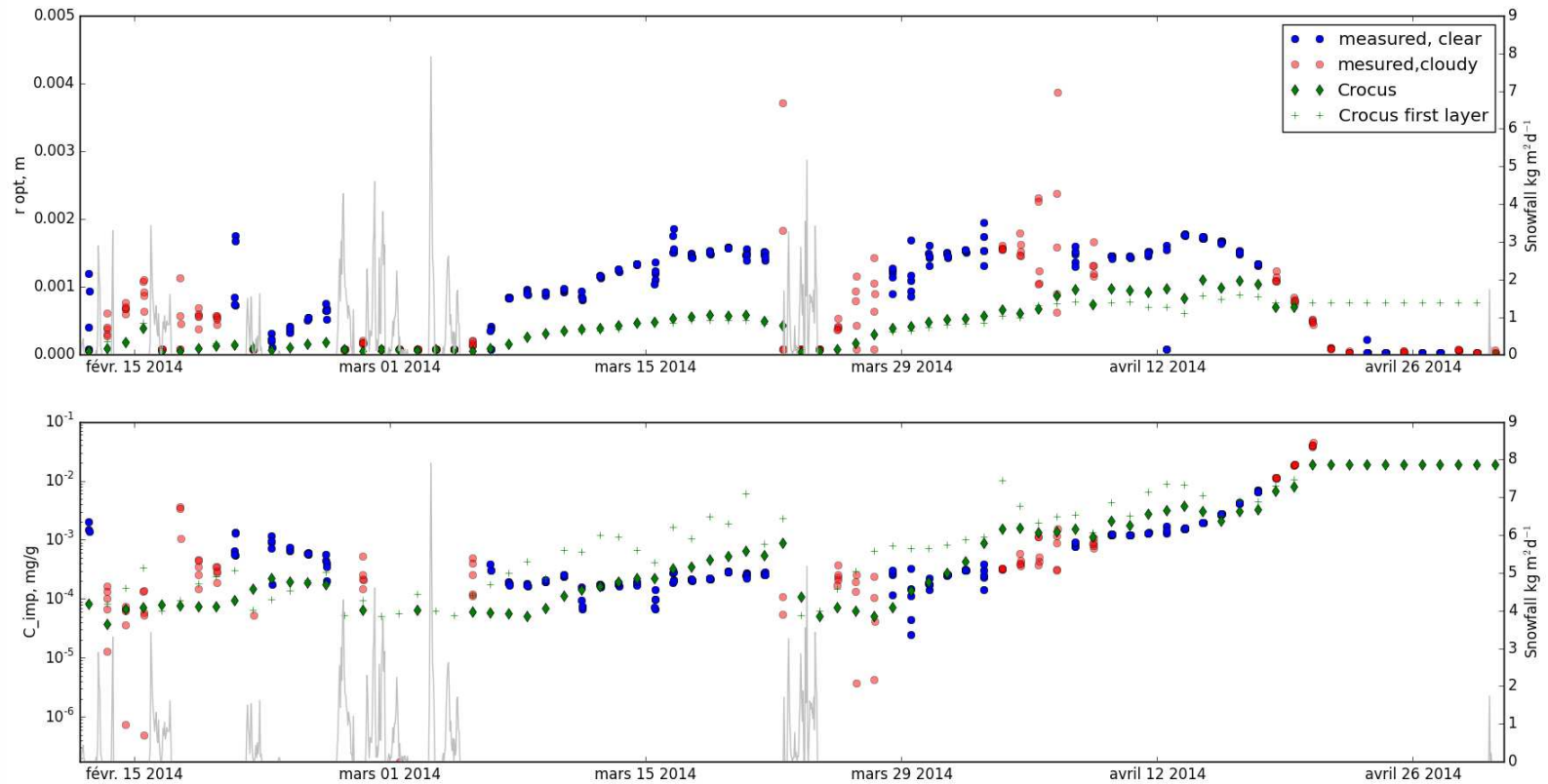


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# Surface snow impurity content



Constant wet and dry deposition fluxes allows a first reasonable representation of the impurity content.

Remarkable  $r_{opt}$  decrease under large impurity loads ... A negative feedback ? Still to be investigated

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## Conclusion and future works

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- Crocus is now the only detailed snow model handling a spectral scheme for radiative transfer in the snow and in the atmosphere together with a prognostic impurity content
- It offers unique capability for optical data assimilation and to account for all the major snow albedo feedbacks in snowpack simulations e.g. needed for accurate climate projections of the Greenland ice sheet (*Tedesco et al., 2016*)
- We still need :
  - force the wet/dry deposition fluxes with an atmospheric chemistry-transport model (e.g. ALADIN-Climat)
  - implement different types of impurity (dust, bc, algae ...)
  - investigate and adjust the evolution laws for impurity (scavenging rates ...)

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Questions ?